### TIME DOMAIN ELECTROMAGNETIC MEASUREMENTS EAST-CENTRAL FLORIDA

Performed For:

Governing Board of the St. Johns River Water Management District P.O. Box 1429 Palatka, Florida 32178-1429

(Contract No. 91G143)

Prepared By:

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(BGI Project #91034)

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This is to certify that I, Jim Hild, have reviewed the figures, tables, and text of the following report, and have retained one copy for my files.

ile an By:

James Hild Professional Geologist State of Florida Registration #PG-0001282

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### EXECUTIVE SUMMARY

Time domain electromagnetic (TDEM) measurements were made in the St. Johns River Water Management District (SJRWMD) in east-central Florida during August 1991. TDEM is a geophysical method that measures from the surface the resistivity layering (geoelectric section) of the subsurface. The objective of the TDEM survey was to infer from the geoelectric sections measured information about water quality in the Upper Floridan Aquifer, such as the first depth of occurrence of ground water with a chloride concentration greater than 1,000 mg/l, and the depth of the 250 mg/l isochlor.

The objective of determining the first depth of occurrence of ground water with a chloride concentration greater than 1,000 mg/l was most readily accomplished. At the 14 sites surveyed that depth varied between the 14 sites from about 180 ft below surface to in excess of 2,000 ft below surface. TDEM measurements are an indirect method for determining water quality, and validity of interpretations are best tested against salinity measurements in wells. At one site (Astronaut High School) a well penetrated the interface with highly saline water, and excellent agreement between depth to ground water of high salinity inferred from TDEM measurements and observed in the well, was obtained.

Determining water quality above the interface with ground water of high salinity, in particular determining the depth of the 250 mg/l isochlor is a more difficult objective to reliably accomplish. This objective is more sensitive to assumptions that necessarily need to be made about chemical composition of ground water and porosity. Also, in present day methods of analysis, surface geophysics resolve distinct boundaries. In real aquifers ground water quality generally changes gradually. Nevertheless, at the majority of sites meaningful information about water quality in the Upper Floridan aquifer was obtained that correspond with available regional data.

The impetus for using surface geophysics for obtaining water quality is its low cost per station compared to drilling, so that higher station density can be afforded. Surface TDEM measurements can, however, not deliver information about water quality to the same level of detail (e.g., chemical composition) or accuracy as can be derived from wells.

### 1.0 INTRODUCTION

I

This report covers the data acquisition, processing, interpretation, and results of a time domain electromagnetic (TDEM) measurements performed at 14 sites in the St. Johns River Water Management District (SJRWMD) in eastcentral Florida. TDEM is a geophysical method that determines from the surface the geoelectric section (resistivity layering) in the subsurface. From the geoelectric section in-turn information about geology and water quality can be inferred, because the electrical resistivity of the Earth depends on lithology, porosity, and concentration of dissolved solids in ground water.

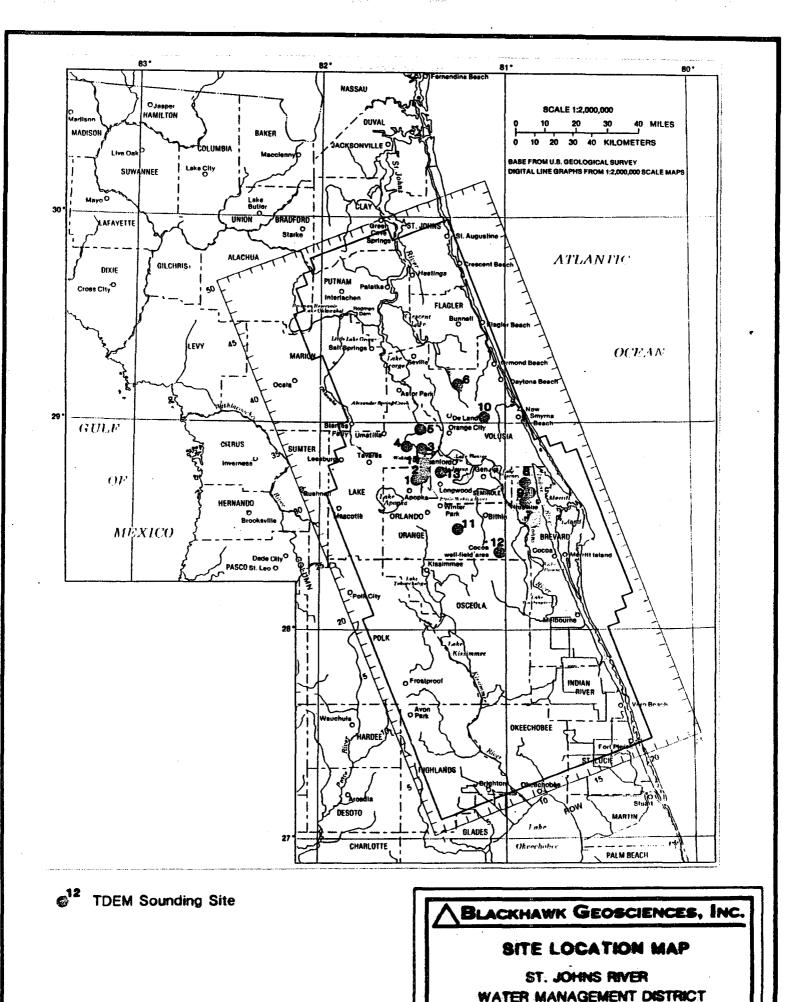
The objectives of the measurements at the 14 sites were to determine

- the interface between fresh water and highly saline water, defined here as ground water with a chloride concentration greater than 1000 mg/l
- the water quality in the Upper Floridan aquifer above the interface with highly saline water. Estimating the depth to the 250 mg/l isochlor particularly was an important objective.

The selection of the locations of the 14 sites for TDEM measurements were made by personnel of SJRWMD, and the overview map of Figure 1-1 shows their location.

The hydrogeologic setting of the 14 sites is expected to be similar. All sites are over the Floridan aquifer system, a system defined (Miller, 1990)on the basis of permeability. In general, the system is at least 10 times more permeable than its bounding upper and lower confining units. The aquifer system is thick and widespread, and the rocks within it generally vary in permeability. It is divided into an Upper Floridan aquifer and a Lower Floridan aquifer, separated by a less permeable unit in most places, and bounded above and below by confining units that may be less permeable.

Figure 1-2 is a stratigraphic column typical of the area, and it also summarizes characteristics of the various units. The geologic formations that comprise the Floridan aquifer system are all carbonate rocks, and most of them are of Eocene age. In places, however, rocks as young as Miocene age are included in the system. The Hawthorn Group and younger primarily clastic material overlies the carbonate rocks at most sites in the SJRWMD. The thicknesses and properties of these units are highly variable. Their composition may vary from permeable sands to impermeable clayey quartz.



PROJECT NO: 91084

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Figure 1-1

# TYPICAL HYDROSTRATIGRAPHIC COLUMN IN WEKIVA RIVER AREA

AGE	FORM	ATION	AQUIFER	THICKNESS	DESCRIPTION
LATE AND POST MIOCENE		SURFICIAL MATERIAL	SURFICIAL AQUIFER	34 ft. to 139 ft.	Sand, Clay, and Coquina
MIOCENE		HAWTHORN	INTERMEDIATE AQUIFER AND CONFINING UNIT	25 ft. to 150 ft.	Sand, Clay, Limestone, and Bolostone, mostly phosphatic
Σ.				Š ft. to 32 ft.	Basal Hawthorn: Dolostone, sandy, phosphatic, hard
		OCALA LIMESTONE		20 ft. to 1 12 ft. 9 ft.	Limestone, relatively pure Coquina, bio- and foraminiferat- Lower Ocata;
			FLORIDAN AQUIFER	to 57.ft.	Limestone, dolomitic, coquincid
EOCENE		AVON PARK LIMESTONE		150 ft. to 330 ft.	Limestone and Dolostone with Peat (disseminated and as beds) or Clay beds
					Avon Park: Dolostone, very hard, tow perosity zone
L				TYPICA	KHAWK GEOSCIENCES, IN L HYDROSTRATIGRAPHK N IN WEKIVA RIVER ARE
				St. Johns PROJECT NO	River Water Management Distric 9: 91094 Figure 1

### 2.0 TECHNICAL APPROACH

The geophysical approach taken to achieve the objectives, stated in Section 1.0, was to determine the geoelectric section of the subsurface, and to infer from that section information about water quality. The resistivity of a water bearing rock is mainly a function of lithology, dissolved solids in ground water, and porosity. Most rock forming minerals are essentially insulators and nearly all electrical current is carried either by free ions in pore water or by exchangeable ions associated with clay particles. To separate the causes of vertical and lateral variation in a geoelectric section requires careful correlation with lithology, and often assumptions about the dominant cause of resistivity variation locally must be made. For the 14 sites in the SJRWMD variation in lithology is mainly expected in the Hawthorn Group and younger surficial units. The composition of these formations can vary from coarse-grained sands and gravels to clays. Thus, in the Hawthorn Group and younger sediments three factors potentially can influence resistivity, - lithology (clay content), porosity and water quality. Without other independent information the causes of lateral and vertical resistivity variation cannot be separated, and no attempt has been made to infer information about water quality from resistivity measurements for the formations above the Floridan aquifer. However, at some sites, one average value of resistivity is obtained for the Hawthorn Group which is typically 100 ft to 200 ft thick, and several hundred feet of the Upper Floridan. In those situations the resistivity obtained probably closely correspond to the real resistivity of the carbonate rocks of the Upper Floridan aquifer.

On the other hand, the lithology of the carbonate rocks comprising the Floridan aquifer system are expected to be uniform. The resistivity of the rocks of the Floridan aquifer will be mainly determined by porosity and dissolved solids concentrations of the pore fluids. Archie's Law is used to express the relationship between formation resistivity, Ro; fluid resistivity, Rw; and porosity,  $\phi$ :

 $F = Ro/Rw = a \phi^{-m}$ 

where F = formation factor and a,m are empirically derived constants dependent on lithology and pore type distribution. Kwader (1986) found a value of m = 1.6 and a = 1 to best fit his many observations from wells completed in the Upper Floridan aquifer in Seminole County.

Fluid resistivity is a function of concentration of dissolved solids and ionic composition. The most common cations in water in the Upper Floridan aquifer are calcium, magnesium and sodium; the most common anions are bicarbonate, chloride and sulfate. Water quality is often expressed in terms of equivalent chloride concentrations. Kwader (1968) established on the basis of many measurements on water samples throughout Seminole County the relation between chloride concentration and fluid resistivity, Rw, given by

CL = 3500/Rw - 153

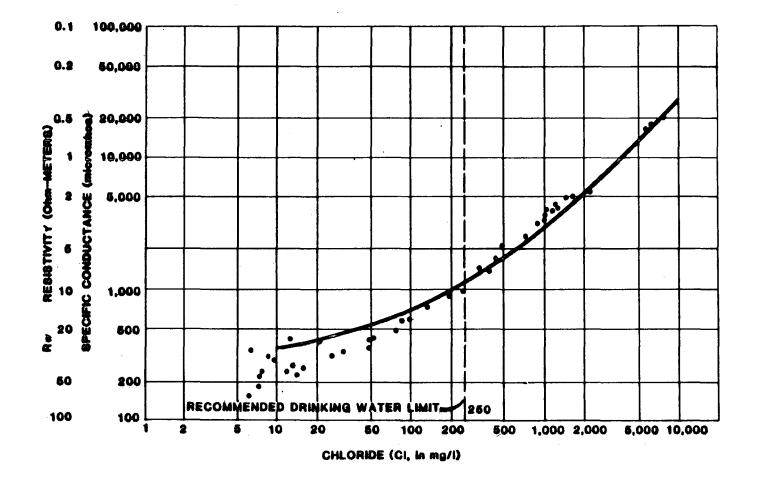
where CL is chloride concentration in mg/l, and Rw is fluid resistivity in ohm-meter.

A graphic presentation of this relation is given in Figure 2-1, and it also shows the data points from which relation [2] was derived. The maximum chloride concentrations for which data points were available to Kwader (1986) was about 10,000 mg/l, and the relation is untested at higher chloride concentrations.

By combining equation [1] and [2] chloride concentration can be related to formation resistivity as a function of porosity, and this relation is displayed in Figure 2-2. Thus, for the Upper Floridan aquifer with an average porosity of 25%, chloride concentrations less than 250 mg/l are expected when its formation resistivity is greater than 80 ohm-m. Chloride concentrations

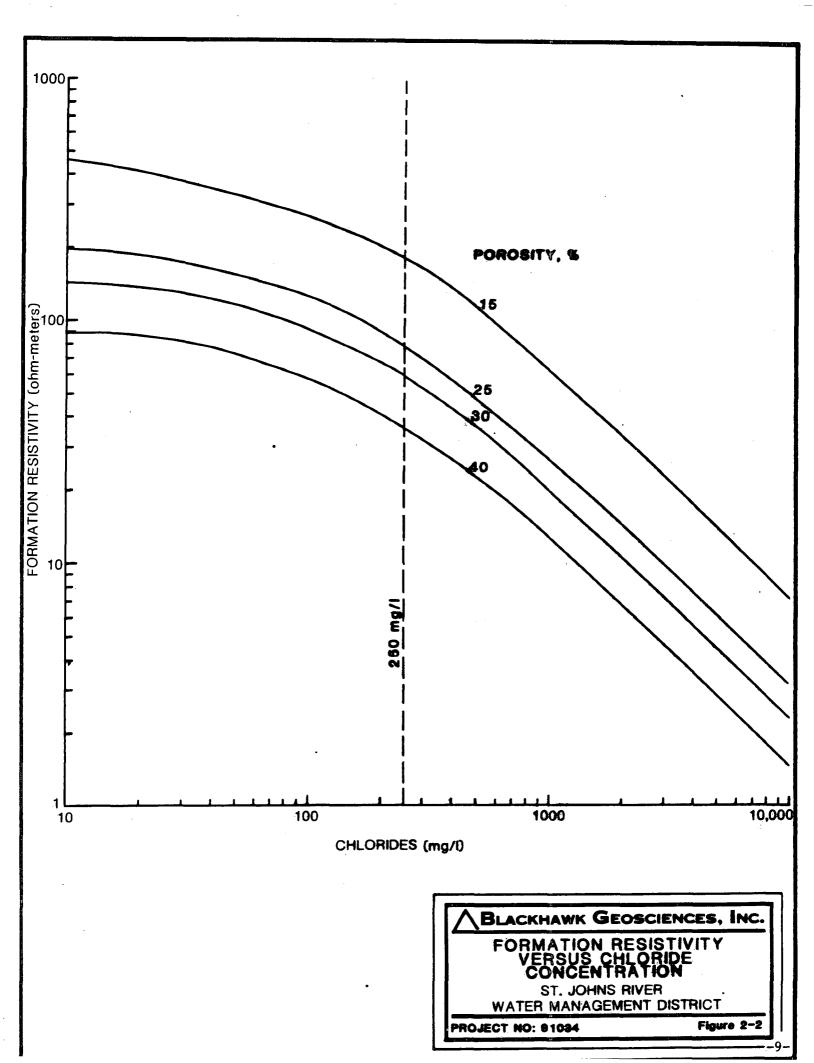
[1]

[2]



(After: Kwader, 1986)

BLACKHAWK GEOSCIENCES, INC. FLUID RESISTIVITY (Rw) VERSUS CHLORIDE CONCENTRATION St. Johns River Water Management District Figure 2-1 PROJECT NO: 91034 φ



greater than 1,000 mg/l would be indicated by formation resistivity values less than about 28 ohm-m, and greater than 10,000 mg/l at formation resistivities less than about 3.2 ohm-m.

It is evident from the above discussion that to derive chloride concentration from a measured value of formation resistivity certain assumptions must be made. The assumptions consistently made for all 14 sites are:

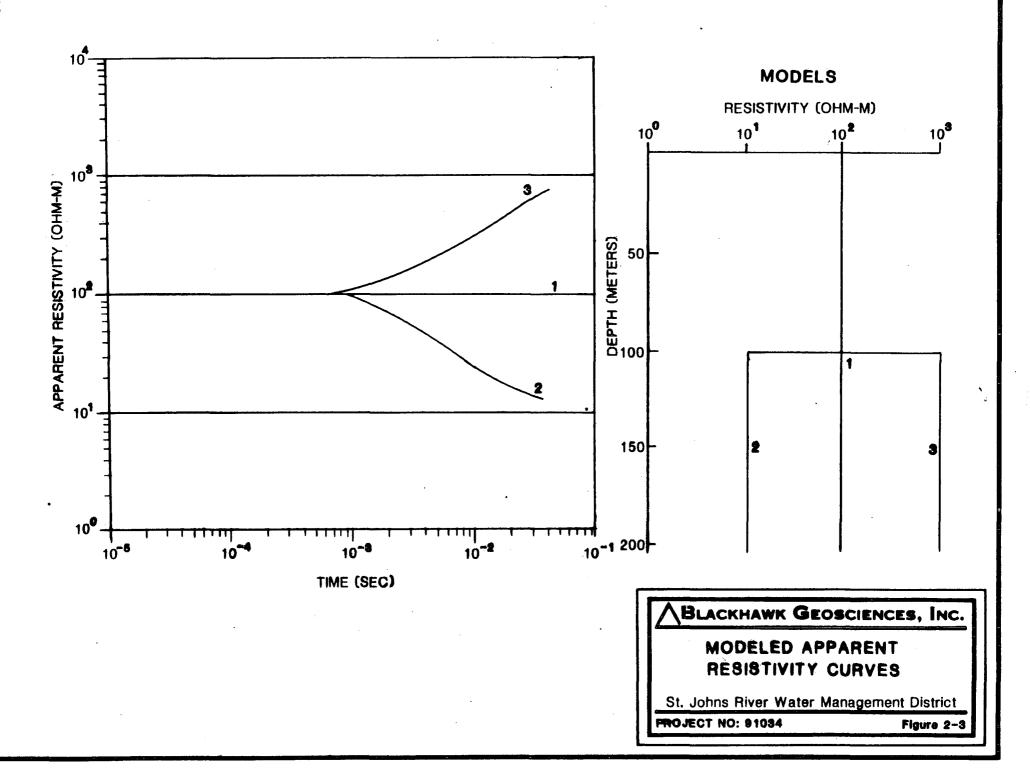
- a) The relation (Fig. 2-1) between fluid resistivity and chloride concentration established by Kwader (1986) for Seminole County is valid also for the 14 sites surveyed in the SJRWMD;
- b) In deriving chloride content from formation resistivity an average porosity of 25% was used for all sites. Information about porosity of the Floridan aquifer is limited. In one published data set, porosities were computed from geophysical logs over the depth interval between 338 ft and 458 ft. Porosities over this depth range varied between 12% to 32% (NW Florida Water Management District, 1983). Since site specific information about porosities was not available, a porosity value of 25% has been used at all sites.
- c) In the Hawthorn Group, resistivity values are influenced by changes in lithology, porosity, and chloride concentration. This precluded inferring meaningful interpretations about chloride concentrations in the Hawthorn Group. Inferences about water quality are, therefore, ideally drawn only for the carbonate rocks below the Hawthorn Group, and for each site an evaluation must be made of the extent clay stringers in the Hawthorn Group may have influenced the average resistivity value measured. Therefore, if a first layer is resolved in the geoelectric section that is less than 300 ft thick, it likely mainly consists of surficial sediments and the Hawthorn Group, and no information about equivalent chloride concentration can be inferred for such layer.

### Measuring the Geoelectric Section

There are several electrical and electromagnetic methods that can determine from the surface the geoelectric section of the subsurface. These various methods vary in lateral and vertical resolution, and sensitivity to geologic noise. The method selected at the 14 sites throughout the SJRWMD was time domain electromagnetic (TDEM) soundings. The TDEM method was selected because (i) for the objectives at hand it has better lateral and vertical resolution and lower sensitivity to geologic noise than other methods, and (ii) it has proven effective for achieving similar objectives in prior surveys in the SJRWMD and other locations. The principles of TDEM soundings are discussed in the technical note in Appendix A. In this section only the deliverables of a TDEM sounding are reviewed. These deliverables are the same for each site.

It is common in all electrical and electromagnetic methods to transform the voltages and electromotive forces measured into apparent resistivities. The reason for that transformation is to better visualize how the geoelectric section, over which a measurement is made, differs from a geoelectric section with an uniform resistivity. This transformation is also made in TDEM, so that for each site an apparent resistivity curve is shown.

Figure 2-3 shows three computed apparent resistivity curves for three different geoelectric sections. In TDEM effective exploration depth increases with time of measurement after turn-off (Appendix A). In model 1 the resistivity is uniform with depth and the apparent resistivity is constant over the entire time interval. In model 2 true resistivity decreases with depth, and the apparent resistivity curves reflect that. The apparent



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resistivities can be seen to decrease with increasing time. In model 3 the resistivity increases with depth and at later time the apparent resistivity curve also shows an increase. Thus, qualitative information about the geoelectric section can be visualized from displaying the data as apparent resistivities.

The function of an apparent resistivity curve can be further explained for sounding 03-01 from site 3 at Wekiva River (Fig. 2-4). The apparent resistivity values can be seen to continuously decrease with increasing time, and to asymptotically approach a value between 10 ohm-m and 20 ohm-m. Thus, from merely viewing the behavior of the apparent resistivity curve, the conclusions can be drawn that (i) the resistivities decrease with depth, and (ii) the resistivity of the lowest layer within the effective exploration depth of the measurement is between 10 ohm-m and 20 ohm-m.

To derive more quantitative information the experimental data points are submitted to an automatic ridge regression inversion program. This inversion program finds the geoelectric section of the subsurface that best matches the observed data. The inversion program requires an initial model for the geoelectric section. A model consists of the number of layers within the effective exploration depth, and the resistivities and thicknesses for each layer. Such an initial model can be obtained in a number of ways, such as

- approximate matching of apparent resistivity curves with model curves from albums of model curves
- from knowledge of the geoelectric section based on resistivity logs run in drill holes
- from conceptual models formed on the basis of known geology and water quality.

The inversion program is then allowed to adjust the model to improve the fit. This involves the adjustment of resistivities and thicknesses of the layers within the geoelectric model. The inversion program does not change the total number of layers submitted for the model, but all other parameters float freely or optionally can be held constant. To determine the influence of number of layers on the solution, separate inversions with a different number of layers may be run.

The geoelectric section obtained from the inversion routine that best matches the experimental data is shown on the right side of Figure 2-4. It consists of a two-layer geoelectric section consisting of an upper layer 87.6 m thick with a resistivity of 50.4 ohm-m. The second layer has a resistivity of 14.2 ohm-m and its thickness extends beyond the effective exploration depth of the measurement. The solid line on Figure 2-4 represents the computed behavior for the two-layer geoelectric section shown on the right, and the experimental data are superimposed on the solid line.

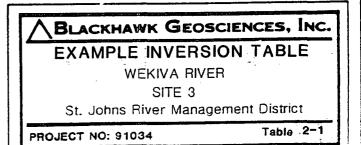
To evaluate the error between the geoelectric section derived from the inversion routine and the experimental data, a tabulation of the inversion and experimental data is also given for each site. The parameters listed on these tables are identified in Table 2-1 for sounding 03-01. Thus, this table lists the error (column 4) between experimental measurements (data, column 2) and calculated data (column 3) for each time gate of measurement (column 1). Also listed on the table is the root mean square (RMS) averaged over all time gates.

### Analysis of Equivalence

The parameters derived for the geoelectric section by the ridge regression inversion are not unique, but generally a range of values will equally fit the observed data within the overall RMS error. This phenomena is called equivalence, and the range of equivalence differs for each parameter of

	DEL: 21	AYERS					
RE	SISTIVITY .	THICKNESS	ELEVAT	ION	CONDUCTANC	E (S)	
	(OHM-M)	(M)	(M)	(FEET)	LAYER	TOTAL	
			12.2	40.0		Solution Geo	electric Sectior
	50.36	87.6	-75.4	-247.3	1.7	1.7	
	14.19						· · · · · · · · · · · · · · · · · · ·
	TIMES	DATA	CALC	% ERROR	STD ERR	Inversion Table	
1	8.90E-05	5.46E+02	5.36E+02	1.955			
2	1.10E-04	4.24E+02	4.14E+02	2.399			
3	1.40E-04	3.21E+02	3.15E+02				
4	1.77E-04	2.46E+02	2.46E+02	0.182			
5	2.20E-04	1 <b>.9</b> 4E+02	1.97E+02	-1.253			
6	2.80E-04	1.51E+02	1.55E+02	<sup>;</sup> -2.519		<b>n</b>	
7	3.55E-04	1.20E+02	1.23E+02	-3.069			
8	4.43E-04	9.80E+01	1.01E+02	-2.880			
9	5.64E-04	8.08E+01	8.16E+01	-0.902			
0	7.13E-04	6.75E+01	6.70E+01	0.826			
11	8.81E-04	5.75E+01	5.72E+01	0.541			
12	1.10E-03	4.99E+01	4.91E+01	1.570		je.	
13	1.41E-03	4.32E+01	4.17E+01	3.603		7	
4	1.80E-03	3.78E+01	3.64E+01	4.045			
5	2.22E-03	3.42E+01	3.28E+01	4.404			
6	2.83E-03	2.94E+01	2.93E+01	0.574			
17	3.55E-03	2.56E+01	2.68E+01	-4.320			
8	4.43E-03	2.43E+01	2.49E+01	-2.452			
19	5.64E-03	2.32E+01	2.30E+01	0.983			
20	7.13E-03	2.20E+01	2.18E+01	1.075		e	
21	8.81E-03	2.06E+01	2.07E+01	-0.497			
22	1.10E-02	1.96E+01	1.97E+01	-0.497			
23	1.41E-02	1.86E+01	1.89E+01	-1.604			
24 25	1.80E-02 2.22E-02	1.76E+01 1.70E+01	1.82E+01 1.77E+01	-3.485 -4.029		· · · · · ·	
25 26	2.22E-02 2.85E-02	1.80E+01	1.72E+01				
20 27	3.60E-02	1.72E+01	1.68E+01	2.225			

R: 237. X: 0. Y: 237. DL: 475. REQ: 264. CF: 1.0000 TDHZ ARRAY, 27 DATA POINTS, RAMP: 190.0 MICROSEC, DATA: 03-01



-14-

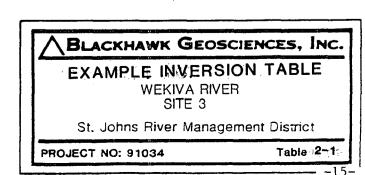
RMS LOG ERROR: 1.66E-02, ANTILOG YIELDS LATE TIME PARAMETERS

\* Blackhawk Geosciences, Incorporated \*

PARAMETER RESOLUTION MATRIX: "F" MEANS FIXED PARAMETER P 1 1.00 P 2 0.00 1.00 T 1 0.00 0.00 1.00 P 1 P 2 T 1

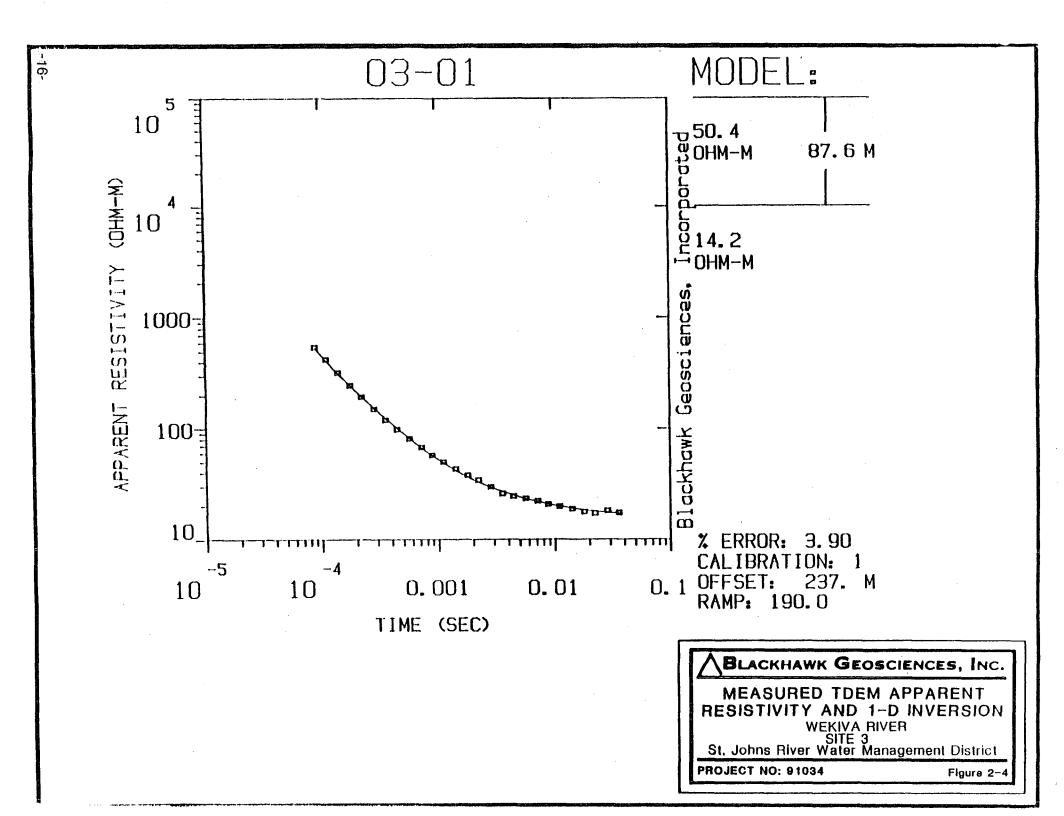
PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

LAYE	R	MINIMUM	BEST	MAXIMUM	<ul> <li>Result of Computation of Equivalence</li> </ul>
RHO	1	46.620	50.357	54.730	
	2	13.612	14.189	14.831	
THICK	1	79.931	87.560	94.326	
DEPTH	1	79.931	87.560	94.326	



**RMS Error** 

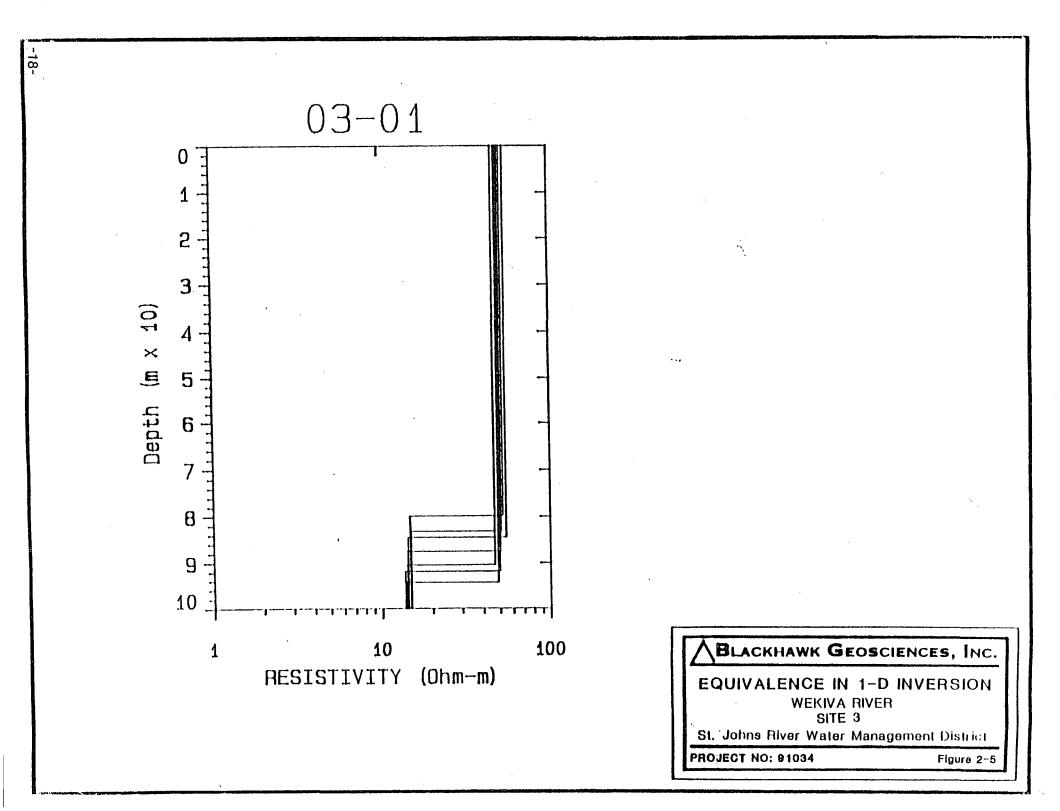
3.8953 %



a geoelectric section. It is a measure of how well each parameter is resolved, and for each sounding the equivalence was evaluated.

The equivalence analysis for sounding 03-01 is shown on Figure 2-5, and the upper and lower bound for each parameter of the geoelectric section is also shown on Table 2-1. Thus, at this site the largest range of equivalence is in determining the depth to the second layer. It may vary from 80 m to 94 m and still result in the same RMS error. The ranges of equivalence for the resistivities of the first and second layer are relatively small.

Examination of the equivalence of the 14 sites throughout the SJRWMD reveals that ranges of equivalence are dependent on the particular geoelectric section encountered. Also when the number of layers increases, the range of equivalence of some parameters in the section may be quite large.



### 3.0 LOGISTICS, DATA PROCESSING AND INTERPRETATION

Logistics and Equipment

1

The general locations for TDEM measurements were determined by SJRWMD. Using the general locations, BGI personnel positioned the transmitter loops and receiver station with compass and measuring string. The transmitter loop locations were selected to mitigate to the maximum extent possible against interferences caused by buried pipelines, other utility lines, fences, power lines and buildings. The transmitter loops were formed of 12-gauge insulated wire laid out in either a square or rectangle on the surface. Dimensions of the loops were determined by the depth of exploration required and availability of open land. At sites 7, 8 and 9 loop sizes of 200 ft by 200 ft were employed, and at the remaining sites the transmitter loop sizes ranged from 1,000 ft by 1,000 ft to 1,500 ft by 1,500 ft, depending on access. Most of the transmitter loops were square in shape, but in some cases access required that rectangular loops be used. The actual dimensions of the transmitter loops at each sounding location are shown on the location maps in Section 4.2 through 4.15. The locations of the loops were plotted on maps by SJRWMD personnel.

The TDEM equipment utilized on the survey was the Geonics EM-37. Some of the characteristics of that instrument are discussed in Appendix A. Table 3-1 summarizes the daily field activities. The field crew consisted of three persons, a two person BGI crew consisting of one geophysicist and one geophysical technician, and a representative of SJRWMD was also present during the field work.

Table 3-1. Daily log of field activities.

<u>Date (1991)</u>	Activity
August 12	Mobilize from Denver, CO to Eustis, FL.
August 13	Read sounding at Site 1 (Wekiva Springs State Park).
August 14	Read soundings at Site 2 (Rock Springs State Preserve) and Site 4 (Simpson Training Center).
August 15	Read sounding at Site 3 (Wekiva River).
August 16	Read additional sounding at Site 4 (Simpson Training Center) and read sounding at Site 5 (Royal Trails).
August 17	Read sounding at Site 6 (Ormond Beach West Well Field).
August 18	Read soundings at Sites 7 (Astronaut High School), 8 (North Brevard Waste Water Treatment Plant), and 9 (Parrish Grove).
August 19	Read sounding at Site 10 (New Smyrna Beach Western Well Field).
August 20	Read sounding at Site 11 (Orange County Landfill).
August 21	Read sounding at Site 12 (Deseret Ranch).
August 22	Read soundings at Sites 13 (Lake Mary) and 14 (Wekiva Springs Road).
August 23	Demobilize from Orlando, FL to Denver, CO.

### Data Acquisition

At all stations the electromotive force (emf) due to the vertical magnetic field in the center of transmitter loops was recorded at several amplifier gains and opposite receiver polarities. Receiver coils with effective areas of  $100 \text{ m}^2$  and  $1,000 \text{ m}^2$  were employed. The data were recorded at base frequencies of 3 Hz and 30 Hz. The current driven through the transmitter loops varied from 13 amps to 24 amps, depending on loop size. All data was recorded on a DAS-54 data logger.

### Mitigation of Inductive Noise Sources

A concern in the acquisition of TDEM data is to recognize interferences caused by induction in metallic structures, such as buried utility lines, fences, grounded power lines and buildings. The primary magnetic field of a transmitter induces eddy current flow in such structures, and these eddy currents in turn cause a secondary magnetic field. That secondary magnetic field can interfere with TDEM measurements, and because it is coherent with the transmitter, it cannot be removed by stacking.

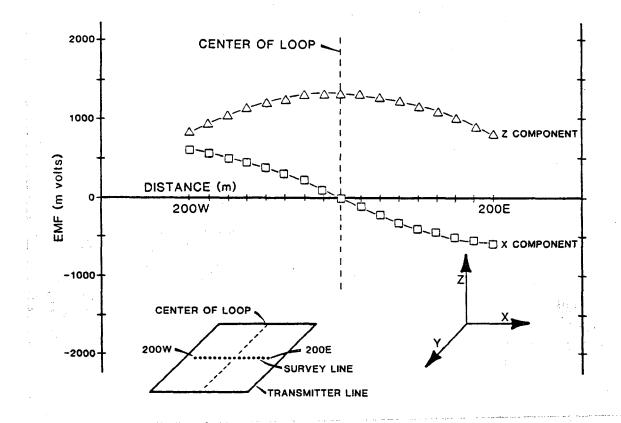
The influence of these inductive noise sources, however, can be recognized from anomalies caused in the spatial behavior of the fields. Figure 3-1 shows a typical behavior of the emf due to the vertical magnetic field on a profile through the center of the loop. This data is typical of that observed over layered ground. Figure 3-1 shows that emf, is relatively flat on a profile through the center of the loop. Measurements at different locations within the loop should nearly coincide for measurements over layered ground. This is illustrated in Figure 3-2 where apparent resistivity curves for soundings at five different locations in the center of the loop have been superimposed. The data virtually coincide and from that observation the conclusion is drawn that inductive noise is minimum.

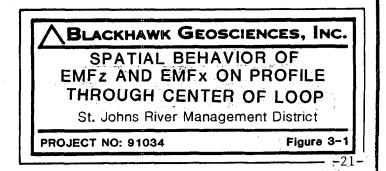
Figure 3-3 on the other hand shows an example of five data sets measured at different locations within the loop with considerable deviation between the data sets. The cause of deviation is inductive interference by metallic structures, and the best way to mitigate this noise is to reposition the transmitter loop to minimize this effect.

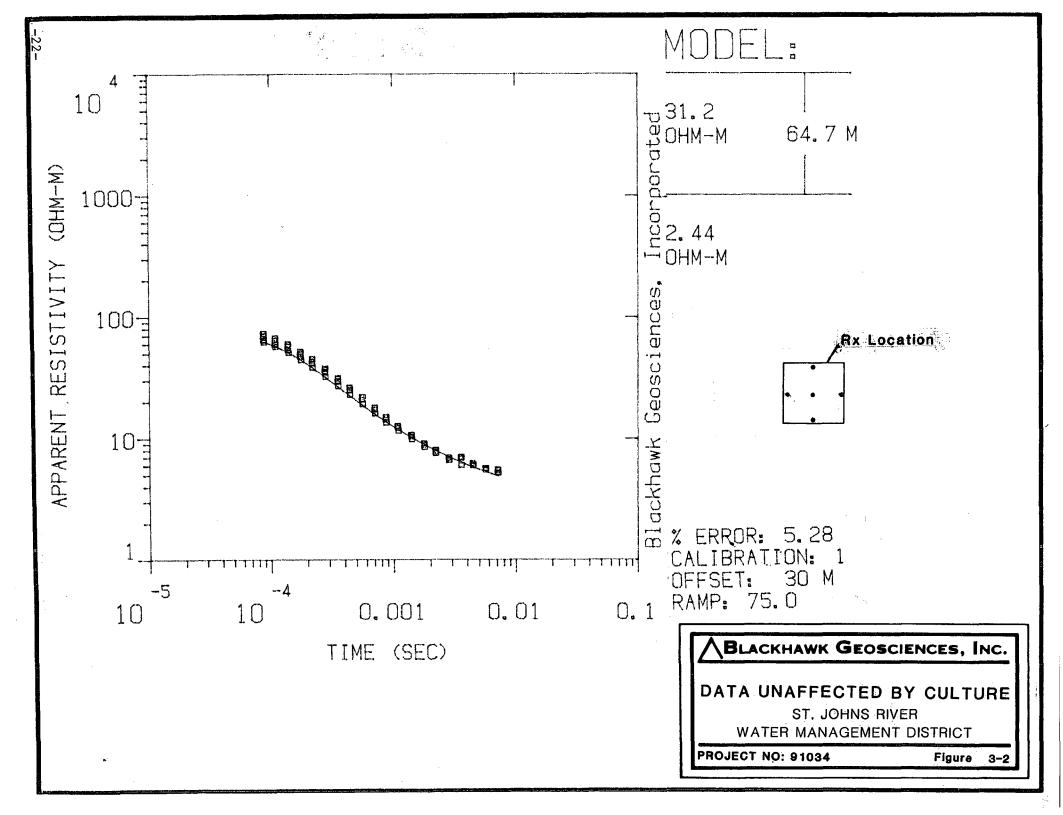
### Data Processing

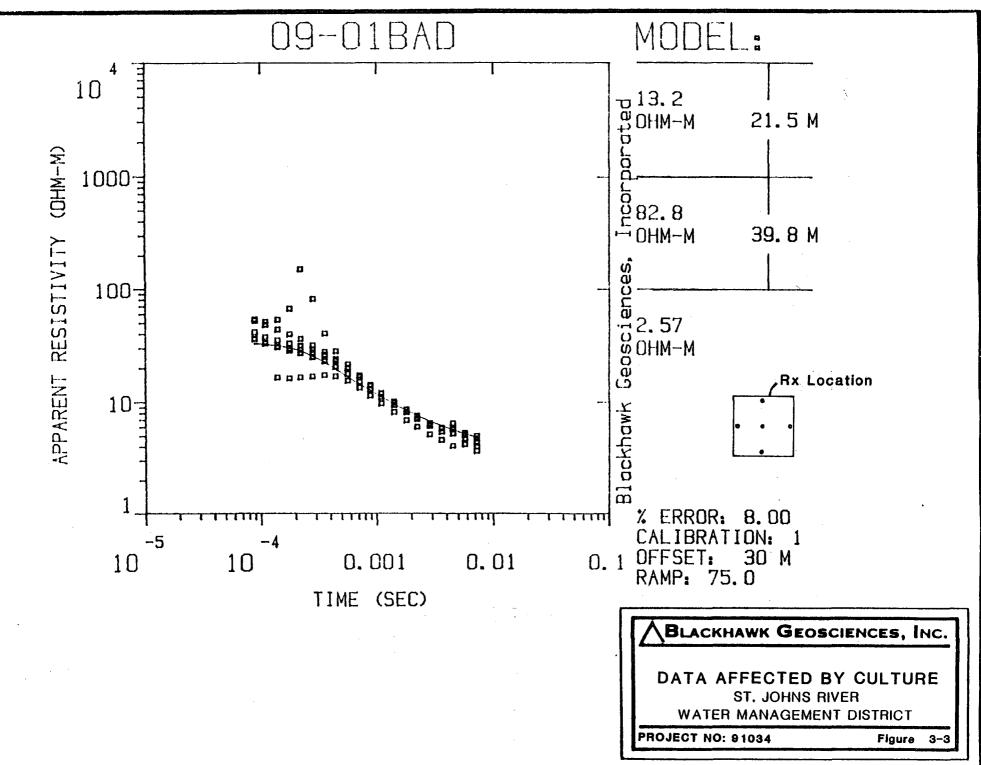
The data stored in the DAS-54 solid memory logger was transferred each day to floppy disks on a computer. The first step in data processing was to average the emf's recorded at opposite receiver polarities. Next, the recordings at different amplifier gains and frequencies are combined to produce one transient decay. The emf's in the various time gates of this decay curve are subsequently entered into a ridge regression inversion program to obtain a one-dimensional geoelectric section that matches the observed decay curve.

# SPATIAL BEHAVIOR OF EMFS









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### 4.0 RESULTS AND DISCUSSION

### 4.1 SUMMARY OF RESULTS

In Sections 4.2 through 4.15 the results of the TDEM measurements at the 14 sites are given. In these sections first the geoelectric section is shown, and subsequently information about first depth of occurrence of highly saline water, water quality in the Upper Floridan aquifer where possible, and depth of occurrence of the 250 mg/l isochlor is inferred from the geoelectric section. In this section the results of all 14 sites are summarized.

Table 4-1.1 lists the sites, the dimensions of the transmitter loops employed, the County in which the site is located, and the longitude and latitude of the center of the loop. In Table 4-1.2 the geoelectric section measured at each site is summarized. It first lists the number of layers of different resistivities used in the inversion. This number corresponds to the number of distinct resistivity boundaries resolved with surface TDEM within the effective exploration depth of the measurement. For each parameter of the geoelectric section (resistivities and thicknesses) the range of equivalences are given in terms of the minimum and maximum value this parameter can assume, and the "best" value.

Table 4-1.3 summarizes the first depth of occurrence of ground water with a chloride concentration greater than 1,000 mg/l. Since information about chloride concentration is derived from resistivity values, the correlation of resistivities to chloride concentrations require assuming a value of porosity. In Table 4-1.3 chloride concentrations are derived for three porosity values - 25%, 30% and 35%. The depth to saline water listed on Table 4-1.3 is expected to represent the top of the boundary between ground water of chloride concentrations less than 1,000 mg/l and ground water with chloride concentrations greater than 1,000 mg/l. Likely chloride concentrations gradually increase below this interface, and gradually decrease above it. The inversion of TDEM data can, at present, only resolve distinct boundaries, and transition zones are not mapped. The depth listed likely represent some average depth within the transition zone.

Water quality measurements were available from a number of sampled wells near TDEM soundings (Toth et al, 1989). The location of these wells is shown on various figures for the individual sites. Table 4-1.4 lists the chloride, hardness, sulfate, and iron concentrations in these wells for the periods of April 1973 to July 1974, and March 1986 and October 1986. The data in this table are referred to on several occasions in discussing the results of individual sites.

Site Name	Site #	County	Loop Size (ft)	Latitude	Longitude
Vekiva Springs State Park	1	Orange	1250 x 1500	28 <sup>0</sup> 44' 10"	81 <sup>0</sup> 28' 53"
Rock Springs State Preserve	2	Orange	1250 x 1450	28 <sup>0</sup> 45' 40"	81 <sup>0</sup> 26' 52"
Vekiva River	3	Lake	1400 x 1600	28 <sup>0</sup> 49' 23"	81 <sup>0</sup> 25' 39"
Simpson Training Center	4-1 4-2	Lake Lake	1100 x 1400 500 x 500	28 <sup>0</sup> 50' 59" 28 <sup>0</sup> 50' 55"	81 <sup>0</sup> 32' 08" 81 <sup>0</sup> 32' 04"
Royal Trails	5	Lake	1000 x 1300	28 <sup>0</sup> 56' 52"	81 <sup>0</sup> 26' 29"
ormond Beach West Well Field	6	Volusia	1200 x 1400	29 <sup>0</sup> 11' 07"	81 <sup>0</sup> 12' 53"
stronaut High School	7	Brevard	200 x 200	28 <sup>0</sup> 37' 27"	80 <sup>0</sup> 50' 53"
North Brevard WWTP	8	Brevard	200 x 200	28 <sup>0</sup> 41' 54"	80 <sup>0</sup> 52' 43"
Parrish Grove	9	Brevard	200 x 200	28 <sup>0</sup> 40' 39"	80 <sup>0</sup> 50' 53"
New Smyrna Beach West Wellfield	10	Volusia	1400 x 1500	29 <sup>0</sup> 01' 09"	81 <sup>0</sup> 07' 38"
Drange County Landfill Site	11	Orange	1500 x 1600	28 <sup>0</sup> 28' 05"	81 <sup>0</sup> 13' 17"
Deseret Ranch	12	Orange	1500 x 1500	28 <sup>0</sup> 23' 20"	81 <sup>0</sup> 00' 51"
Lake Mary	13	Seminole	1100 x 1100	28 <sup>0</sup> 46' 07"	81 <sup>0</sup> 20' 38"
Jekiva Springs Road	14	Lake	950 x 950	28 <sup>0</sup> 47' 56"	81 <sup>0</sup> 25' 52"

Table 4-1.1 Geographic information about measurement sites

				Laver 1						Layer 2					Laver 3							Laver 4		
	Number of Modeled Layers in Geoelectric	Ρı	Resisti (ohm-me	ivity sters)		Thickr h <sub>1</sub> (mete		ρz	Resist (ohm-ma			Thick h <sub>2</sub> (met		þ <b>3</b>	Resist (ohan-aa	ivity eters)		Thick h <sub>3</sub> (met		ρ4	Resist (oh <b>n-n</b>	ivity sters)		
Site Name	Section	Hin	Best	Max	Min	Best	Max	Nin	Best	Max	Nin	9est	Мах	Nin	Best	Max	Nin	8es t	Max	Hin	Best	Max		
) Wekiva Springs State Park	2	53	54	56	61	63	66	550	640	780														
) Rock Springs State Preserve	3	29	31	33	120	140	150	220	1100	12,000	150	170	180	14	16	18								
Wekiva River	2	47	50	55	80	88	94	14	14	15														
Simpson Training Center	3	36	48	69	24	41	77	190	465	2568	129	154	174	4.1	6.8	11.4								
Royal Trails	4	64	78	88	57	90	130	180	240	410	310	360	400	1.7	3.3	6.4	18	44	110	20	25	33		
Ormond Beach West Well Field	3	30	39	51	13	23	48	92	98	110	290	320	330	1.9	2.3	2.7								
Astronaut High School	2	27	30	32	53	55	57	2.3	2.5	2.7														
North Brevard WMTP	2	33	35	38	55	56	57	1.5	1.6	1.6														
Parrish Grove	2	29	31	33	62	65	67	2.2	2.4	2.7														
)) New Sayrna Beach West Weilfield	3	30	35	38	61	80	100	110	150	310	180	210	230	7.3	9.0	11.0								
) Orange County Landfill Site	3	31	33	34	80	88	98	440	710	1300	620	640	660	7.2	11	16								
) Deseret Ranch	3	58	190	1900	9,9	17	28	25	26	27	340	400	460	6.2	9.4	14						54 A.		
) Lake Mary	2	43	60	81	21	37	64	240	290	390	320	340	370	4.4	5.5	6.8						``.		
) Wekiva Springs Road	3	21	23	25	32	43	57	45	49	55	230	250	270	7.2	8.7	10.0								

<u>Table 5-1.2</u> Summary of Geoelectric Sections with Range of Equivalence	

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## **Table 4-1.3**

# First Depth of Occurrence and Chloride Concentration of Layer with Chloride Concentrations Greater than 1,000 mg/l Inferred From Geoelectric Section Derived from TDEM Soundings

		Formation Resistivity (ohm-m)	First Depth of Occurrence (ft below surface)	Estimated Chloride Concentration (mg/l) at 3 Porosities					
	· · · · · · · · · · · · · · · · · · ·			Porosity					
				25%	30%	35%			
(1)	Wekiva Springs State Park	Not Detected							
(2)	Rock Springs State Preserve	16	1002	1,850	1,340	1,010			
(3)	Wekiva River	14	287	2,140	1,540	1,170			
4)	Simpson Training Cente	r 6.8	639	4,580	3,490	2,690			
5)	Royal Trails	3.3	1465	9,590	7,130	5,540			
6)	Ormond Beach West Wellfield	2.3	1122	>10,000	>10,000	8,010			
7)	Astronaut High School	2.5	181	>10,000	9,460	7,360			
8)	North Brevard WWTP	1.6	183	>10,000	>10,000	>10,000			
9)	Parrish Grove	2.4	212	>10,000	9,860	7,670			
10)	New Smyrna Beach WWF	9.0	949	3,420	2,520	1,930			
11)	Orange County Landfill	10.6	2389	2,882	2,110	1,620			
12)	Deseret Ranch	9.4	1355	3,270	2,400	1,840			
13)	Lake Mary	5.5	1248	5,700	4,210	3,260			
14)	Wekiva Springs Road	8.7	960	3,540	2,610	2,000			

					Chlo	Chlorides mg/lHardness mg/lSulfate mg/l		<u>/l</u>	Iron #g/l							
<u>COUNTY</u>	<u>Well owner</u>	WELL ID	LAT	LONG	Apr 73- Jul 74	Mar <u>86</u>	0ct <u>86</u>	Apr 73- Jul 74	Mar <u>86</u>	0ct <u>86</u>	Apr 73- Jul 74	Mar <u>86</u>	0ct <u>86</u>	Apr 73- Jul 74	Mar <u>86</u>	0ct <u>86</u>
Lake	Wekiva Falls	L-0048	284740	812517	240	296	300	370	400		200	210	230	0	50	20
	Wekiva Falls	L-0049	<b>28</b> 4740	812517		312	300		390			210	210			
	Woodard	L-0073	284747	812515		352	340		440			225	220		50	60
	Dorton's Barn	L-0047	284753	812621		57	71		190			50	57		188	540
	Mock	L-0076	284800	812523		244	220		370			190	190		50	120
	Major Realty	L-0038	284933	812558	282	176	220	420	308		165	145	150	210	452	430
Orange	Fl Wek. Pres.	OR0054 OR0035	284541 284429	812652 812720	200 40	208 45	220 77	324 156	900 150		180 45	145 18	160 50	30 20	50 50	60 10

# Table 4-1.4. Chloride, hardness, sulfate, and iron concentrations in sampled wells for the period April 1973 to July 1974, March 1986 and October 1986 (after Toth et al, 1989)

### 4.2 WEKIVA SPRINGS STATE PARK (SITE 1)

### Location and Geoelectric Section

The detailed location of this sounding is shown in Figure 4-2.1. The interpreted geoelectric section is shown in Figure 4-2.2, and it consists of a two layer section. The resistivity of the upper layer is 54 ohm-m and it is 209 ft thick; the resistivity of the second layer is 640 ohm-m and its thickness extends beyond the effective exploration depth of the measurement.

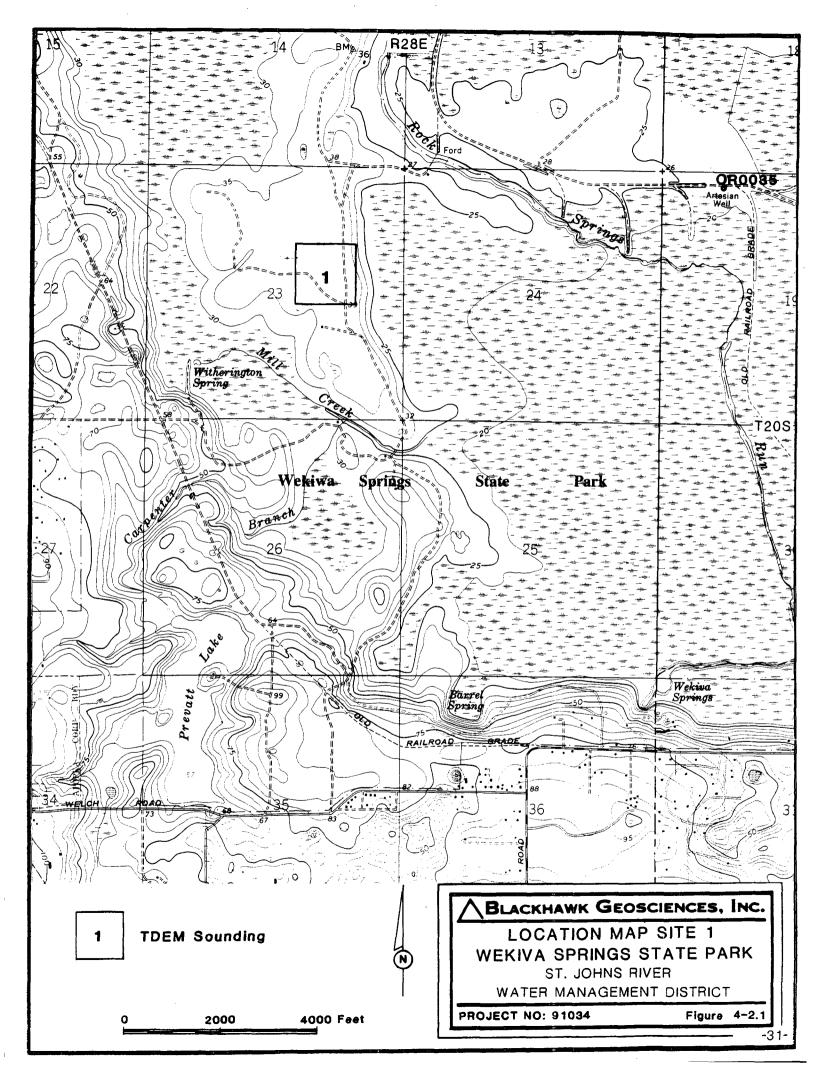
## Depth of Occurrence of High Salinity Water

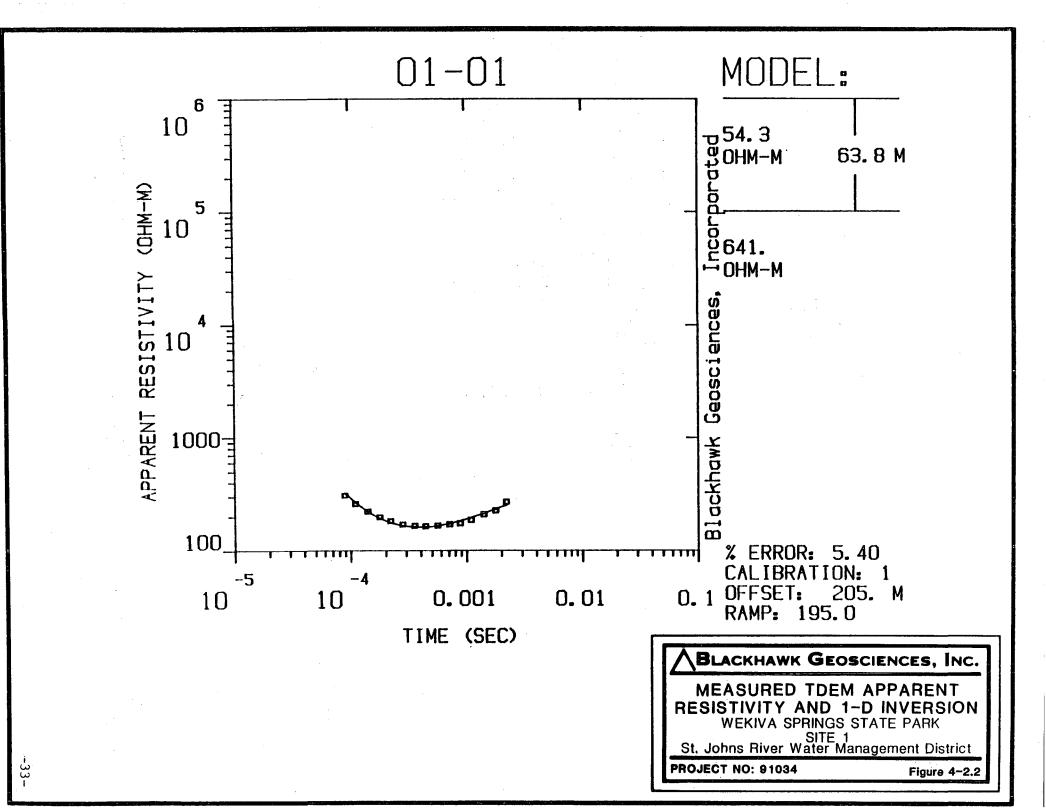
Thus, resistivity values characteristic of saline water were not encountered in this sounding within the effective exploration depth of the measurement. Since one objective of this TDEM survey is to determine the depth of occurrence of the interface with highly saline water, an evaluation was made of the minimum depth a layer of low resistivity would have been detected.

To estimate the minimum depth of occurrence of a low resistivity layer, a series of computations were made placing a low resistivity layer (characteristic of saline water) at various depths below the geoelectric section observed. The minimum depth at which model curves with this low resistivity layer inserted diverges from observed data, represents the effective depth of exploration. The effective exploration depth for detecting a layer with a resistivity of 5 ohm-m or less (more than 5,000 mg/l chloride) was computed to be approximately 2,500 ft. The effective exploration depth decreases somewhat for detecting strata of lower salinity. At a resistivity of 28 ohm-m corresponding to 1,000 mg/l chloride, the effective exploration is about 1,800 ft. The influence of salinity on effective exploration depth is shown in Table 4-2.1

Table 4-2.1.	Influence of salinity and resistivity on effective
	exploration depth for detecting a unit of low
	resistivity placed below the geoelectric section
	observed at Site 1.

Resistivity of Unit of Low Resistivity (ohm-m) Placed Below Geoelectric Section Measured	Equivalent Chloride (mg/l) at 25% Porosity	Effective Depth of Exploration (ft) for Detecting Layer of Low Resistivity (high salinity)				
2.7	10,000	2,500				
5.0	5,000	2,500				
10	3,000	2,150				
20	1,200	2,000				
80	250	1,500				





### Depth of Occurrence of 250 mg/l Isochlor

Another objective of the survey was to determine the depth to the 250 mg/l isochlor. The upper 209 ft with a resistivity of 54 ohm-m likely represents the Hawthorn Group and younger sediments, as well as a portion of the upper Ocala aquifer. The resistivity value of this layer is probably influenced as much by lithology as by chloride concentration. However, the resistivity value of 640 ohm-m below 209 ft is characteristic of the Upper Floridan aquifer from 209 ft below surface to about 1,500 ft below surface (see Table 4-2.1). The chloride concentration of the ground water below about 209 ft is expected to be lower than 250 mg/l based on a formation resistivity of 640 ohm-m, assuming an average porosity of 25%. Measurements in wells in the general area by Toth et al (1989) show the chloride concentration to be below 25 mg/l.

The depth of occurrence of the 250 mg/l isochlor is below the effective depth of exploration of the measurement. It was computed to occur at a depth greater than 1,500 ft below surface.

### Accuracy of Measurement and Interpretation

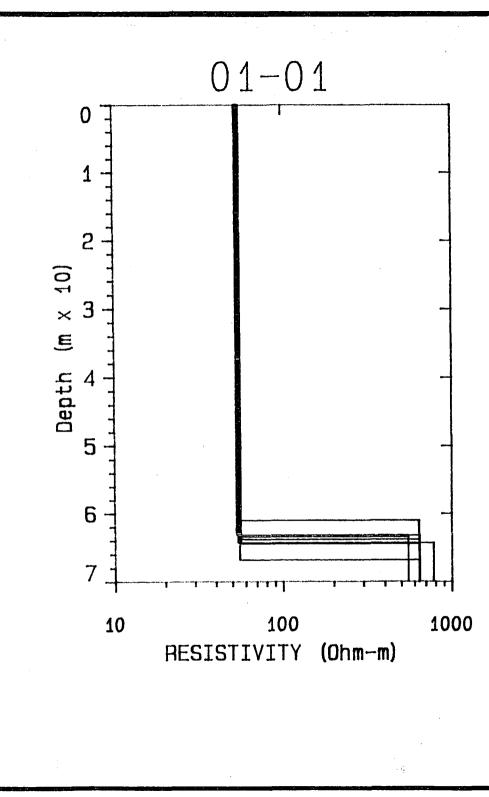
Figure 4-2.3 shows the evaluation of equivalence of the TDEM sounding at this site, and the inversion table (Table 4-2.2) lists the upper and lower boundary of the parameters of the geoelectric section. The important parameter influencing the information about chloride concentration of the Upper Floridan aquifer is the resistivity of the second layer, and the value of that resistivity is determined to be between 550 ohm-m and 776 ohm-m. That range of resistivity is well above the value of 80 ohm-m below which, at an average porosity of 25%, chloride concentrations greater than 250 mg/l are expected (Fig. 4-2.2). The conclusion, inferred from the measurements, that the chloride concentration between 209 ft and 1,500 ft below surface is less than 250 mg/l, is not influenced by accuracy of measurements.

Accuracy of measurements, however, is but one factor influencing validity of the conclusions reached. More important may be the validity of the assumptions made. To conclude that the 250 mg/l isochlor occurs at a depth of 1,500 ft below surface assumes that the average porosity remains near 25%, and the validity of that assumption cannot be tested.

#### Summary of Results of TDEM Measurements at the Wekiva Springs State Park

From the TDEM measurements the following information about water quality was inferred:

- The depth of occurrence of highly saline water was beyond the effective exploration depth of the measurements. With computer modeling, the minimum depth of occurrence of ground water with a chloride concentration of 1,000 mg/l was determined to be about 1,800 ft.
- 2) The chloride concentration in the Upper Floridan aquifer between a depth of 209 ft and 1,500 ft below surface was inferred to be less than 250 mg/l. The depth of occurrence of the 250 mg/l isochlor was inferred to be below 1,500 ft below surface.



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BLACKHAWK GEOSCIENCES, INC. EQUIVALENCE IN 1-D INVERSION WEKIVA SPRINGS STATE PARK SITE 1 St. Johns River Water Management District PROJECT NO: 91034 Figure 4-2.3 01-01

(S) TOTAL

1.2

MODEL: 2 LAYERS

RE	SISTIVITY T	HICKNESS	ELEVATI	ION ·	CONDUCTANCE
	(OHM-M)	(M)	(M)	(FEET)	LAYER
			10.7	35.0	
	54.28	63.8	-53.2 -	174.5	1.2
6	41.08				
	TIMES	DATA	CALC	% ERROR	STD ERR
1	8.90E-05	3.06E+02	3.25E+02	-5.616	
2	1.10E-04	2.58E+02	2.66E+02	-3.048	
3	1.40E-04	2.21E+02	2.21E+02	-0.034	
4	1.77E-04	1.97E+02	1.93E+02	2.364	
5	2.20E-04	1.83E+02	1.76E+02	3.667	
6	2.80E-04	1.70E+02	1.66E+02	2.971	
7	3.55E-04	1.66E+02	1.61E+02	2.831	
8	4.43E-04	1.64E+02	1.61E+02	1.814	
9	5.64E-04	1.66E+02	1.65E+02	0.493	
10	7.13E-04	1.71E+02	1.73E+02	-1.174	
11	8.81E-04	1.74E+02	1.83E+02	-4.824	
12	1.10E-03	1.87E+02	1.96E+02	-4.339	
13	1.41E-03	2.09E+02	2.13E+02	-1.734	
14	1.78E-03	2.26E+02	2.32E+02	-2.383	
15	2.21E-03	2.69E+02	2.51E+02	7.239	

R: 205. X: 0. Y: 205. DL: 409. REQ: 228. CF: 1.0000 TDHZ ARRAY, 15 DATA POINTS, RAMP: 195.0 MICROSEC, DATA: 01-01 WEKIVA SPRINGS STATE PARK

RMS LOG ERROR: 2.28E-02, ANTILOG YIELDS 5.4013 % LATE TIME PARAMETERS

\* Blackhawk Geosciences, Incorporated \*

PARAMETER RESOLUTION MATRIX: "F" MEANS FIXED PARAMETER P 1 0.31 P 2 0.04 0.01 T 1 -0.23 -0.04 0.18 P 1 P 2 T 1

36.



Table 4-2.2

PROJECT NO: 91034

# PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

	LAYER		MINIMUM	BEST	MAXIMUM	
	RHO	1	52.698 550.121	54.277 641.082	55.910 776.137	
a La constante La	THICK	1	61.054	63.848	66.847	
:	DEPTH	1	61.054	63.848	66.847	

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BLACKHAWK GEOSCIENCES, INC. INVERSION TABLE WEKIVA SPRINGS STATE PARK SITE 1 St. Johns River Management District PROJECT NO: 91034 Table 4-2.2

-37

# 4.3 ROCK SPRINGS STATE PARK (SITE 2)

# Location and Geoelectric Section

The location of the Rock Springs sounding is shown in Figure 4-3.1. The interpreted geoelectric section is shown in Figure 4-3.2 and it consists of three layers. The upper layer with a resistivity of 31 ohm-m is 455 ft thick, the second layer is 547 ft thick and displays a resistivity of 1,088 ohm-m. The lower layer has a resistivity of 16 ohm-m and extends to the effective exploration depth of the measurements.

### Geologic Interpretation of Geoelectric Section

The upper layer of 31 ohm-m likely represents surficial sediments, the Hawthorn Group, and portions of the upper Floridan aquifer. The transition from 31 ohm-m to 1,088 ohm-m may relate to the middle semi-confining unit reported (Tibbals 1990) to occur in this general area at a depth of about 375 ft below surface. The discrepancy in depth between 375 ft reported by Tibbals (1990) and the resistivity boundary at 455 ft may be due to (i) the regional information displayed by Tibbals (1990) versus the local measurements of the TDEM sounding, and (ii) resistivity boundaries may reflect both changes in porosity as well as salinity which two parameters need not necessarily coincide.

# Depth of Occurrence of High Salinity Water

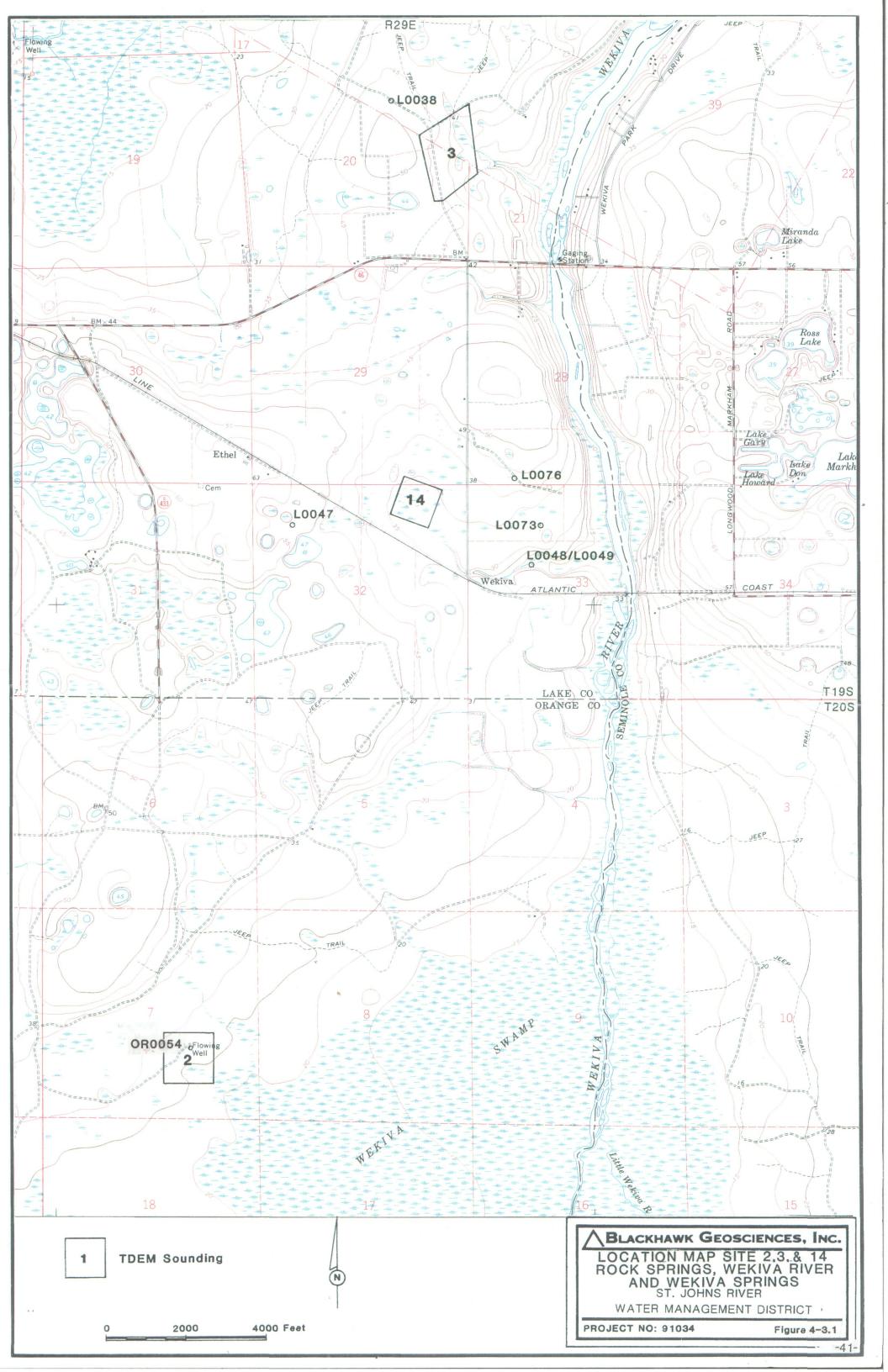
The lowest resistivity encountered in the geoelectric section is 16 ohmm, and the top of this layer occurs at a depth of 977 ft below msl. Assuming an average porosity for the carbonate rocks of 25%, the chloride content of ground water at that depth is expected to be in excess of 1,850 mg/l (Fig. 2-2).

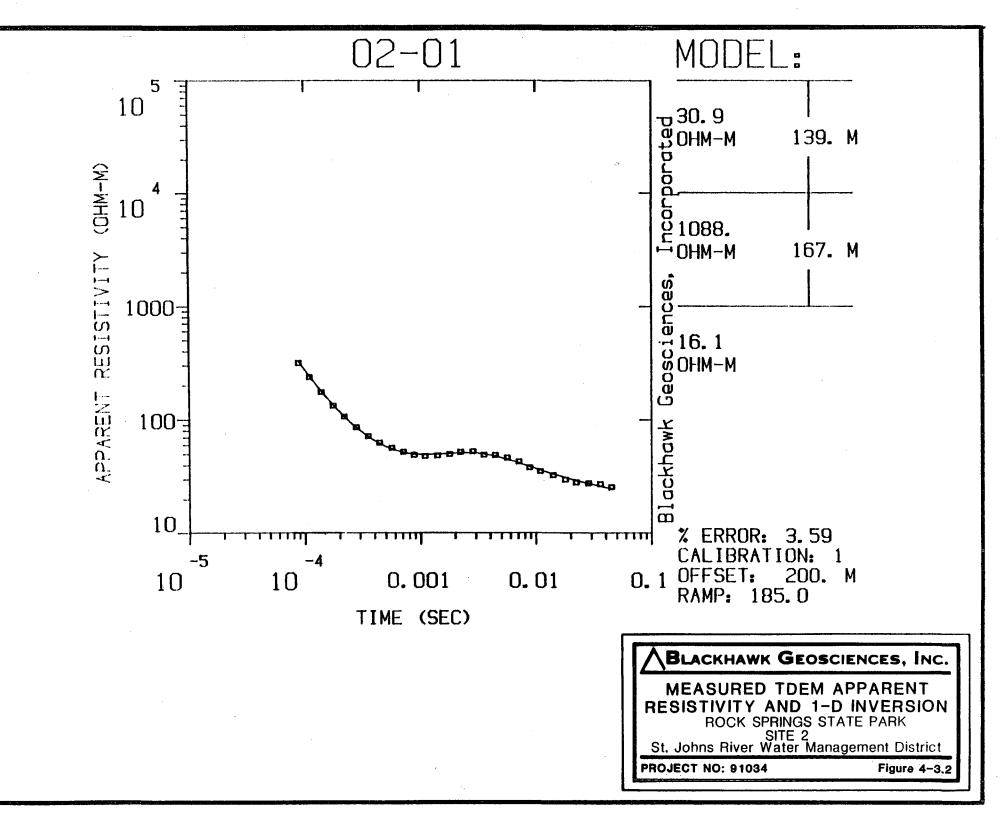
# Depth of Occurrence of 250 ppm Isochlor

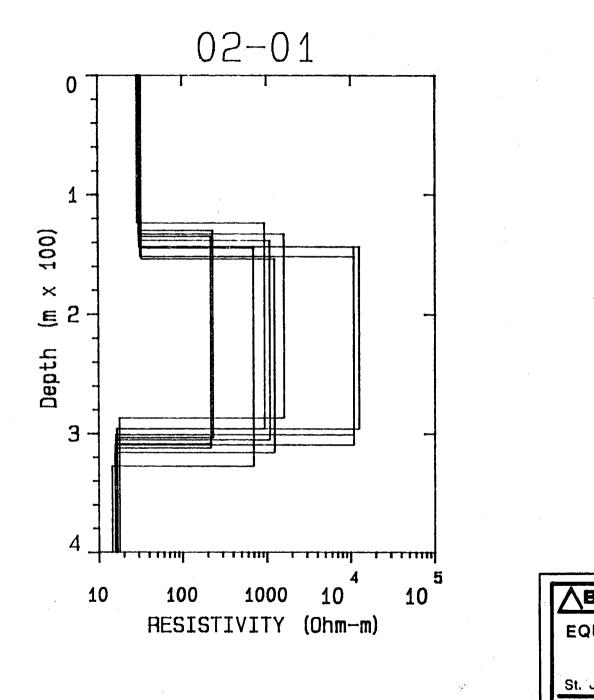
The upper layer in the geoelectric section has a resistivity of 31 ohm-m. This layer likely represents surficial sediments and the Hawthorn Group, as well as portions of the Upper Floridan aquifer. Because this layer appears several hundred feet thicker than the Hawthorn Group and younger sediments, the resistivity of 31 ohm-m likely also is characteristic of the first few hundred feet of the carbonate rocks of the Upper Floridan aquifer. Assuming an average porosity of 25%, 31 ohm-m corresponds to chloride concentrations in excess of 500 mg/l chloride. Toth et al (1989) measured chloride concentrations between 208 and 220 mg/l in March and October of 1986. In addition, he observed sulfate concentrations between 145 to 160 mg/l. The water quality measured in the wells shown on the map in Figure 4-3.1 are given in Table 4-1.4.

There is thus a discrepancy between chloride concentration inferred from the geoelectric section measured with TDEM and measured in wells by Toth et al (1989). Reasons for this disagreement might be:

- 1) The resistivity of 31 ohm-m for the upper 455 ft, which includes surficial sediments, the Hawthorn Group, and the Upper Floridan, may be influenced by some clay layers in both the Hawthorn Group and surficial sediments which would lower the resistivity.
- 2) In inferring chloride concentration from formation resistivity an average porosity of 25% is assumed, and the chemical position of the ground water is assumed similar to that measured by Kwader (1986) for wells in Seminole County. The relative high levels of sulfates measured by Toth et al (1989) indicate that the chemical composition of the ground water at site 2 may differ from the wells on which the data in Figures 2-1 and 2-2 are based.







BLACKHAWK GEOSCIENCES, INC.			
EQUIVALENCE IN 1-D INVERSION ROCK SPRINGS STATE PARK			
SITE 2			
St. Johns River Water Management District			
PROJECT NO: 91034 Figure 4-3.3	3		

3) The time difference of approximately five years between measurements on wells and the TDEM survey.

## Accuracy of Measurement and Interpretation

Figure 4-3.3 shows the evaluation of equivalence for the TDEM sounding at this site and the inversion table lists the upper and lower boundary of the parameters of the geoelectric section. The important parameter influencing information inferred about chloride concentration is the resistivity of the upper layer. The range of equivalence in the resistivity of that layer is between 29 ohm-m and 33 ohm-m. Variation in resistivity over that range does not negate the conclusion that the chloride concentration is expected to exceed 250 mg/l chloride in the Upper Floridan aquifer.

The main source of error in deriving water quality is not in the TDEM measurements, but in the assumptions necessary to infer from resistivity measurements information about chloride concentration.

## <u>Summary of Results of TDEM Measurements at the Rock Springs State Park</u> (Site 2)

From the TDEM measurements the following information about aquifer characteristics and water quality was derived:

- 1) The depth of occurrence of saline water was interpreted to occur at elevations below 977 ft below msl. The chloride concentration at that depth is expected to exceed 1,850 mg/l.
- 2) On the basis of the resistivity value measured for the surficial sediments, the Hawthorn Group, and the Upper Floridan aquifer, the chloride concentration is expected to exceed 250 mg/l for the entire aquifer. This is, however, not consistent with water samples measured in wells in 1986 and reported by Toth et al (1989). They showed chloride concentrations just below 250 mg/l but also relatively high sulfate concentrations were measured. There also is a possibility that the resistivity value measured is influenced by clay stringers in the Hawthorn Group and younger sediments.
- 3) An increase in resistivity (from about 31 ohm-m to 1,088 ohm-m) is observed at a depth of 430 ft below msl. This change is interpreted to be related to a change in porosity in the carbonate rocks. It may represent a middle semi-confining unit.

MODEL: 3 LAYERS

RE	SISTIVITY	THICKNESS	ELEVAT	ION	CONDUCTANC	E (S)
	(OHM-M)	(M)	(M)	(FEET)	LAYER	TOTAL
			7.6	25.0		
	30.91	138.7	-131.1	-430.0	4.5	4.5
10	88.34	166.8	-297.8	-977.1	0.2	4.6
	16.08					
	TIMES	DATA	CAL C	% F0000		
	THES	DATA	CALC	% ERROR	STD ERR	
1	8.90E-05	3.16E+02	3.24E+02	-2.470		
2	1.10E-04	2.37E+02	2.42E+02	-2.150		
3	1.40E-04	1.75E+02	1.77E+02	-1.491		
4	1.77E-04	1.33E+02	1.34E+02	-1.347		
5	2.20E-04	1.06E+02	1.07E+02	-0.283		
6	2.80E-04	8.54E+01	8.51E+01	0.387		
7	3.55E-04	7.13E+01	7.07E+01	0.892		
8	4.43E-04	6.25E+01	6.16E+01	1.530		
9	5.64E-04	5.63E+01	5.51E+01	2.077		
10	7.13E-04	5.20E+01	5.14E+01	1.287		
11	8.81E-04	4.89E+01	4.96E+01	~1.380		
12	1.10E-03	4.78E+01	4.91E+01	-2.556		
13	1.41E-03	4.84E+01	4.95E+01	-2.291		
14	1.80E-03	4.98E+01	5.05E+01	-1.323		
15	2.22E-03	5.18E+01	5.11E+01	1.367		
16	2.85E-03	5.25E+01	5.11E+01	2.888		
17	3.55E-03	4.91E+01	4.99E+01	-1.695		
18	4.43E-03	4.87E+01	4.78E+01	1.812		
19	5.64E-03	4.62E+01	4.48E+01	3.212		
20	7.13E-03	4.28E+01	4.17E+01	2.637		
21	8.81E-03	3.81E+01	3.89E+01	-1.908		
22	1.10E-02	3.54E+01	3.62E+01	-2.244		
23	1.41E-02	3.27E+01	3.35E+01	-2.267		•
24	1.80E-02	2.98E+01	3.12E+01	-4.325		
25	2.22E-02	2.82E+01	2.94E+01	-3.939		
26	2.85E-02	2.77E+01	2.76E+01	0.467		
27	3.60E-02	2.73E+01	2.61E+01	4.673		
28	4.49E-02	2.57E+01	2.49E+01	3.315	~	
						<u>∧</u> e



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R: 200. X: 0. Y: 200. DL: 400. REQ: 222. CF: 1.0000 TDHZ ARRAY, 28 DATA POINTS, RAMP: 185.0 MICROSEC, DATA: 02-01 ROCK SPRINGS STATE PRESERVE

RMS LOG ERROR: 1.53E-02, ANTILOG YIELDS 3.5855 % LATE TIME PARAMETERS

\* Blackhawk Geosciences, Incorporated \*

PARAMETER RESOLUTION MATRIX: "F" MEANS FIXED PARAMETER P 1 1.00 P 2 0.00 0.01 P 3 0.00 0.00 1.00 T 1 0.00 -0.02 0.00 1.00 T 2 0.00 0.03 0.00 0.00 1.00 P 1 P 2 P 3 T 1 T 2

# PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

LAYER		MINIMUM BEST		MAXIMUM	
RHO	1	29.271	30.905	32.561	
	2	219.345	1088.344	12556.611	
	3 14.423		16.085	17.715	
тніск	1	124.244	138.679	154.120	
	2	152.240	166.771	182.799	
DEPTH	1	124.244	138.679	154.120	
	2	287.075	305.449	327.784	



# 4.4 WEKIVA RIVER (SITE 3)

# Location and Geoelectric Section

The detailed location of this sounding is shown in Figure 4-4.1. The interpreted geoelectric section is shown in Figure 4-4.2, and consists of two layers. The upper layer has a resistivity of 50 ohm-m and a thickness of 287 ft. The lower layer has a resistivity of 14 ohm-m and was interpreted at a depth of 287 ft below surface or 247 ft below msl.

### Geologic Interpretation of Geoelectric Section

The upper layer with a resistivity of 50 ohm-m and a thickness of 287 ft likely represents the Hawthorn Group and younger sediments, as well as a portion of the Upper Floridan aquifer.

# Depth of Occurrence of High Salinity Water

The lowest resistivity encountered within the effective exploration depth of the measurement is 14 ohm-m, and it is first encountered at a depth of 247 ft below msl. Assuming an average porosity of 25%, the chloride concentration of ground water at that depth is expected to exceed 2,140 mg/l.

#### Depth of Occurrence of 250 mg/l Isochlor

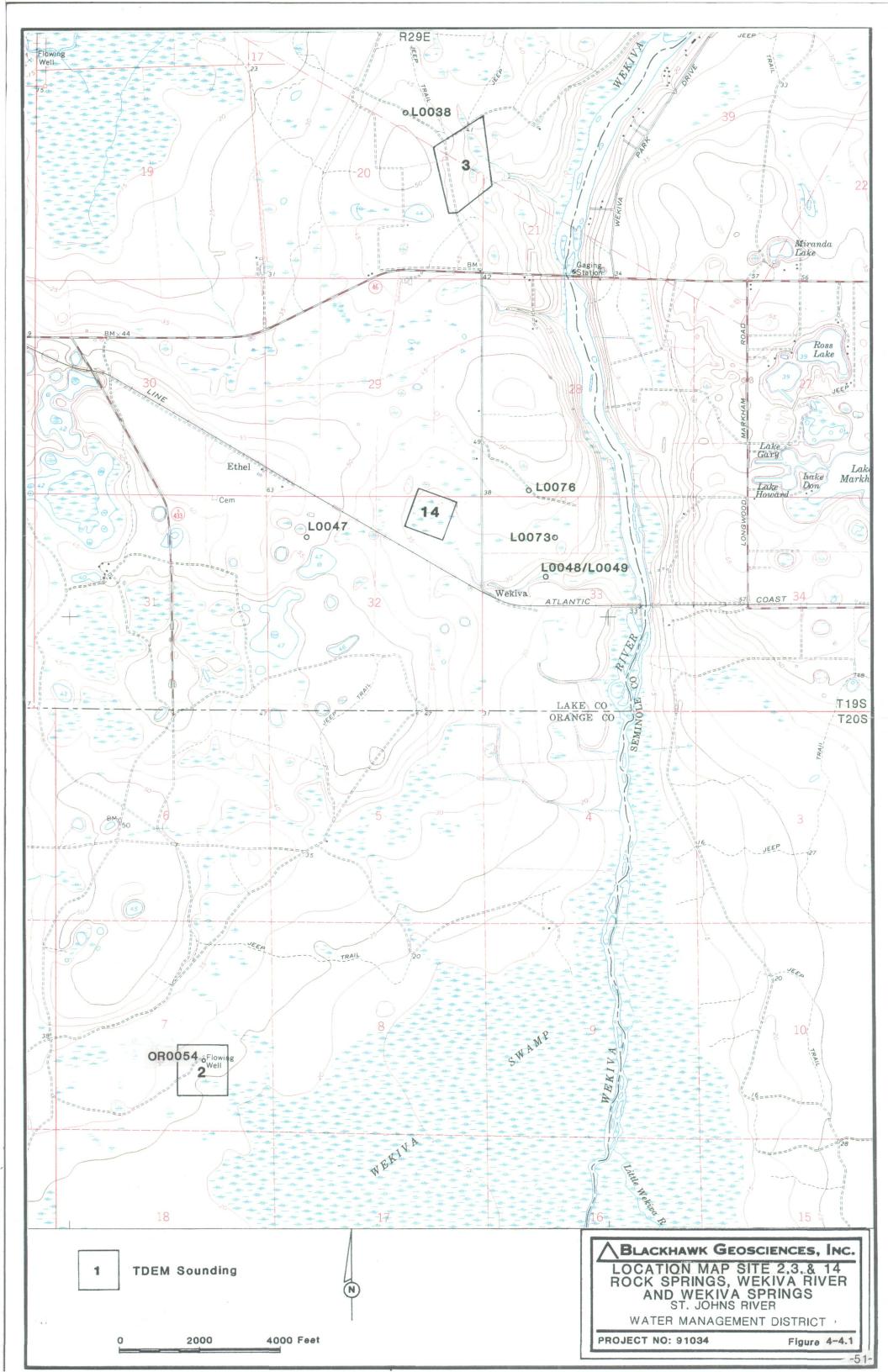
The upper layer in the geoelectric section has a resistivity of 50 ohm-m, and this layer is interpreted to be 287 ft thick. This layer likely includes part of the Upper Floridan aquifer below the Hawthorn Group. Even though this layer appears to be 100 ft to 200 ft thicker than the Hawthorn Group and younger sediments, the resistivity of 50 ohm-m may be influenced by clay layers in the Hawthorn Group and younger sediments. For that reason no information about chloride concentration has been inferred for this upper layer.

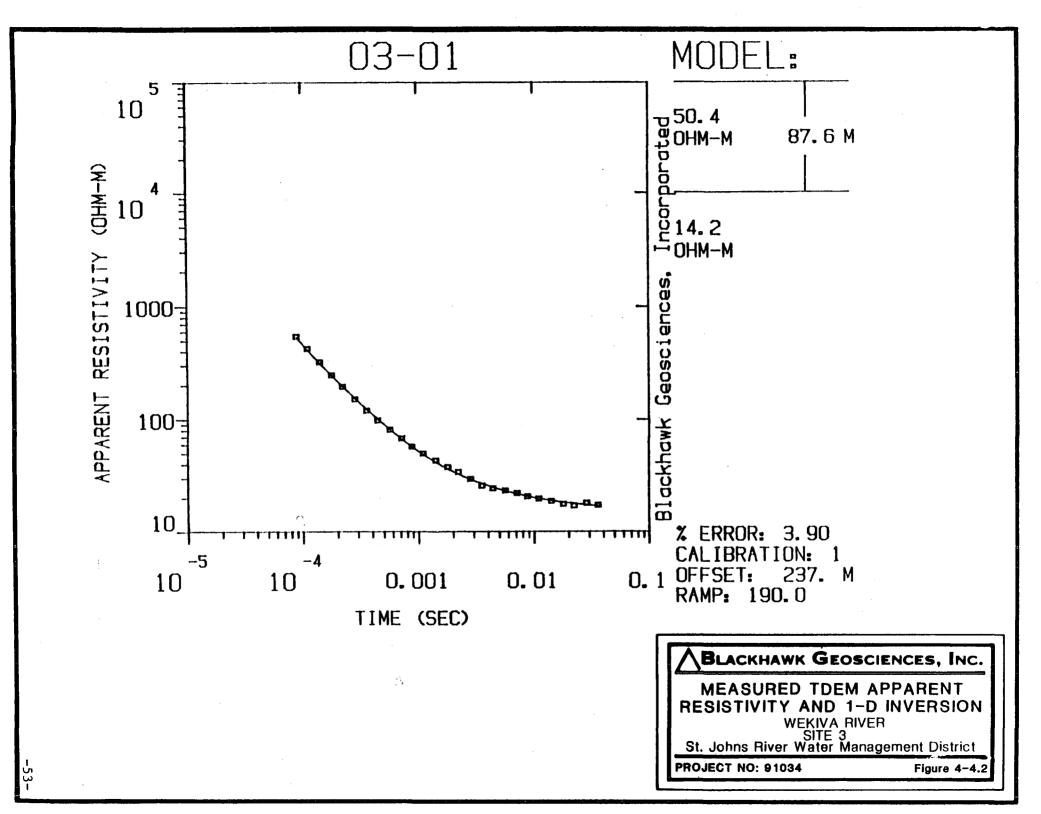
Toth et al (1989) in 1986 measured chloride concentrations between 176 mg/l (March) and 220 mg/l (October) in Well L-0038 nearby (Table 4-1.3). The sulfate concentrations for the same measurement periods were 145 mg/l and 150 mg/l.

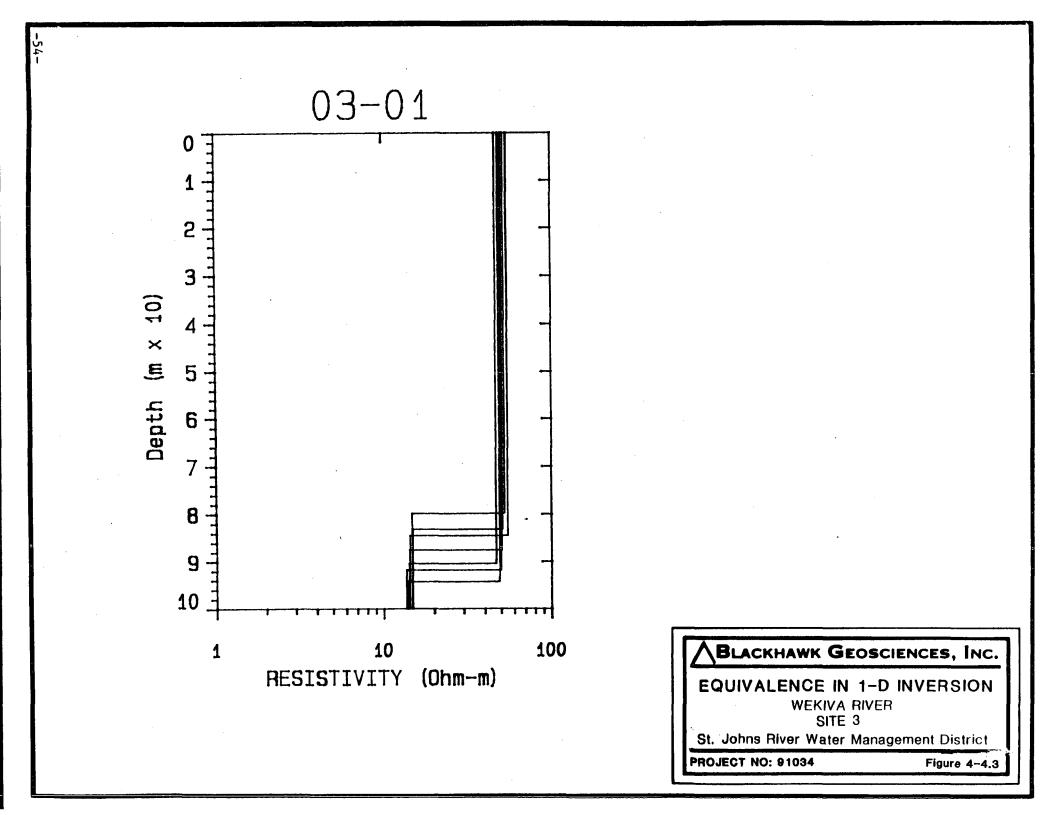
### Accuracy of Measurement and Interpretation

Figure 4-4.3 shows the evaluation of equivalence for the TDEM sounding at this site and the inversion table (Table 4-4.1) lists the upper and lower boundaries of the parameters of the geoelectric section. The important parameter influencing information inferred about chloride concentration is the resistivity of the upper layer. The range of equivalence in the resistivity of that layer is between 46 ohm-m and 55 ohm-m.

The main source of error in deriving water quality is not in the geophysical measurement, but in the assumptions necessary to infer from resistivity measurements information about chloride concentration. Since the resistivity of the first layer is expected to be influenced by clay layers in the Hawthorn Group and younger sediments, no information about chloride concentration is inferred for the first 100 ft to 150 ft of the Upper Floridan aquifer.







M	DEL 2					
-	ODEL: 2	LATERS				
RI	ESISTIVITY	THICKNESS	ELEVAT	ION	CONDUCTANCE	(\$)
	(OHM-M)	(M)	(M)	(FEET)	LAYER	TOTAL
			12.2	40.0		
	50.36	87.6	-75.4	-247.3	1.7	1.7
	14.19					
	TIMES	DATA	CALC	% ERROR	STD ERR	
1		5.46E+02				
2		4.24E+02				
3	1.40E-04	3.21E+02	3.15E+02	1.981		
4		2.46E+02				
5	2.20E-04	1.94E+02				
6	2.80E-04		1.55E+02			
7		1.20E+02				
8		9.80E+01				
9		8.08E+01				
10	7.13E-04	6.75E+01	6.70E+01	0.826		
11		5.75E+01				
12		4.99E+01				
13		4.32E+01		3.603		
14				4.045		
15		3.42E+01		4.404		
16		2.94E+01				
17		2.56E+01				
18		2.43E+01				
19		2.32E+01	2.30E+01	0.983		
20	7.13E-03	2.20E+01	2.18E+01	1.075		
21	8.81E-03	2.06E+01	2.07E+01	-0.497		
22	1.10E-02	1.96E+01	1.97E+01	-0.497		
23	1.41E-02	1.86E+01	1.89E+01	-1.604		
24	1.80E-02	1.76E+01	1.82E+01	-3.485		
25	2.22E-02	1.70E+01	1.77E+01	-4.029		
26	2.85E-02		1.72E+01	4.460		
27	3.60E-02	1.72E+01	1.68E+01	2.225		

03-01

R: 237. X: 0. Y: 237. DL: 475. REQ: 264. CF: 1.0000 TDHZ ARRAY, 27 DATA POINTS, RAMP: 190.0 MICROSEC, DATA: 03-01

	CIENCES, INC.
INVERSION TAI WEKIVA RIVER	
SITE 3	1
St. Johns River Manage	ment District
PROJECT NO: 91034	Table 4-4.1

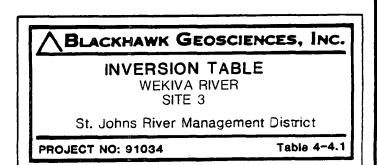
RMS LOG ERROR: 1.66E-02, ANTILOG YIELDS 3.8953 % LATE TIME PARAMETERS

\* Blackhawk Geosciences, Incorporated \*

PARAMETER RESOLUTION MATRIX: "F" MEANS FIXED PARAMETER P 1 1.00 P 2 0.00 1.00 T 1 0.00 0.00 1.00 P 1 P 2 T 1

PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

LAYER		MINIMUM	BEST	MAXIMUM
RHO	1	46.620	50.357	54.730
	2	13.612	14.189	14.831
THICK	1	79.931	87.560	94.326
DEPTH	1	79.931	87.560	94.326



# Summary of Results of TDEM Measurements at Wekiva River (Site 3)

From the TDEM measurements the following information about aquifer characteristics and water quality was derived:

- 1) The depth of occurrence of highly saline water was interpreted to occur at an elevation of 247 ft below msl. Below that elevation chloride concentrations were inferred to exceed 2,140 mg/l.
- 2) A resistivity value of 50 ohm-m was measured for the first 287 ft below surface. Since this layer will include the Hawthorn Group and younger sediments, as well as several hundred feet of the Upper Floridan aquifer, no information about chloride concentration was inferred above the interface with highly saline water.

### 4.5 SIMPSON TRAINING CENTER (SITE 4)

# Location and Geoelectric Section

The detailed location of this sounding is shown in Figure 4-5.1. Two soundings were made at this location to assure that the readings were not influenced by cultural interferences. One sounding employed a 1,200 ft by 1,200 ft loop, and one sounding was made with a 500 ft by 500 ft loop. In Figure 4-5.2 the two sounding curves are superimposed. The two curves merge at later times and diverge at earlier time, and this divergence is due to differences in geometric factors caused by different loop sizes. The two data sets were jointly inverted to the three layer geoelectric section shown on Figure 4-5.2.

### Geologic Interpretation of Geoelectric Section

The upper layer in the geoelectric section is interpreted to correspond to the Hawthorn Group. Its thickness of 135 ft derived locally at the TDEM stations compares with a regional thickness of 100 ft on the map of Tibbals (1990). The resistivities of the second (466 ohm-m) and third (6.8 ohm-m) layer represent different strata in the carbonate rocks of the Floridan aquifer.

## Depth of Occurrence of Highly Saline Water

The lowest resistivity encountered within the effective exploration depth of the measurement is 6.8 ohm-m, and it is first encountered at a depth of 639 ft below surface, or at an elevation of 559 ft below msl. Assuming an average porosity of 25%, the chloride concentration of ground water below that depth is expected to exceed 4,580 mg/l.

### Depth of Occurrence of 250 mg/l Isochlor

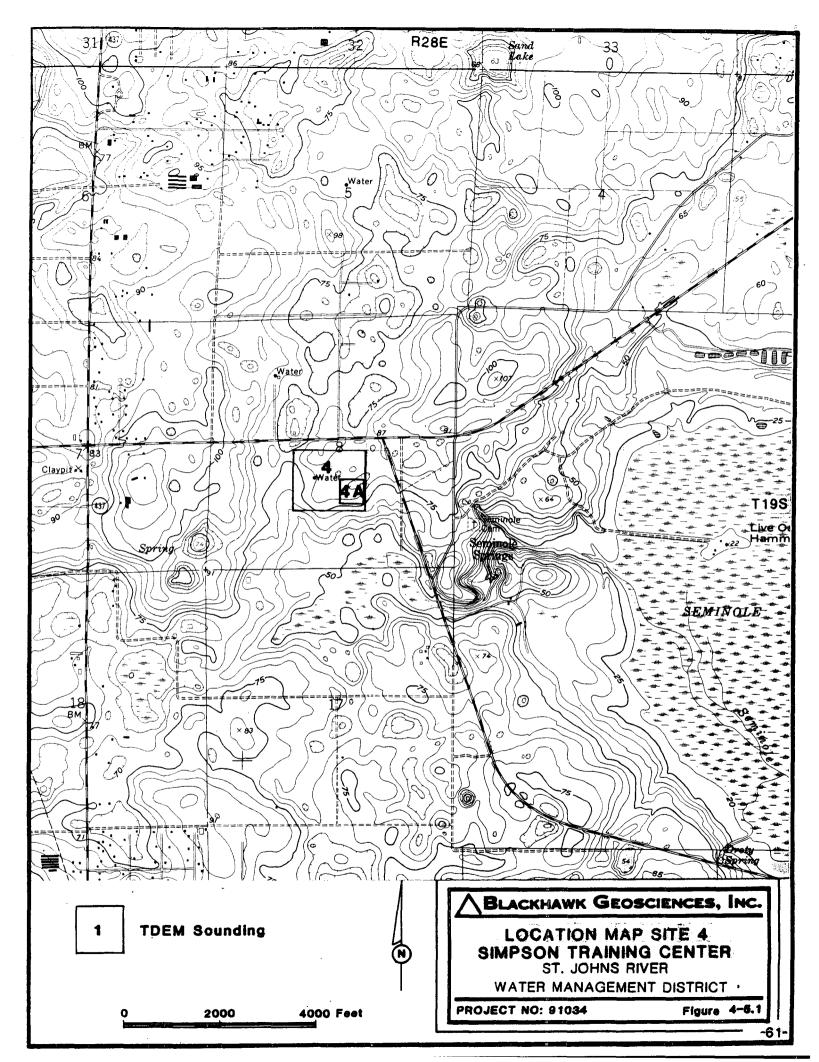
The second layer with a resistivity of 466 ohm-m between 55 ft and 559 ft below msl corresponds to a chloride concentration less than 100 mg/l assuming an average porosity of 25% and validity of the relation displayed on Figure 2-2. The occurrence of the 250 mg/l isochlor is expected to approximately coincide with the elevation of the boundary between the layer of 466 ohm-m and 6.8 ohm-m interpreted at 559 ft below msl. Very likely there will be a transition zone from ground water of low to high chloride concentrations. The thickness of such zones are not measured by TDEM.

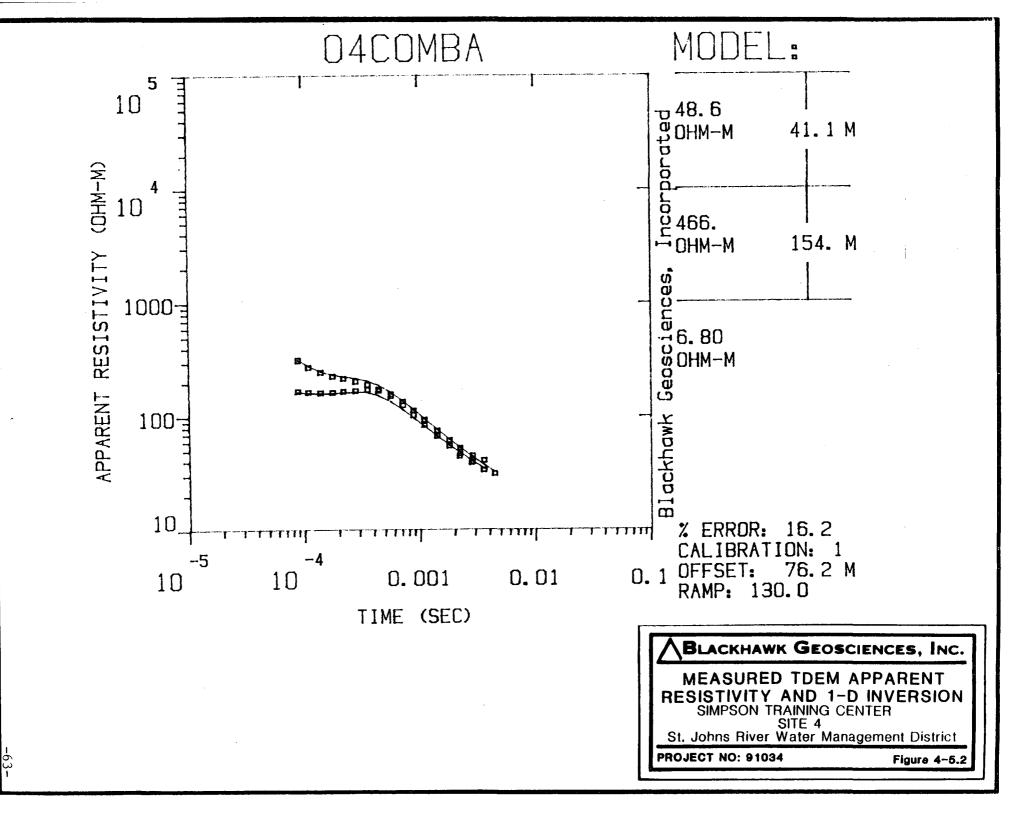
### Accuracy of Measurement and Interpretation

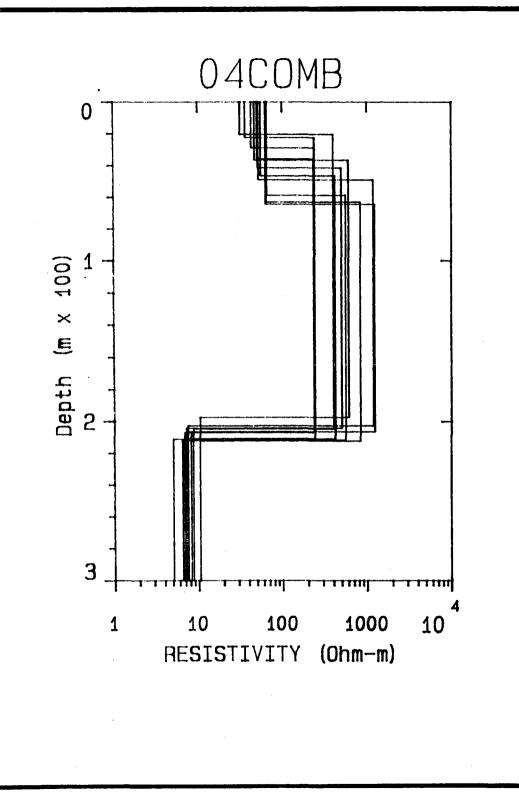
Figure 4-5.3 shows the computation of equivalence for the TDEM sounding at this site and the inversion table (Table 4-5.1) lists the upper and lower boundary of the parameters of the geoelectric section. The important parameters influencing information inferred about chloride concentration is the resistivity of the second layer. The range of equivalence in the resistivity of that layer is between 190 ohm-m and 2569 ohm-m. That variation, although large in resistivity, does not negate the conclusion about water quality within the interval between elevations of 55 ft and 559 ft below msl. Using an average porosity of 25% and the relation in Figure 2-2, chloride concentrations less than 250 mg/l are predicted for resistivities greater than 80 ohm-m.

### <u>Summary of Results of TDEM Measurements at the Simpson Training Center</u> (Site 4)

From the TDEM measurements the following information about aquifer characteristics and water quality was derived:







BLACKHAWK GEOSCIENCES, INC. EQUIVALENCE IN 1-D INVERSION SIMPSON TRAINING CENTER SITE 4 St. Johns River Water Management District PROJECT NO: 91034 Figure 4-5.3

# 04COMBA

MODEL: 3 LAYERS

RE	SISTIVITY	THICKNESS	ELEVAT	ION	CONDUCTANCE	(\$)
	(OHM-M)	(M)	(M)	(FEET)	LAYER	TOTAL
			24.4	80.0		
	48.63	41.1	-16.7	-54.8	0.8	0.8
4	65.58	153.6	-170.3	-558.8	0.3	1.2
	6.80					
	TIMES	DATA	CALC	% ERROR	STD ERR	
1				1.006		
2		1.65E+02	1.62E+02			
3	1.40E-04	1.63E+02	1.60E+02	1.907		
4	1.77E-04	1.64E+02		1.429		
5	2.20E-04		1.65E+02	1.640		
6	2.80E-04	1.70E+02	1.67E+02	1.525		
7	3.55E-04	1.74E+02	1.64E+02	6.096		
8	4.43E-04	1.70E+02	1.53E+02	10.918		
9	5.64E-04	1.58E+02	1.35E+02	16.708		
10	7.13E-04	1.36E+02	1.15E+02	18.213		
11	8.81E-04	1.13E+02	9.74E+01	15.674		
12	1.10E-03	9.30E+01	8.17E+01	13.878		
13	1.41E-03	7.48E+01	6.68E+01	12.112		
14	1.80E-03	6.13E+01	5.54E+01	10.561		
15	2.22E-03	5.23E+01	4.73E+01	10.567		
16	2.85E-03	4.46E+01	3.98E+01	12.096		
17	3.57E-03	4.05E+01	3.43E+01	18.329		
18	8.90E-05	3.19E+02	3.20E+02	-0.388		
19	1.10E-04	2.74E+02	2.80E+02	-2.129		
20	1.40E-04	2.48E+02	2.52E+02	-1.357		
21	1.77E-04	2.29E+02	2.35E+02	-2.859		
22	2.20E-04	2.19E+02	2.26E+02	-3.456		
23	2.80E-04	2.05E+02	2.18E+02	-6.027		
24	3.55E-04	1.92E+02	2.07E+02	-7.117		
25	4.43E-04	1.74E+02	1.89E+02	-8.154		
26	5.64E-04	1.52E+02	1.64E+02	-7.311		
27	7.13E-04	1.25E+02	1.38E+02	-8.860		
28	8.81E-04	1.02E+02	1.15E+02	-11.224		
2 <b>9</b>	1.10E-03	8.35E+01	9.51E+01	-12.237		
						<u>,                                    </u>

BLACKHAWK GEOSCIENCES, INC. INVERSION TABLE SIMPSON TRAINING CENTER SITE 4 St. Johns River Management District PROJECT NO: 91034 Table 4-5.1

30	1.41E-03	6.68E+01	7.64E+01	-12.567
31	1.80E-03	5.46E+01	6.25E+01	-12.579
32	2.20E-03	4.41E+01	5.30E+01	-16.790
33	2.22E-03	4.66E+01	5.26E+01	-11.383
34	2.80E-03	3.87E+01	4.41E+01	-12.294
35	2.85E-03	3.94E+01	4.35E+01	-9.456
36	3.55E-03	3.34E+01	3.72E+01	-10.186
37	4.43E-03	3.09E+01	3.20E+01	-3.494

R: 76. X: 0. Y: 76. DL: 152. REQ: 84. CF: 1.0000 R: 185. X: 0. Y: 185. DL: 369. REQ: 0. CF: 1.0000 TDHZ ARRAY, 17 DATA POINTS, RAMP: 130.0 MICROSEC, DATA: 04COMBA TDHZ ARRAY, 20 DATA POINTS, RAMP: 185.0 MICROSEC, DATA: 04-01A SIMPSONS TRAINING CENTER LOOP 2

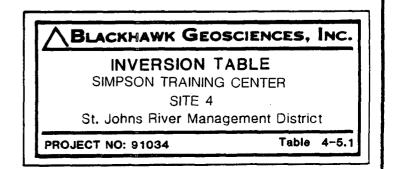
RMS LOG ERROR: 6.51E-02, ANTILOG YIELDS 16.1742 % LATE TIME PARAMETERS

\* Blackhawk Geosciences, Incorporated \*

PARAMETER RESOLUTION MATRIX: "F" MEANS FIXED PARAMETER P 1 0.98 P 2 -0.06 0.14 P 3 0.01 -0.01 0.98 T 1 -0.04 -0.19 0.01 0.92 T 2 0.01 0.06 0.00 0.02 0.99 P 1 P 2 P 3 T 1 T 2

PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

LAYER		MINIMUM BEST		MAXIMUM
RHO	1	35.934	48.635	68.908
	2	189.780	465.575	2568.555
	3	4.054	6.803	11.407
THICK	1	23.827	41.085	77.377
	2	129.939	153.636	174.971
DEPTH	1	23.827	41.085	77.377
	2	185.469	194.721	209.405



- 1) The depth of occurrence of saline water was interpreted at a depth of 559 ft below msl, and below that depth chloride concentration in excess of 4,580 mg/l are expected.
- 2) The interface between boundaries of different resistivities at 559 ft below msl was also inferred to approximately represent the 250 mg/l isochlor. Between elevations of 55 ft and 559 ft below msl, chloride concentrations less than 250 mg/l were predicted.

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## 4.6 ROYAL TRAILS (SITE 5)

### Location and Geoelectric Section

The detailed location map of this sounding is shown in Figure 4-6.1. The sounding data and the inverted geoelectric section is shown in Figure 4-6.2 and consists of a four layer section. Interference by buried telephone lines along at least two sides of the transmitter loop influenced data quality in earlier time gates, as evidenced by some scatter in the data at times between  $10^{-4}$  sec and  $10^{-3}$  sec. At times greater than  $10^{-3}$  sec little or no scatter in the data is observed.

#### Geologic Interpretation of Geoelectric Section

The first layer in the geoelectric section is 295 ft thick and has a resistivity of 78 ohm-m. Very likely this layer includes the Hawthorn Group and younger sediments as well as a portion of carbonate rocks of the Upper Floridan aquifer. The second, third, and fourth layers represent different horizons within the carbonate rocks of the Floridan aquifer.

#### Depth of Occurrence of Saline Water

The lowest resistivity encountered within the effective exploration depth of the measurement is 3.3 ohm-m and it is first encountered at a depth of 1,465 ft below surface, or at an elevation of 1,410 ft below msl. Assuming an average porosity of 25%, the chloride concentration of ground water at that depth is expected to exceed 9,590 mg/l. The geoelectric section shows this layer to be 144 ft thick, and the resistivity to increase below this layer. The most probable interpretation of an increase in resistivity at an elevation of 1,554 ft below msl is a decrease in average porosity of the carbonate rocks.

### Depth of Occurrence of 250 mg/l Isochlor

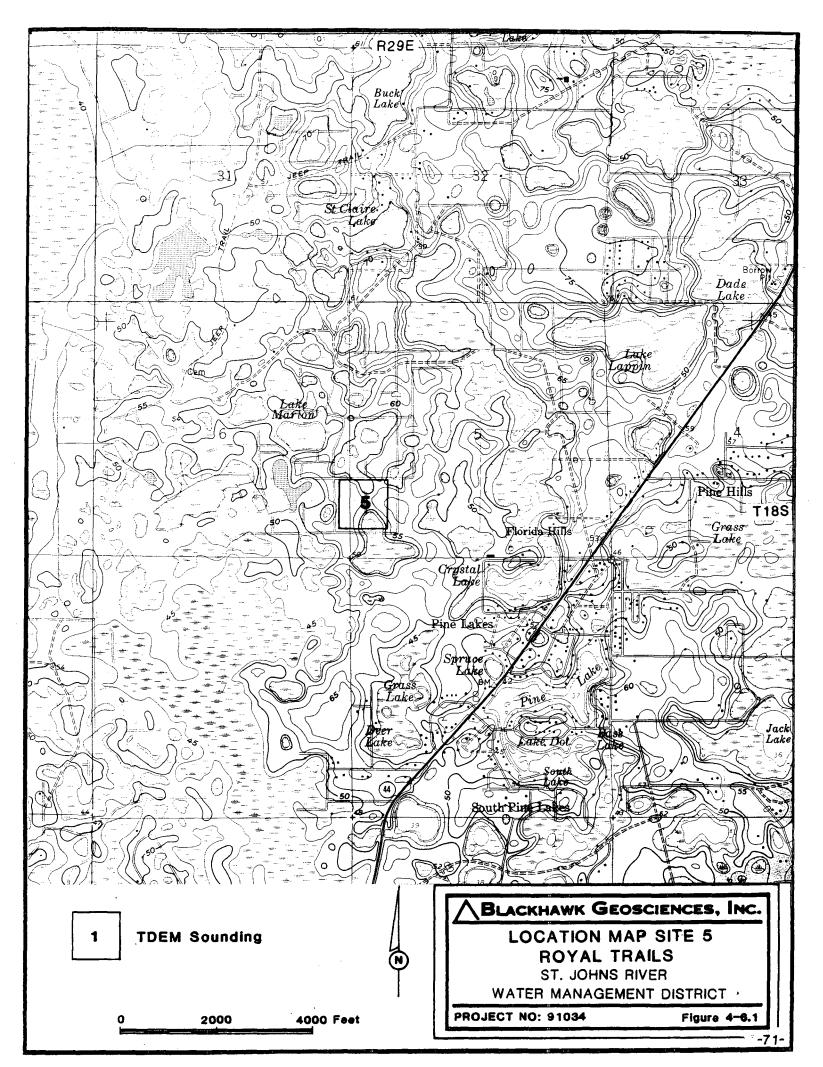
The resistivities of the first two layers is 78 ohm-m and 241 ohm-m. The first layer is 295 ft thick and likely includes the Hawthorn Group and likely 100 ft to 200 ft of the Upper Floridan aquifer. The resistivity of 78 ohm-m may be influenced by some clay layers in the Hawthorn Group which would lower its resistivity. The second layer with a resistivity of 241 ohm-m, occurring between elevations of 240 ft below msl and 1,410 ft below msl, represents the major section of the Floridan aquifer and the water quality in that section is inferred to be below 100 mg/l chloride. The depth of occurrence of the 250 mg/l is expected to coincide with the resistivity boundary (between 241 ohm-m and 3.3 ohm-m) measured at an elevation of 1,410 ft below msl.

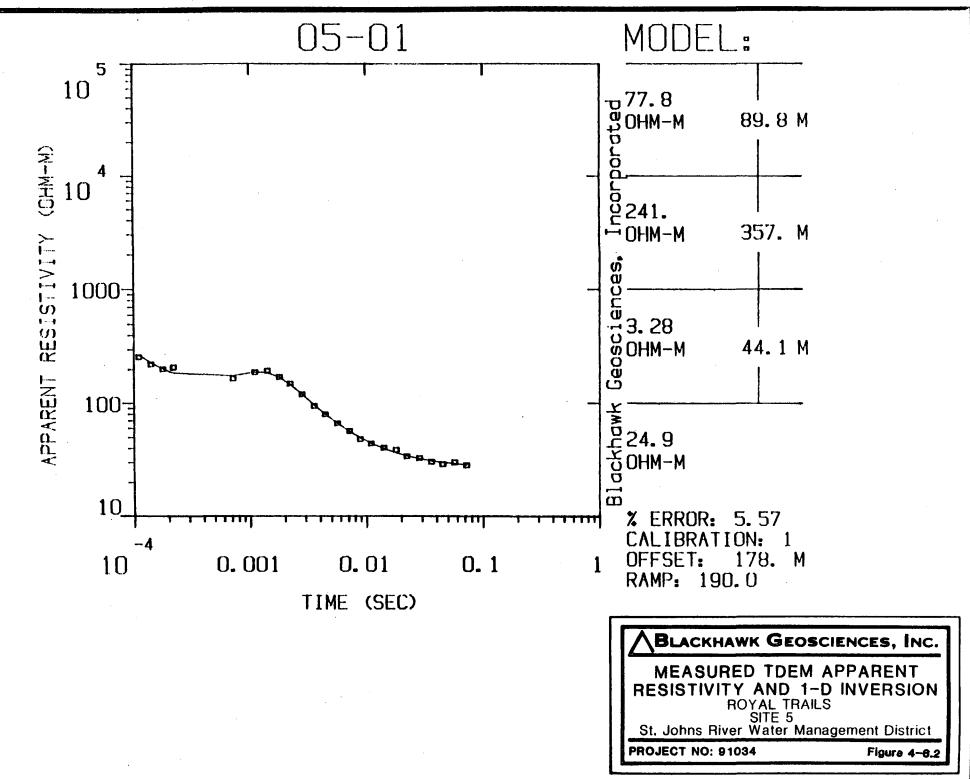
#### Accuracy of Measurement and Interpretation

Figure 4-6.3 shows the evaluation of equivalence for the TDEM sounding at this site, and the inversion table (Table 4-6.1) lists the upper and lower boundary of the parameters of the geoelectric section.

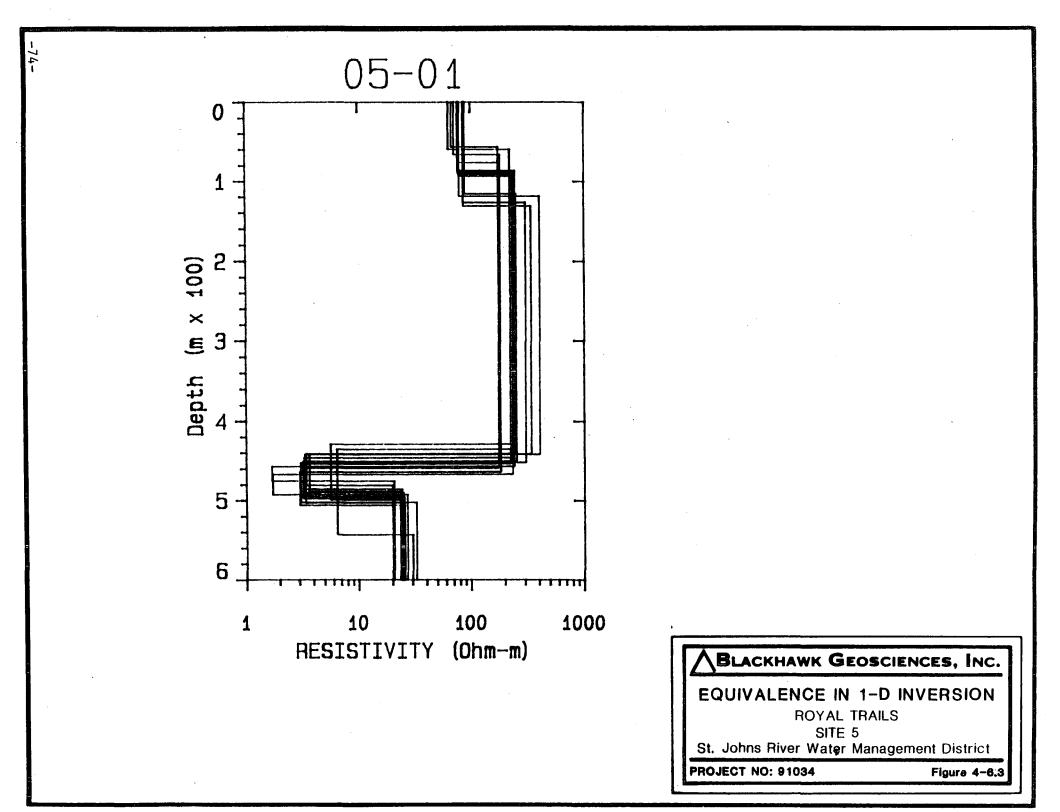
The range of equivalence in determining the depth to the layer of low resistivity (3.3 ohm-m) can be seen on Figure 4-6.3 to be about  $\pm 40 \text{ m}$  (130 ft) or about 10% of total depth. The upper and lower bound in the resistivity of this layer is between 1.7 ohm-m and 6.4 ohm-m, and although large, does not negate the conclusion that this layer represents carbonate rock of high salinity. At a resistivity of 6.4 ohm-m chloride concentrations greater than 4,800 mg/l are expected.

The equivalence in the resistivity of the second layer, representing the major section of the Floridan aquifer, is between 176 ohm-m and 406 ohm-m. Over this range of resistivities, chloride concentrations less than 250 mg/l





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05-01

MODEL: 4 LAYERS

RESISTIVITY	THICKNESS	ELEVATION		CONDUCTANCE	(S)	
(OHM-M)	(M)	(M)	(FEET)	LAYER	TOTAL	
		16.8	55.0			
77.80	89.8	-73.0	-239.7	1.2	1.2	
240.60	356.7	-429.7	-1409.9	1.5	2.6	
3.28	44.1	-473.8	-1554.4	13.4	16.1	
24.95						

TIMES		DATA	CALC	% ERROR	STD ERR
1	1.10E-04	2.57E+02	2.70F+02	-5.032	
2	1.40E-04				
3	1.77E-04				
4	2.20E-04				
	•				
5	7.13E-04				
6	1.10E-03	1.87E+02	1.89E+02	-0.577	
7	1.41E-03	1.93E+02	1.86E+02	3.960	
8	1.78E-03	1.69E+02	1.69E+02	0.051	
9	2.21E-03	1.48E+02	1.46E+02	1.584	·
10	2.80E-03	1.19E+02	1.19E+02	-0.069	
11	3.55E-03	9.38E+01	9.61E+01	-2.412	
12	4.43E-03	7.90E+01	7.95E+01	-0.701	
13	5.64E-03	6.57E+01	6.59E+01	-0.185	
14	7.13E-03	5.63E+01	5.60E+01	0.509	
15	8.81E-03	4.78E+01	4.93E+01	-3.143	
16	1.10E-02	4.38E+01	4.41E+01	-0.707	
17	1.41E-02	4.00E+01	3.95E+01	1.196	
18	1.80E-02	3.85E+01	3.63E+01	6.042	
19	2.22E-02	3.37E+01	3.41E+01	-1.088	
20	2.85E-02	3.25E+01	3.22E+01	1.094	
21	3.60E-02	3.02E+01	3.08E+01	-1.930	
22	4.49E-02	2.88E+01	2.98E+01	-3.442	
23	5.70E-02	2.99E+01	2.89E+01	3.225	
24	7.19E-02	2.81E+01	2.83E+01	-0.824	

R: 178. X: 0. Y: 178. DL: 356. REQ: 198. CF: 1.0000 TDHZ ARRAY, 24 DATA POINTS, RAMP: 190.0 MICROSEC, DATA: 05-01 ROYAL TRAILS



RMS LOG ERROR: 2.36E-02, ANTILOG YIELDS 5.5747 % LATE TIME PARAMETERS

\* Blackhawk Geosciences, Incorporated \*

PARAMETER RESOLUTION MATRIX:

-

"F"	MEANS	FIXED	PARAM	TER			
Р 1	0.93						
P 2	-0.03	0.36					
P 3	0.01	-0.02	0.49				
Р4	-0.01	0.04	0.06	0.74			
т 1	-0.15	-0.35	0.04	-0.02	0.50		
τ2	0.04	0.12	0.04	-0.01	0.13	0.95	
т3	0.00	0.00	-0.44	-0.12	-0.01	0.02	0.41
	P 1	I P 2	2 P 3	5 P 4	• т 1	T 2	Т 3

# PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

LAYER		MINIMUM	BEST	MAXIMUM
RHO	1	63.813	77.801	88.385
	2	175.784	240.598	406.399
	3	1.700	3.283	6.413
	4	20.082	24.946	32.670
THICK	1	56.979	89.810	131.813
	2	310.000	356.679	402.312
	3	18.314	44.064	107.642
DEPTH	1	56.979	89.810	131.813
	2	429.396	446.490	467.501
	3	475.793	490.554	543.355



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are predicted assuming 25% porosity and validity of the relation shown on Figure 2-2.

# Summary of Results of TDEM Measurements at Royal Trails (Site 5)

From the TDEM measurements the following information about water quality was derived:

- The depth of occurrence of saline water was interpreted at an elevation of 1,410 ft below msl, and below that depth chloride concentrations in excess of 9,590 mg/l are inferred.
- 2) This interface at 1,410 ft below msl is also inferred to be the elevation of occurrence of the 250 mg/l isochlor. The main aquifer section between elevations of 240 ft and 1,410 ft below msl is expected to contain ground water with chloride concentrations less than 250 mg/l chloride.

### 4.7 ORMOND BEACH - WEST WELL FIELD (SITE 6)

### Location and Geoelectric Section

The detailed location map of this sounding is shown in Figure 4-7.1. The TDEM data and the inverted geoelectric section is shown in Figure 4-7.2 and consists of a three layer section.

#### Geologic Interpretation of Geoelectric Section

Well MW-2 is located approximately 1.5 to 2 miles west of sounding 6, and in Figure 4-7.3 the lithologic drilling log and two geophysical logs run in the well are compared with the geoelectric section derived from the TDEM sounding. The comparison provides a key for inferring geology from the geoelectric section assuming that the geology can be interpolated over a distance of about two miles:

- The upper layer of 39 ohm-m and 76 ft thick corresponds to sands and clays of surficial sediments overlaying the carbonate rocks.
- 2) The resistance log in the borehole shows no change in trend below the casing from about 80 ft to the depth of completion of the hole (290 ft). This is consistent with the geoelectric section derived from TDEM which indicates a resistivity of 98 ohm-m from a depth of about 76 ft to 1,122 ft below surface.

#### Depth of Occurrence of Saline Water

The lowest resistivity encountered within the effective depth of the measurement is 2.3 ohm-m, and it is first encountered at a depth of 1,122 ft below surface or at an elevation of 1,087 ft below msl. Assuming an average porosity of 25% the chloride concentration at that depth is expected to exceed 10,000 mg/l.

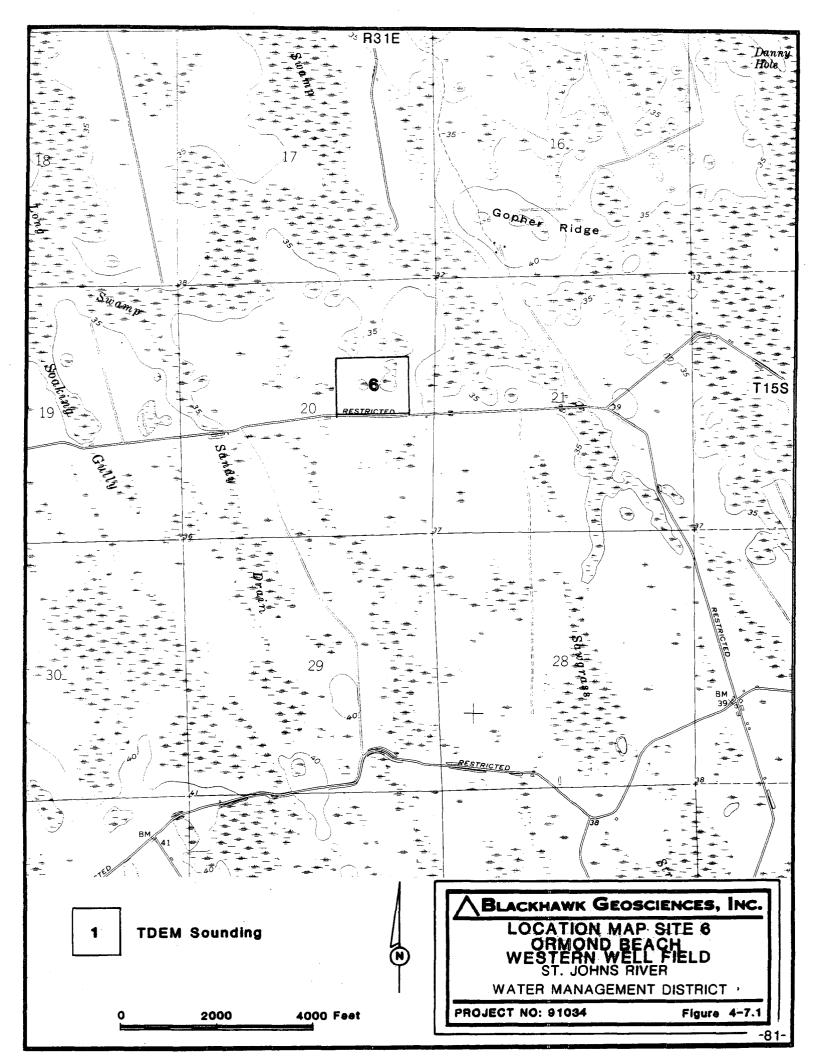
# Depth of Occurrence of 250 mg/l Isochlor

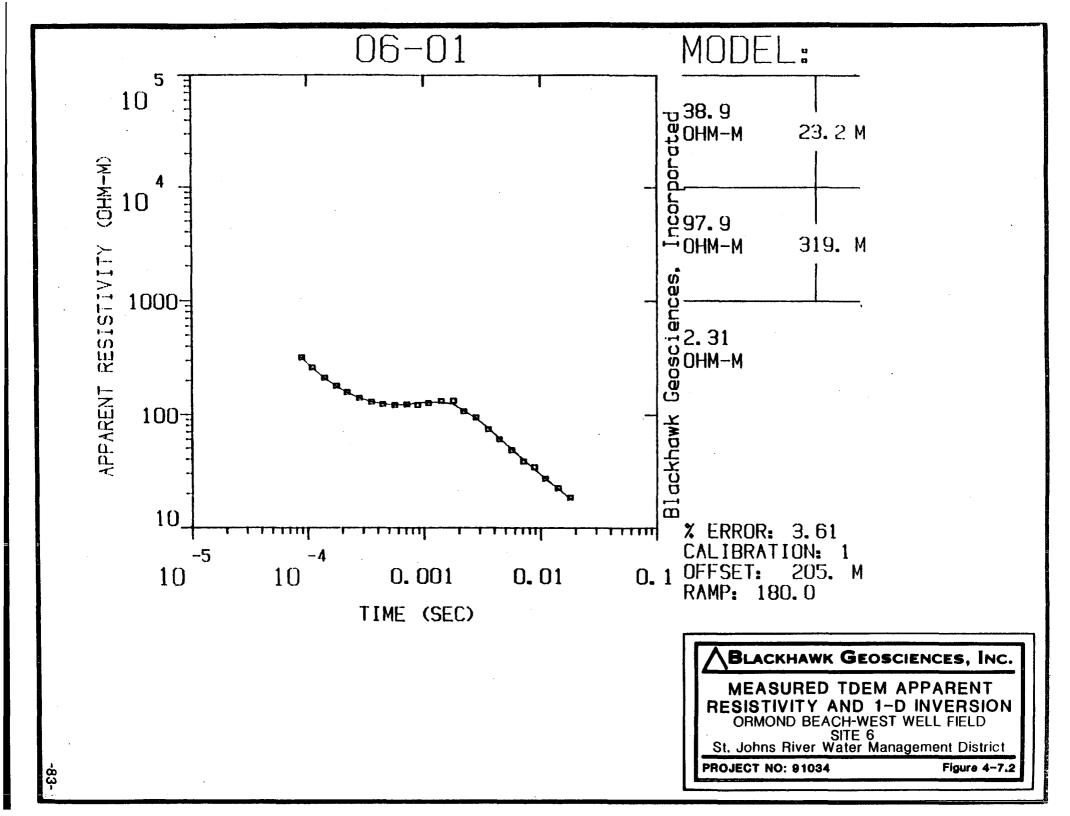
The resistivity of the carbonate rocks of the Floridan aquifer below the surficial sediments interpreted from the TDEM data is 98 ohm-m. The resistivity of this layer characterizes the carbonate rocks between an elevation of about 41 ft below msl to 1,087 ft below msl. Assuming an average porosity of 25% a chloride concentration less than 250 mg/l is inferred from the geoelectric section. This is consistent with information published by Tibbals (1990) which shows the upper Floridan to contain less than 250 ppm chloride at this location. The depth of occurrence of the 250 ppm chloride is expected to coincide with the resistivity boundary mapped at an elevation of 1,087 ft below msl.

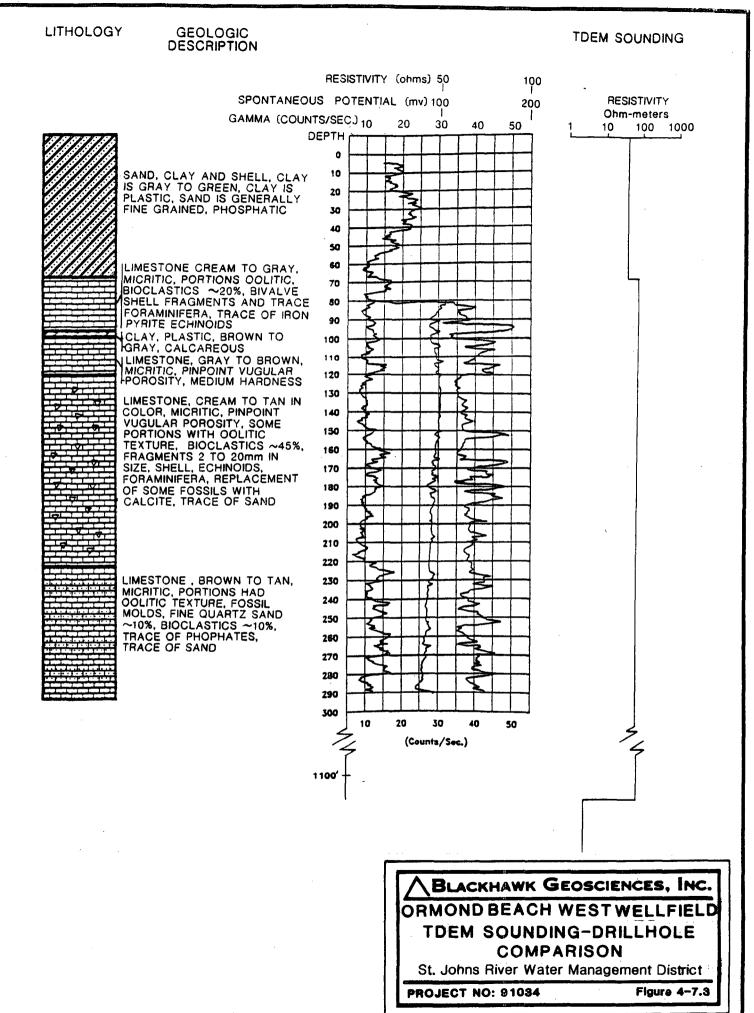
### Accuracy of Measurement and Interpretation

Figure 4-7.4 shows the evaluation of equivalence of the TDEM sounding at this site and the inversion table (Table 4-7.1) lists the upper and lower bounds of the parameters of the geoelectric section. The range of equivalence indicated for the depth of the highly saline layer is less than 2% of total depth, or about  $\pm$  20 ft. The resistivity range of the second layer, characteristic of the major section of the Floridan aquifer, may vary between 92 ohm-m and 109 ohm-m. Variation within this range does not negate the conclusion that the chloride concentration is below 250 mg/l.

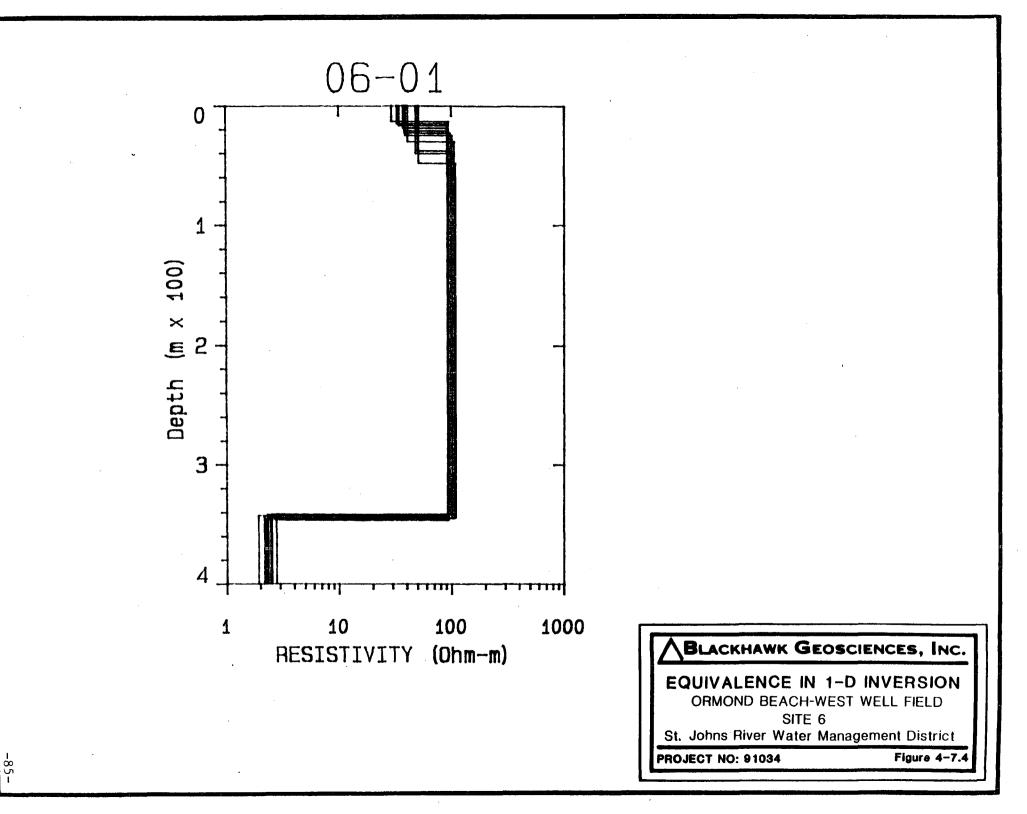
Very likely a transition zone where the chloride concentration increases gradually from less than 250 mg/l to in excess of 10,000 mg/l will occur above the interface determined at an elevation of 1,087 ft below msl. Such a transition zone is not mapped with TDEM and may be as large as 100 ft to 200 ft.







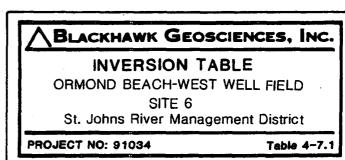
-84-



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MODEL: 3 LAYERS
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RE	SISTIVITY	THICKNESS	ELEVAT	ION	CONDUCTANCE	(S)
	(OHM-M)	(M)	(M)	(FEET)	LAYER	TOTAL
			10.7	35.0		
	38.89	23.2	-12.5	-41.2	0.6	0.6
	97.89	318.7	-331.3 -	1086.8	3.3	3.9
	2.31					
	TIMES	DATA	CALC	% ERROR	STD ERR	
1	8.90E-05	3.20E+02	3.19E+02	0.269		
2	1.10E-04	2.61E+02	2.58E+02			
3	1.40E-04					
4	1.77E-04	1.79E+02	1.78E+02	0.712		
5	2.20E-04					
6	2.80E-04					
7	3.55E-04					
8	4.43E-04					
9	5.64E-04		1.21E+02			
10	7.13E-04					
11						
12	1.10E-03	1.26E+02				
13	1.41E-03					
14	1.80E-03					
15	2.20E-03					
16						
17			7.57E+01			
18						
19	5.64E-03					
20	7.13E-03					
21						
22						
23						
24	1.80E-02	1.84E+01	1.81E+01	1.627		

R: 205. X: 0. Y: 205. DL: 409. REQ: 228. CF: 1.0000 CLHZ ARRAY, 24 DATA POINTS, RAMP: 180.0 MICROSEC, DATA: 06-01 ORMOND BEACH WEST WELL FIELD



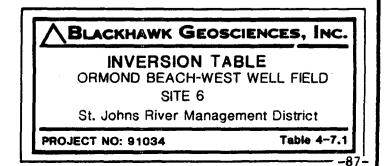
RMS LOG ERROR: 1.54E-02, ANTILOG YIELDS 3.6093 % LATE TIME PARAMETERS

\* Blackhawk Geosciences, Incorporated \*

PARAMETER RESOLUTION MATRIX: "F" MEANS FIXED PARAMETER P 1 1.00 P 2 0.00 1.00 P 3 0.00 0.00 1.00 T 1 0.00 0.00 0.00 0.99 T 2 0.00 0.00 0.00 0.00 1.00 . P 1 P 2 P 3 T 1 T 2

## PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

LAYER MINIMUM		MINIMUM	BEST	MAXIMUM
RHO	1	29.528	38.891	51.365
	2	92.173	97.886	109.125
	3	1.928	2.313	2.776
THICK	1	13.426	23.217	48.435
	2	296.246	318.721	333.004
DEPTH	1	13.426	23.217	48.435
	2	341.472	341.938	346.430



# Summary of Results of TDEM Measurements at Ormond Beach - West Well Field (Site 6)

From the TDEM measurements the following information about water quality was derived:

- 1) The depth of occurrence of saline water was interpreted at an elevation of 1,087 ft below msl, and below that depth chloride concentrations in excess of 10,000 mg/l chloride are expected.
- 2) The interface at 1,087 ft below msl is also inferred as the elevation of occurrence of the 250 mg/l isochlor. The main aquifer section between elevations of 41 ft below msl and 1,087 ft below msl is expected to contain ground water with chloride concentrations less than 250 mg/l.

#### 4.8 ASTRONAUT HIGH SCHOOL (SITE 7)

#### Location and Geoelectric Section

The detailed location map of this sounding is shown in Figure 4-8.1. The TDEM data and the inverted geoelectric section is shown in Figure 4-8.2and consists of a two layer section.

## <u>Comparison of Resistivity Log in a Nearby Well with TDEM Derived Geoelectric</u> <u>Section</u>

In Figure 4-8.3 the resistivity log in a nearby well is compared with the geoelectric section derived from TDEM surveys. The two data sets agree about the depth of occurrence (181 ft below surface) of a low resistivity layer in the section, but are in disagreement about absolute values. Above this depth the resistivity varies greatly in the resistivity log. The upper layer of about 30 ohm-m measured with the TDEM likely represents surficial sediments, the Hawthorn Group, and portions of the upper carbonate rocks of the Floridan aquifer. It is an average resistivity of all the various sand and clay layers encountered within the upper 181 ft.

#### Depth of Occurrence of Saline Water

The lowest resistivity encountered within the effective depth of the measurement is 2.5 ohm-m, and it is first encountered at a depth of 181 ft below surface or at an elevation of 156 ft below msl. Assuming an average porosity of 25% the chloride concentration below that depth is expected to exceed 10,000 mg/l.

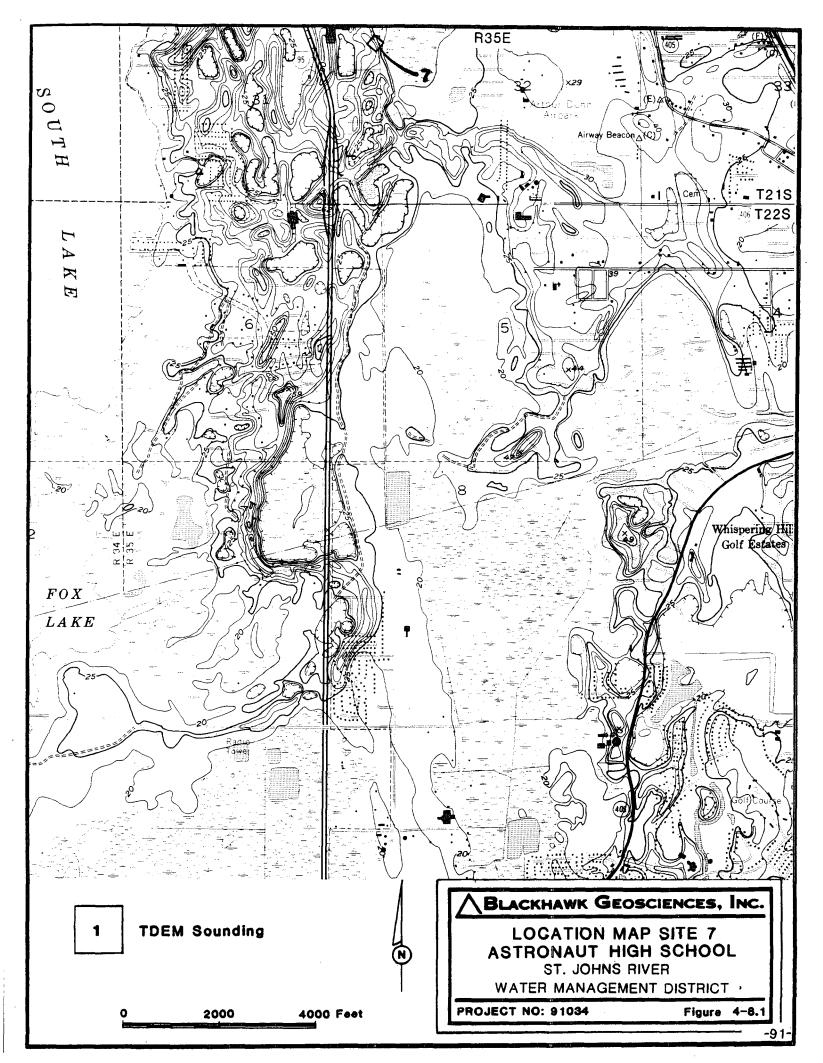
### Depth of Occurrence of 250 mg/l Isochlor

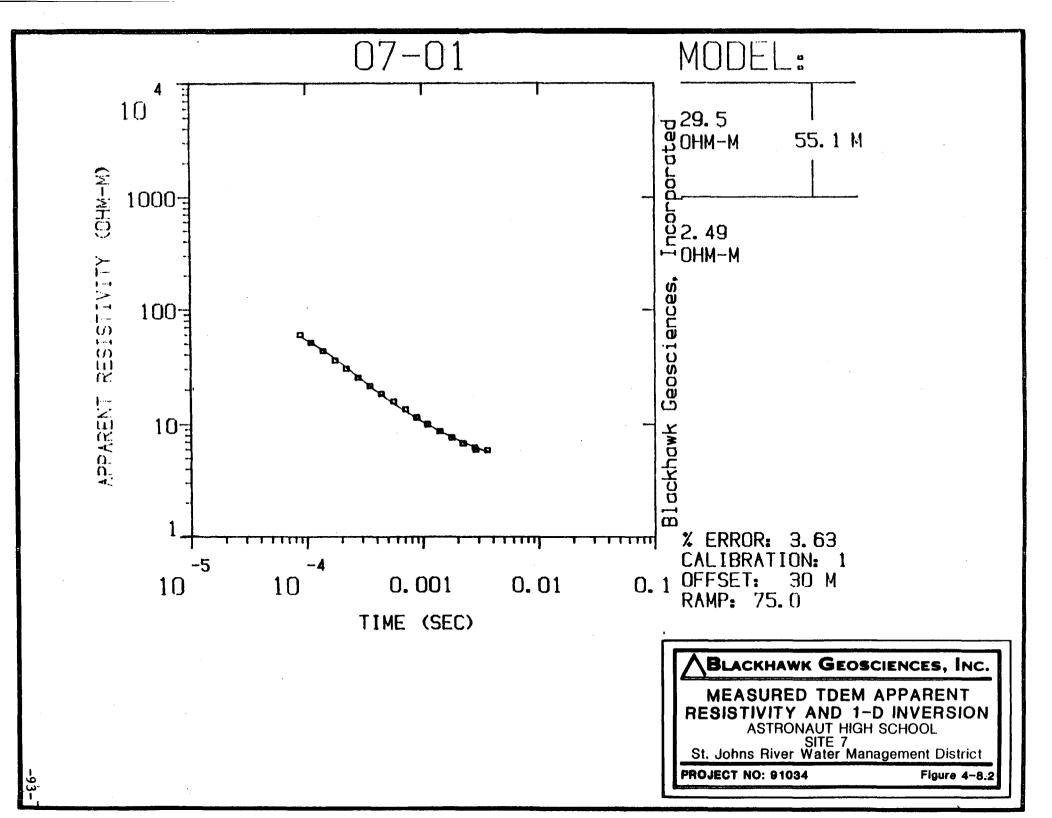
The resistivity of the upper 181 ft likely include portions of the carbonate rocks below the Hawthorn Group. However, since the upper 181 ft includes the Hawthorn Group and younger sediments, the resistivity of 30 ohm-m may be influenced by clay layers. Chloride concentration can, therefore, not be inferred for this layer.

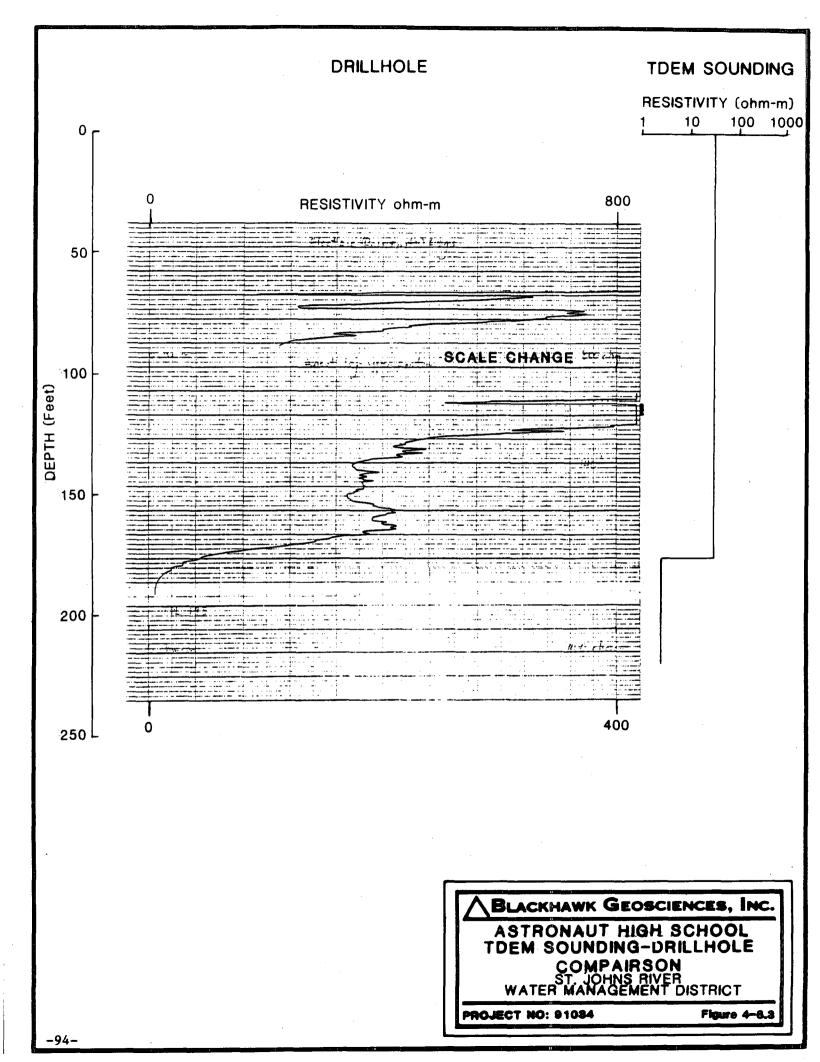
#### Accuracy of Measurement and Interpretation

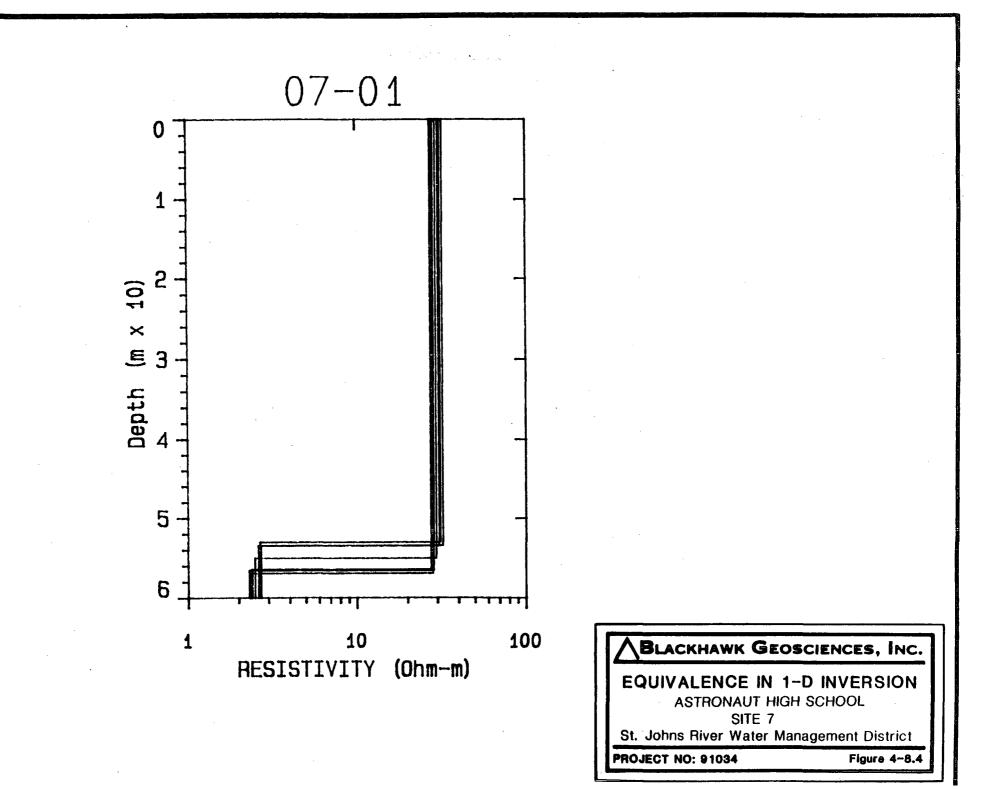
Figure 4-8.4 shows the evaluation of equivalence of the TDEM sounding at this site, and the inversion table (Table 4-8.1) lists the upper and lower bounds of the parameters of the geoelectric section. The range of equivalence of all parameters is relatively small. The error of measurement in determining the elevation of the interface with saline water is expected to be  $\pm$  5% of total depth, or about  $\pm$  10 ft.

The range of resistivity of the upper layer may vary between 27 ohm-m and 32 ohm-m. Since clay layers in the Hawthorn Group and younger sediments likely influence this range of resistivity measured, no conclusions about chloride concentration can be drawn.









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MODEL: 2 LAYERS
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R	ESISTIVITY T	HICKNESS	ELEVAT	ION	CONDUCTANCE	(S)
	(OHM-M)	(M)	(M)	(FEET)	LAYER	TOTAL
			7.6	25.0		
	29.53	55.1	-47.5	-155.7	1.9	1.9
	2.49					
	TIMES	DATA	CALC	% ERROR	STD ERR	
1	8.90E-05	5.98E+01	5.78E+01	3.300		
2	1.10E-04	5.09E+01	5.12E+01	-0.497		
3	1.40E-04	4.27E+01	4.36E+01	-2.021		
4	1.77E-04	3.54E+01	3.67E+01	-3.639		
5	2.20E-04	3.01E+01	3.10E+01	-2.869		
6	2.80E-04	2.51E+01	2.56E+01	-1.818		
7	3.55E-04	2.13E+01	2.13E+01	-0.112		
8	4.43E-04	1.83E+01	1.80E+01	1.604		
9	5.64E-04	1.58E+01	1.51E+01	<sup>~</sup> 4.104		
10	7.13E-04	1.35E+01	1.29E+01	4.285		
11	8.81E-04	1.15E+01	1.13E+01	1.631		
12	8.90E-04	1.14E+01	1.12E+01	1.440		
13	1.10E-03	1.00E+01	9.90E+00	1.211		
14	1.10E-03	9.88E+00	9.88E+00	-0.050		
15	1.40E-03	8.60E+00	8.64E+00	-0.495		
16	1.41E-03	8.64E+00	8.60E+00	0.451		
17	1.77E-03	7.63E+00	7.66E+00	-0.454		
18	1.80E-03	7.53E+00	7.61E+00	-1.053		
19	2.20E-03	6.75E+00	6.91E+00	-2.295		
20	2.22E-03	6.73E+00	6.88E+00	-2.151		
21	2.80E-03	6.20E+00	6.23E+00	-0.426		
22	2.85E-03	5.90E+00	6.18E+00	-4.485		
23	3.60E-03	5.86E+00	5.64E+00	3.797		

R: 30. X: 0. Y: 30. DL: 60. REQ: 33. CF: 1.0000 CLHZ ARRAY, 23 DATA POINTS, RAMP: 75.0 MICROSEC, DATA: 07-01 ASTRONAUT HIGH SCHOOL

RMS LOG ERROR: 1.55E-02, ANTILOG YIELDS LATE TIME PARAMETERS 3.6305 %



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\* Blackhawk Geosciences, Incorporated \*

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PARAMETER RESOLUTION MATRIX:

"F" MEANS FIXED PARAMETER

P 1 0.99

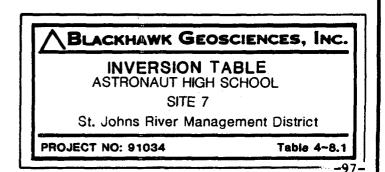
P 2 -0.01 0.99

T 1 0.00 0.00 1.00

P 1 P 2 T 1
```

PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

LAYER		MINIMUM	BEST	MAXIMUM
RHO	1	27.462	29.528	32.165
	2	2.305	2.490	2.704
THICK	1	53.037	55.074	56.953
DEPTH	1	53.037	55.074	56.953
		· · · ·		



## Summary of Results of TDEM Measurements at Astronaut High School (Site 7)

From the TDEM measurements the following information about water quality was derived:

- 1) The depth of occurrence of saline water was interpreted to occur at an elevation of 156 ft below msl, and below that depth chloride concentrations in excess of 10,000 mg/l chloride are expected.
- 2) Because no distinction could be made between the resistivity values of the Hawthorn Group, surficial sediments, and the carbonate rocks of the Upper Floridan, no information about chloride concentration could be inferred above the highly saline water interface.

## 4.9 NORTH BREVARD WASTE WATER TREATMENT PLANT (SITE 8)

## Location and Geoelectric Section

The detailed location map of this sounding is shown in Figure 4-9.1. The TDEM data and the inverted geoelectric section is shown in Figure 4-9.2and consists of a two layer section.

## Depth of Occurrence of Saline Water

The lowest resistivity encountered within the effective depth of the measurement is 1.6 ohm-m, and it is first encountered at a depth of 183 ft below surface or at an elevation of 168 ft below msl. Assuming an average porosity of 25% the chloride concentration below that depth is expected to exceed 10,000 mg/l.

#### Depth of Occurrence of 250 mg/l Isochlor

The resistivity of the upper 183 ft likely include portions of the carbonate rocks below the Hawthorn Group and younger sediments. Since the resistivity of 36 ohm-m may have been influenced by clay layers, no information about chloride concentration can be inferred for this upper layer.

#### Accuracy of Measurement and Interpretation

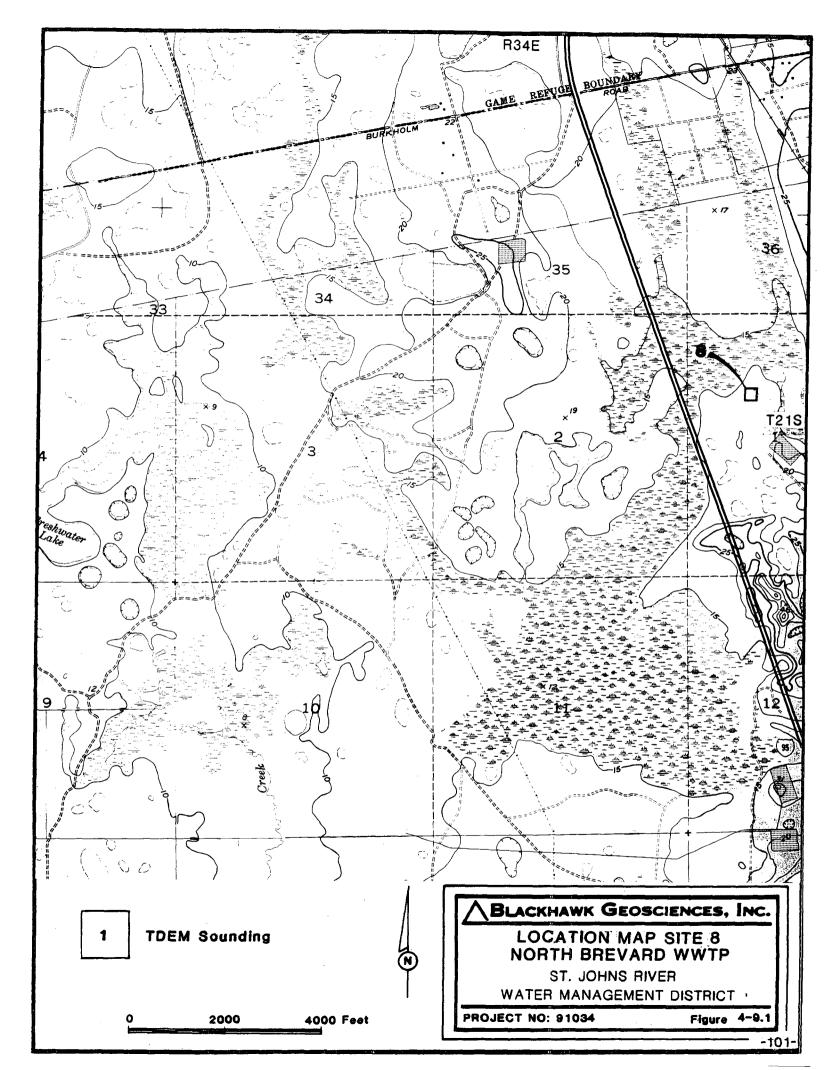
Figure 4-9.3 shows the evaluation of equivalence of the TDEM sounding at this site, and the inversion table (Table 4-9.1) lists the upper and lower bounds of the parameters of the geoelectric section. The range of equivalence of all parameters is relatively small. The error of measurement in determining the elevation of the interface with saline water is expected to be  $\pm$  5% of total depth, or about  $\pm$  10 ft.

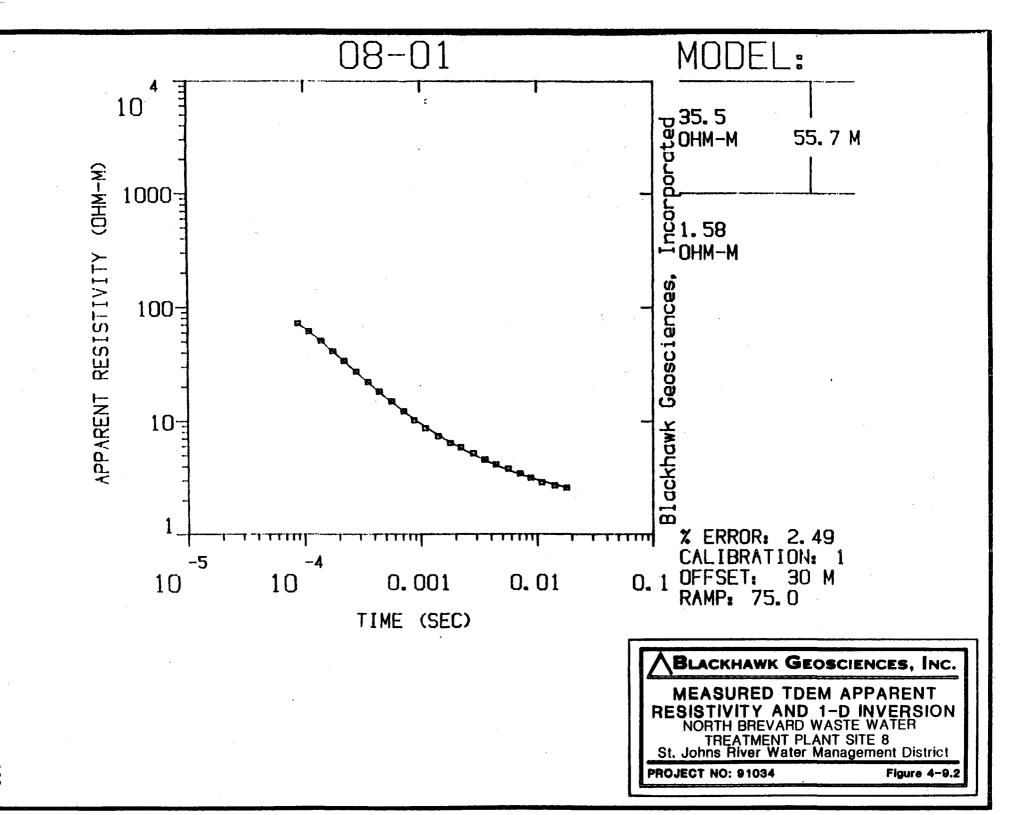
The range of resistivity of the upper layer may vary between 33 ohm-m and 38 ohm-m. Clay layers in the Hawthorn Group and younger sediments may be the cause of this value of resistivity, and hence no information about chloride concentration is derived for the upper 183 ft.

#### Summary of Results of TDEM Measurements at North Brevard Waste Water Treatment Plant (Site 8)

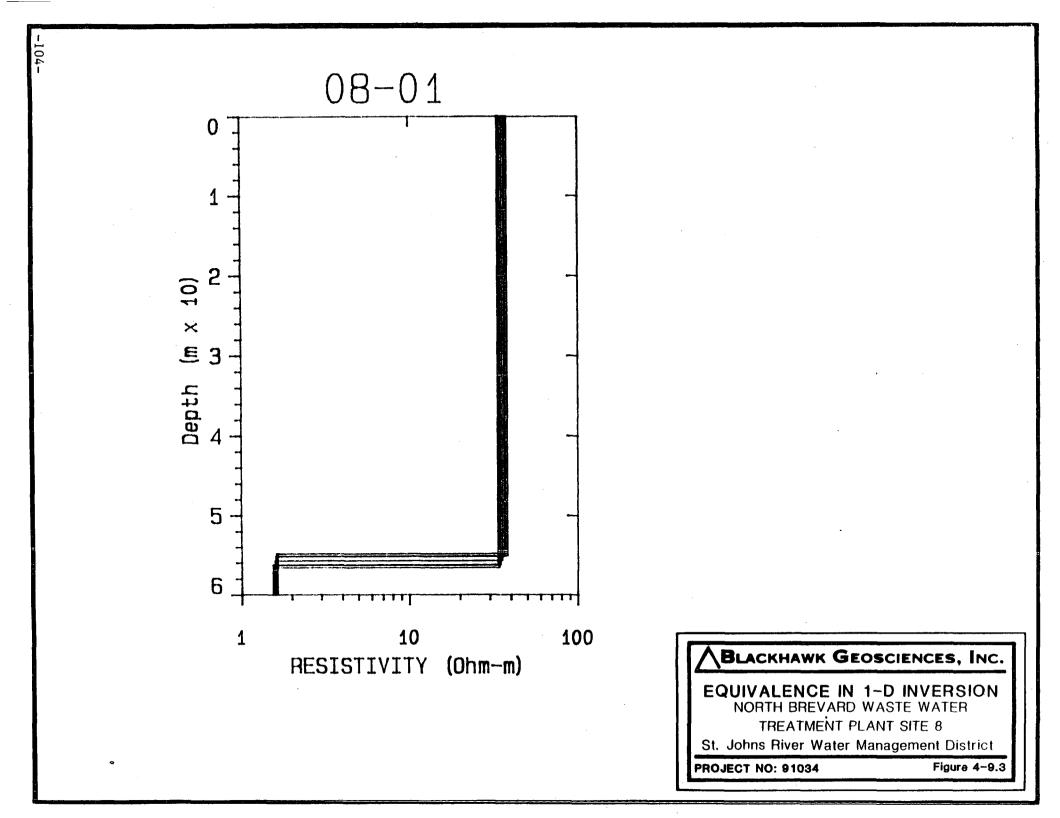
From the TDEM measurements the following information about water quality was derived:

- 1) The depth of occurrence of saline water was interpreted to occur at an elevation of 168 ft below msl, and below that depth chloride concentrations in excess of 10,000 mg/l chloride are expected.
- 2) Because no distinction could be made between the resistivity value of the Hawthorn Group, surficial sediments, and the carbonate rocks of the Upper Floridan, no information about chloride concentration could be inferred above the highly saline water interface.





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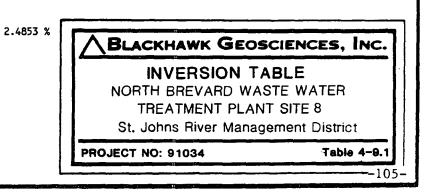


MODEL: 2 LAYERS

R	ESISTIVITY	THICKNESS	ELEVAT	ION	CONDUCTANCE	(S)
	(OHM-M)	(M)	(M)	(FEET)	LAYER	TOTAL
			4.6	15.0		
	35.50	55.7	-51.2	-167.9	1.6	1.6
	1.58					
•	TIMES	DATA	CALC	% ERROR	STD ERR	
1	8.90E-05	7.19E+01	7.16E+01	0.332		
2	1.10E-04	6.13E+01	6.17E+01	-0.655		
3	1.40E-04	5.06E+01	5.07E+01	-0.242		
4	1.77E-04	4.11E+01	4.12E+01	-0.042		
5	2.20E-04	3.40E+01	3.37E+01	0.748		
6	2.80E-04	2.72E+01	2.70E+01	0.731		
7	3.55E-04	2.20E+01	2.18E+01	0.827		
8	<b>4.43</b> E-04	1.81E+01	1.80E+01	0.851		
9	5.64E-04	1.49E+01	1.47E+01	1.054		
10	7.13E-04	1.21E+01	1.22E+01	-0.440		
11	8.81E-04	1.02E+01	1.04E+01	-2.464		
12	1.10E-03	8.63E+00	8.91E+00	-3,125		
13	1.41E-03	7.34E+00	7.54E+00	-2.533		
14	1.80E-03	6.38E+00	6.50E+00	-1.731		
15	2.21E-03	5.88E+00	5.76E+00	1.960		
16	2.83E-03	5.21E+00	5.06E+00	2.784		
17	3.55E-03	4.58E+00	4.53E+00	1.065		
18	4.43E-03	4.17E+00	4.11E+00	1.657		
19	5.64E-03	3.82E+00	3.72E+00	2.627		
20	7.13E-03	3.46E+00	3.41E+00	1.558		
21	8.81E-03	3.18E+00	3.18E+00	0.099		
22	1.10E-02	2.89E+00	2.97E+00	-2.561		
23	1.41E-02	2.72E+00	2.77E+00	-1.627		
24	1.80E-02	2.59E+00	2.61E+00	-0.544		

R: 30. X: .0. Y: 30. DL: 60. REQ: 33. CF: 1.0000 CLHZ ARRAY, 24 DATA POINTS, RAMP: 75.0 MICROSEC, DATA: 08-01 NORTH BREVARD WWTP

RMS LOG ERROR: 1.07E-02, ANTILOG YIELDS



## LATE TIME PARAMETERS

\* Blackhawk Geosciences, Incorporated \*

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PARAMETER RESOLUTION MATRIX:

"F" MEANS FIXED PARAMETER

P 1 1.00

P 2 0.00 1.00

T 1 0.00 0.00 1.00

P 1 P 2 T 1
```

PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

LAYER MI		MINIMUM	NIMUM BEST	
RHO	1	33.439	35.499	37.946
	2	1.531	1.580	1.633
THICK	1	54.849	55.749	56.632
DEPTH	1	54.849	55.749	56.632

## **BLACKHAWK GEOSCIENCES, INC.**

INVERSION TABLE

NORTH BREVARD WASTE WATER TREATMENT PLANT SITE 8

St. Johns River Management District

PROJECT NO: 91034

#### 4.10 PARRISH GROVE (SITE 9)

## Location and Geoelectric Section

The detailed location map of this sounding is shown in Figure 4-10.1. Two soundings were made at this site. The data in the first sounding was distorted due to a buried pipeline near one side of the transmitter loop. A second sounding was repositioned 200 ft away from the pipeline and undistorted data were obtained at that location. This is the location shown in Figure 4-10.1. The TDEM data and the inverted geoelectric section is shown in Figure 4-10.2 and consists of a two layer section.

## Depth of Occurrence of Saline Water

The lowest resistivity encountered within the effective depth of the measurement is 2.4 ohm-m, and it is first encountered at a depth of 212 ft below surface or at an elevation of 182 ft below msl. Assuming an average porosity of 25% the chloride concentration below that depth is expected to exceed 10,000 mg/l.

#### Depth of Occurrence of 250 mg/l Isochlor

The resistivity of the upper 212 ft likely include portions of the carbonate rocks below the Hawthorn Group and younger sediments. Since the resistivity of 31 ohm-m may have been influenced by clay layers, no conclusions about chloride concentrations can be drawn for the upper 212 ft.

#### Accuracy of Measurement and Interpretation

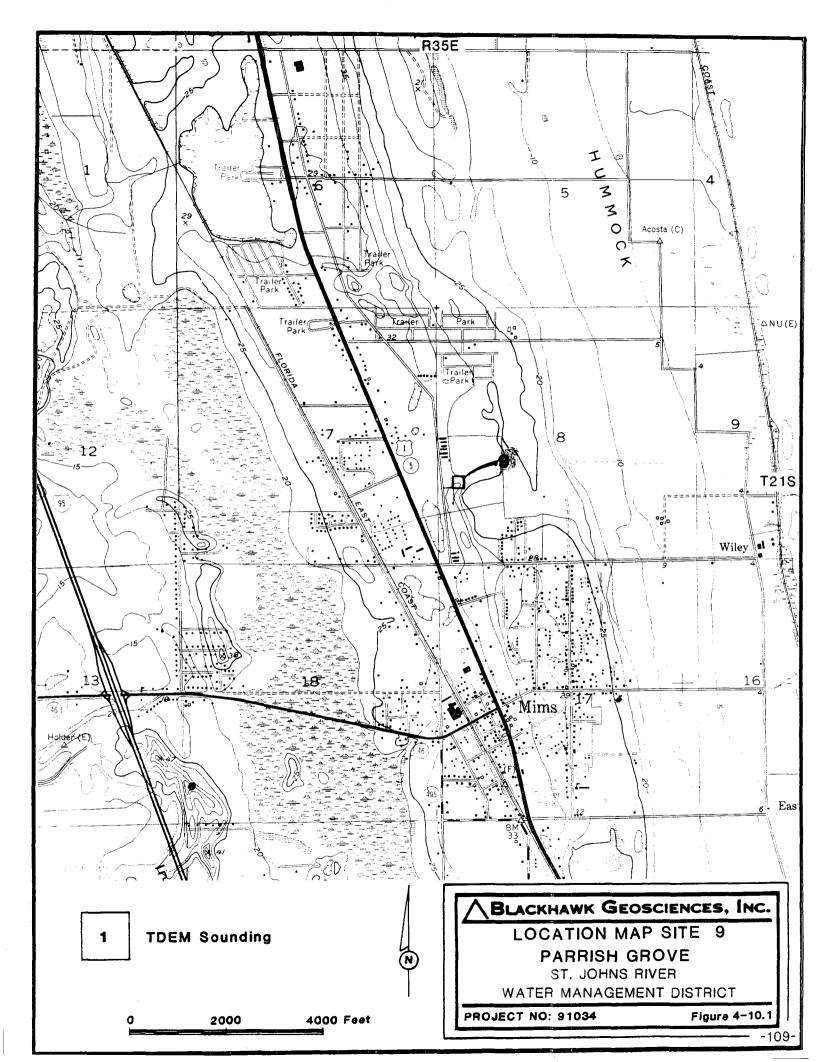
Figure 4-10.3 shows the evaluation of equivalence of the TDEM sounding at this site, and the inversion table (Table 4-10.1) lists the upper and lower bounds of the parameters of the geoelectric section. The range of equivalence of all parameters is relatively small. The error of measurement in determining the elevation of the interface with saline water is expected to be  $\pm$  5% of total depth, or about  $\pm$  10 ft.

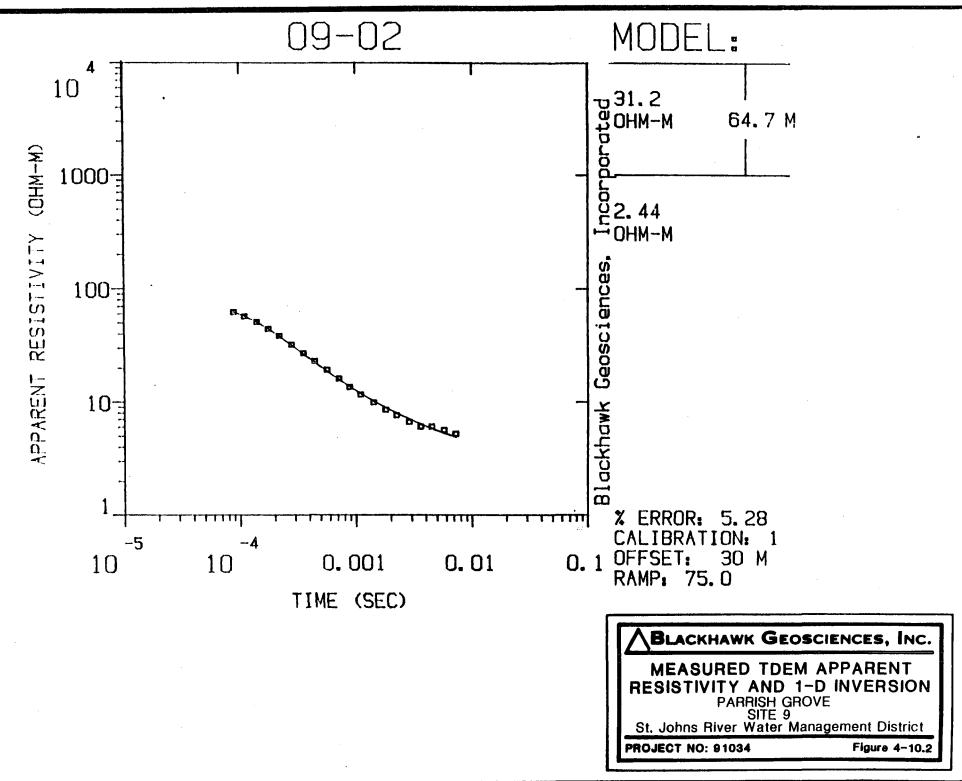
The range of resistivity of the upper layer may vary between 29 ohm-m and 33 ohm-m. Since clay layers in the Hawthorn Group and younger sediments may be the cause of this resistivity value, no information about chloride concentrations were inferred for the upper 212 ft.

#### Summary of Results of TDEM Measurements at Parrish Grove (Site 9)

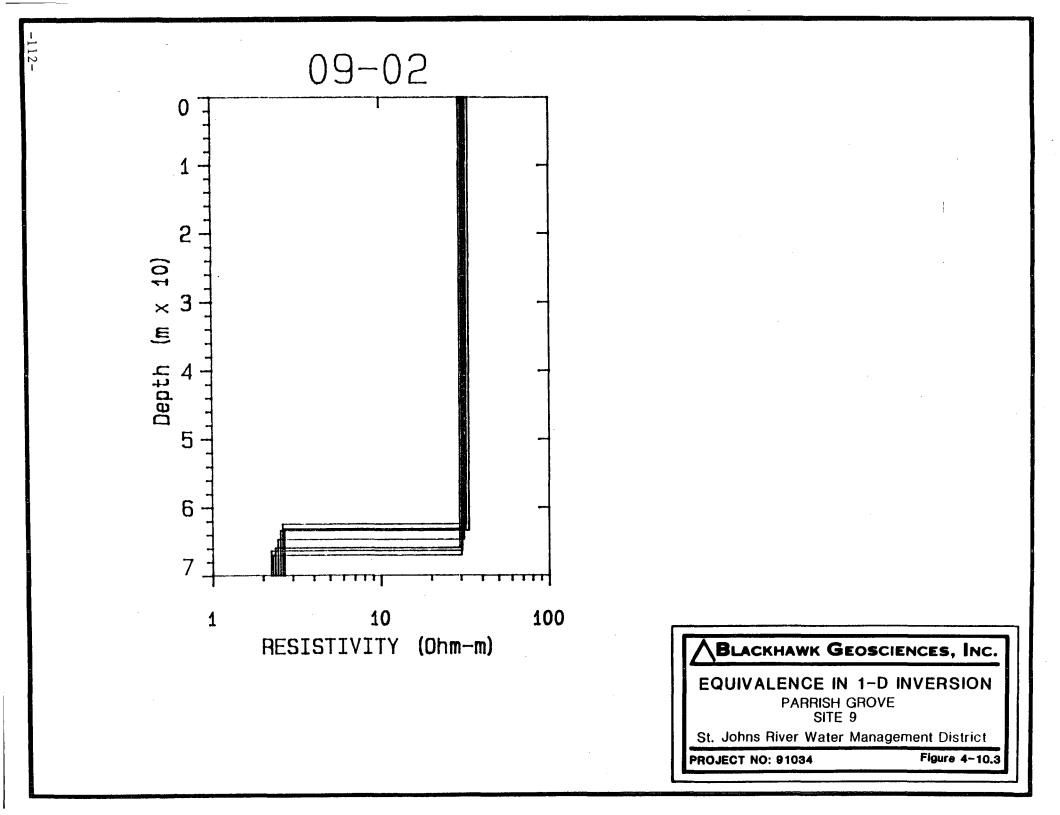
From the TDEM measurements the following information about water quality was derived:

- 1) The depth of occurrence of saline water was interpreted to occur at an elevation of 182 ft below msl, and below that depth chloride concentrations in excess of 10,000 mg/l chloride are expected.
- 2) Because no distinction could be made between the resistivity value of the Hawthorn Group, the surficial sediments, and the carbonate rocks of the Upper Floridan, no information about chloride concentration could be inferred above the highly saline water interface.





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MODEL: 2 LAYERS
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RE	SISTIVITY T	HICKNESS	ELEVATI	ON	CONDUCTANCE	(\$)
	(OHM-M)	(M)	(M)	(FEET)	LAYER	TOTAL
			9.1	30.0		
	31.19	64.7	-55.6 -	182.4	2.1	2.1
	2.44					
	TIMES	DATA	CALC	% ERROR	STD ERR	
1	8.90E-05	6.21E+01	6.24E+01	-0.450		
2	1.10E-04	5.70E+01	5.73E+01	-0.527		
3	1.40E-04	5.08E+01	5.09E+01	-0.211		
4	1.77E-04	4.42E+01	4.43E+01	-0.165		
5	2.20E-04	3.84E+01	3.81E+01	0.675		
6	2.80E-04	3.22E+01	3.18E+01	0.984		
7	3.55E-04	2.70E+01	2.65E+01	1.804		
8	4.43E-04	2.30E+01	2.23E+01	2.728		
9	5.64E-04	1.92E+01	1.86E+01	3.076		
10	7.13E-04	1.60E+01	1.57E+01	1.944		
11	8.81E-04	1.35E+01	1.36E+01	-0.797		
12	1.10E-03	1.16E+01	1.18E+01	-2.079		
13	1.41E-03	9.84E+00	1.01E+01	-2.838		
14	1.80E-03	8.47E+00	8.84E+00	-4.201		
15	2.22E-03	7.55E+00	7.91E+00	-4.524		
16	2.85E-03	6.58E+00	7.01E+00	-6.107		
17	3.60E-03	6.01E+00	6.32E+00	-4.899		
18	4.49E-03	6.02E+00	5.78E+00	4.047		
19	5.70E-03	5.63E+00	5.29E+00	6.296		
20	7.19E-03	5.21E+00	4.90E+00	6.427		

R: 30. X: 0. Y: 30. DL: 60. REQ: 33. CF: 1.0000 CLHZ ARRAY, 20 DATA POINTS, RAMP: 75.0 MICROSEC, DATA: 09-02 PARRISH GROVE

RMS LOG ERROR: 2.23E-02, ANTILOG YIELDS 5.2792 %

\* Blackhawk Geosciences, Incorporated \*

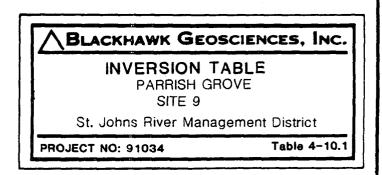


PARAMETER RESOLUTION MATRIX:

"F" MEANS FIXED PARAMETER P 1 1.00 P 2 0.00 1.00 T 1 0.00 0.00 1.00 P 1 P 2 T 1

## PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

LAYE		MINIMUM	BEST	MAXIMUM	
RHO	1	29.407	31,190	33.340	
	2	2.231	2.444	2.674	
THICK	1	62.447	64,725	67.029	
DEPTH	1	62.447	64.725	67.029	



## 4.11 NEW SMYRNA BEACH WESTERN WELL FIELD (SITE 10)

#### Location and Geoelectric Section

The detailed location of this sounding is shown in Figure 4-11.1. The interpreted geoelectric section is shown in Figure 4-11.2 and consists of three layers.

## Geologic Interpretation of Geoelectric Section

The upper layer in the geoelectric section is 263 ft thick and has a resistivity of 35 ohm-m. Since the regional information by Tibbals (1990) indicate the Hawthorn Group and younger sediments to be 100 ft thick, the first layer in the geoelectric section, therefore, also likely will include portions of the upper carbonate rocks of the Floridan aquifer. The resistivity value of this layer is as much determined by lithology as by water quality, and no inferences about water quality are drawn for this depth interval.

Geophysical logs were available for a well drilled in the same section. In Figure 4-11.3 the geoelectric section derived from TDEM soundings and the resistivity log run in the well are compared. Over the depth interval between the top of the casing (about 100 ft) and about 280 ft below surface, the average value of resistivity displayed by the log is about 50 ohm-m, and at about 280 ft below surface this increases to about 100 ohm-m. There is thus correspondence between resistivity logs and TDEM derived geoelectric sections manifested by:

- i) reasonable agreement in depth of occurrence of change in resistivity (263 ft versus 280 ft); and
- ii) both resistivity log and TDEM derived geoelectric section show an increase in resistivity at that depth.

There is some difference in absolute values of resistivity measured.

#### Depth of First Occurrence of Saline Water

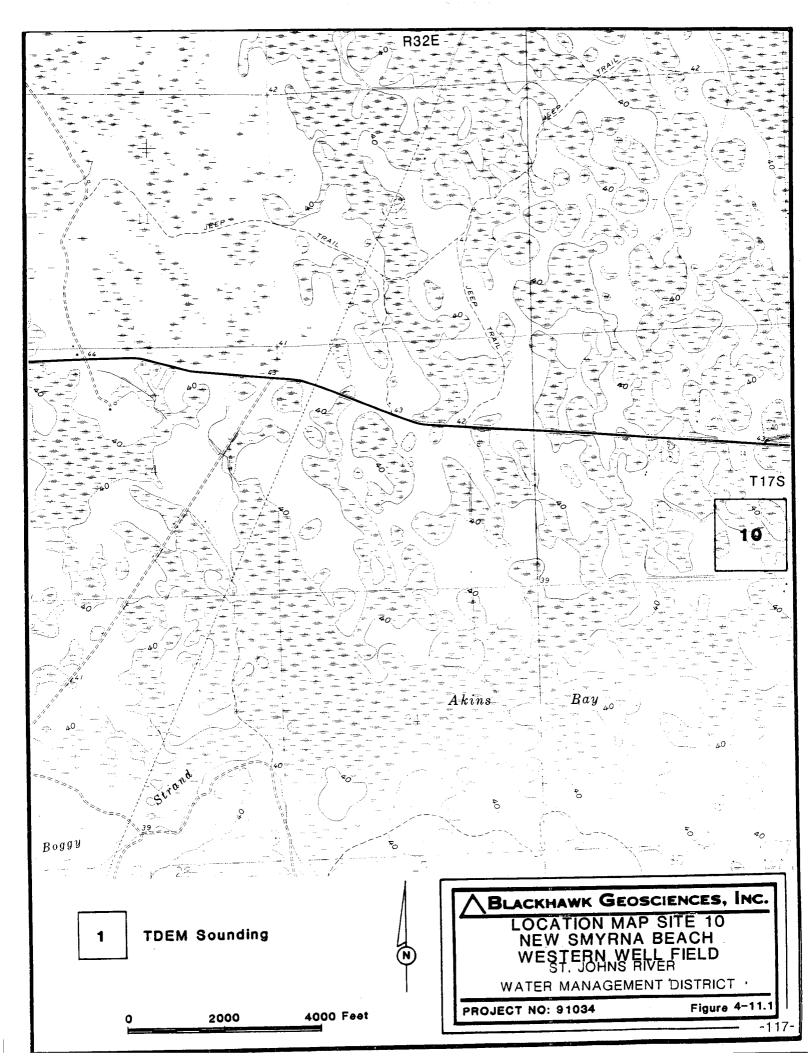
The lowest resistivity encountered in the geoelectric section is 9 ohm-m and it occurs at a depth of 949 ft below surface or at an elevation of 909 ft below msl. Assuming a porosity of 25% this would correspond to a chloride concentration in excess of 3,420 mg/l.

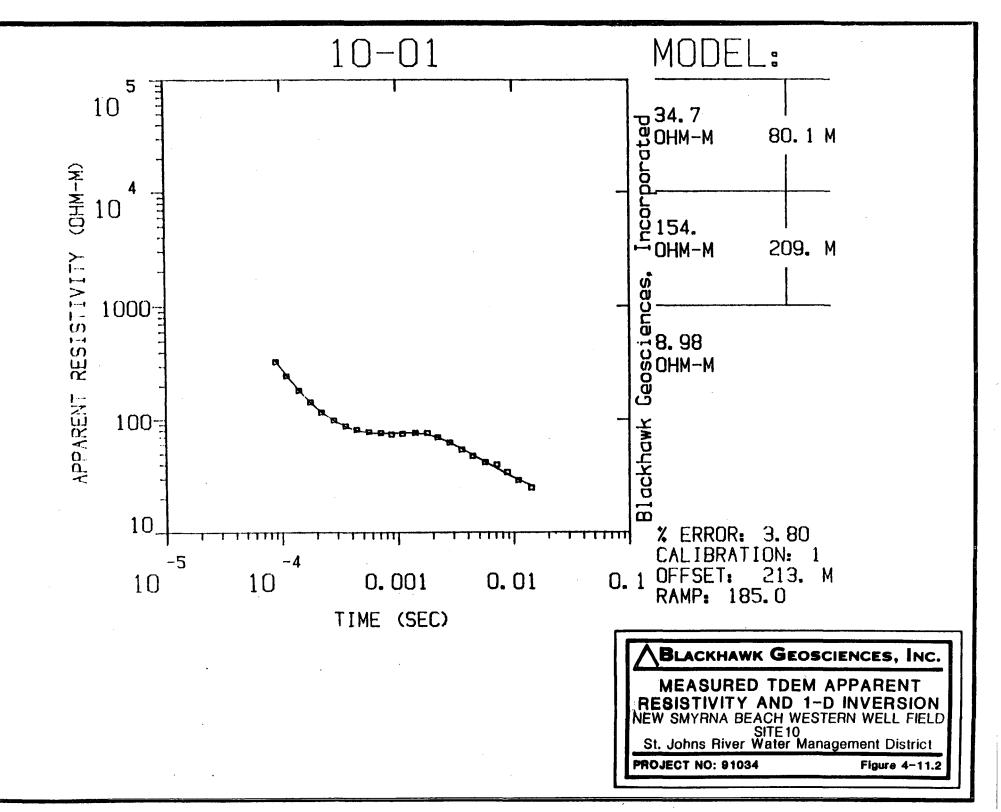
#### Depth of Occurrence of 250 mg/l Isochlor

The resistivity of the second layer in the geoelectric section in the elevation interval between 223 ft below msl and 909 ft below msl is 154 ohm-m. Assuming a porosity of 25% the ground water in the carbonate rocks over this interval is expected to have a chloride concentration less than 250 mg/l. The depth to the 250 mg/l isochlor is expected to coincide with the resistivity boundary between 154 ohm-m and 9 ohm-m at an elevation of 909 ft below msl. Likely a transition zone from ground water with chloride concentrations less than 250 mg/l to in excess of 3,000 mg/l will occur. This transition zone is not determined by the geoelectric section derived from TDEM.

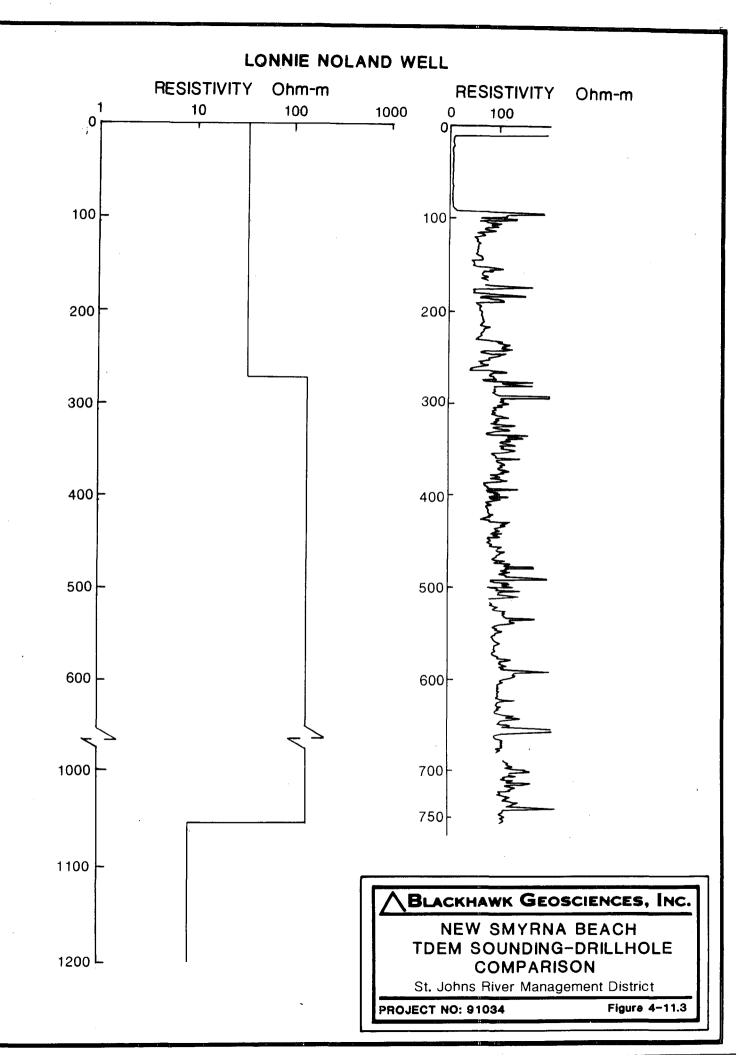
## Accuracy of Measurement and Interpretation

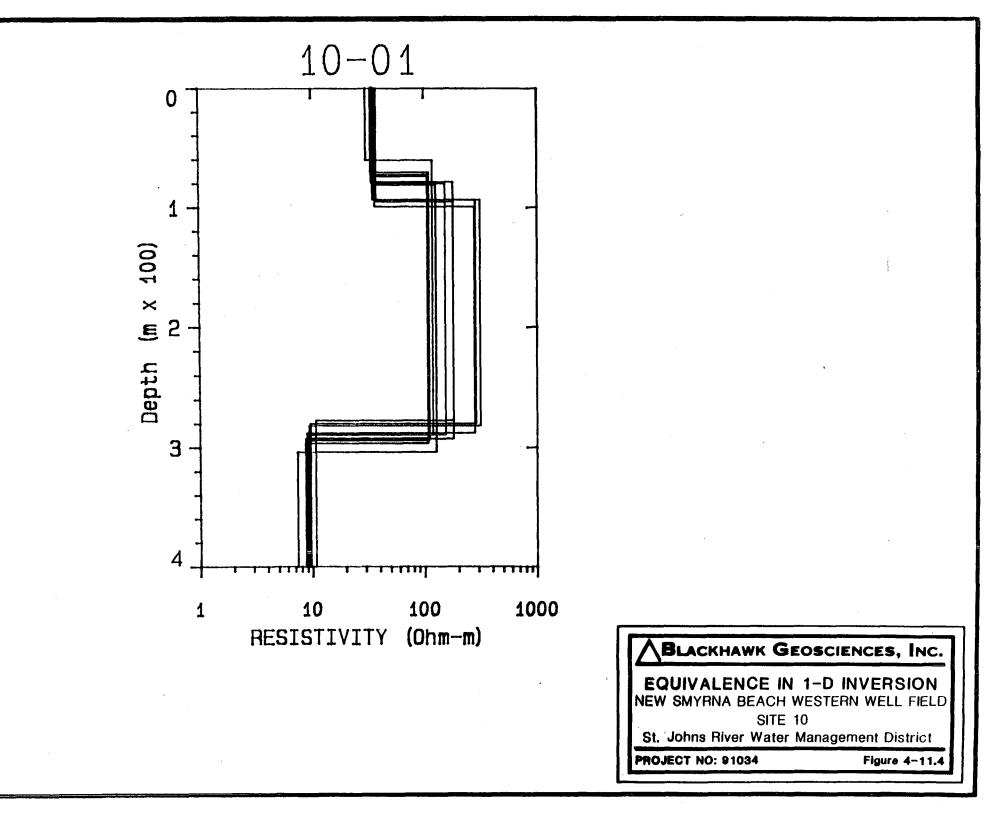
Figure 4-11.4 shows the evaluation of equivalence for the TDEM sounding at this site, and the inversion table (Table 4-11.1) lists the upper and lower bounds of the parameters of the geoelectric section. The range in the depth to the saline layer with a chloride concentration in excess of 3,420 mg/l is about 10% of total depth, or about  $\pm$  90 ft. The range of equivalence in the resistivity of the second layer, the main section of the Floridan aquifer, is





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MODEL: 3 LAYERS

RE	SISTIVITY	THICKNESS	ELEVAT	ION	CONDUCTANCE	(\$)
	(OHM-M)	(M)	(M)	(FEET)	LAYER	TOTAL
			12.2	40.0		
	34.69	80.1	-67.9	-222.7	2.3	2.3
· 1	53.69	209.1	-277.0	-908,8	1.4	3.7
	8.98					
	TIMES	DATA	CALC	% ERROR	STD ERR	
1	8.90E-05	3.31E+02	3.36E+02	-1.536		
2	1.10E-04	2.46E+02	2.51E+02	-1.794		
3	1.40E-04	1.83E+02	1.85E+02	-1.244		
4	1.77E-04	1.44E+02	1.43E+02	0.330		
5	2.20E-04	1.16E+02	1.17E+02	-1.026		
6	2.80E-04	9.82E+01	9.83E+01	-0.045		
7	3.55E-04	8.73E+01	8.66E+01	0.855		
8	4.43E-04	8.10E+01	8.00E+01	1.258		
9	5.64E-04	7.74E+01	7.62E+01	1.594		
10	7.13E-04	7.60E+01	7.51E+01	1.125		
11	8.81E-04	7.37E+01	7.55E+01	-2.380		
12	1.10E-03	7.46E+01	7.63E+01	-2.106		
13	1.41E-03	7.61E+01	7.61E+01	-0.092		
14	1.80E-03	7.58E+01	7.36E+01	3.066		
15	2.21E-03	6.92E+01	6.91E+01	0.045		
16	2.80E-03	6.20E+01	6.24E+01	-0.547		
17	3.55E-03	5.38E+01	5.50E+01	-2.147		
18	4.43E-03	4.74E+01	4.84E+01	-1.946		
19	5.64E-03	4.16E+01	4.20E+01	-0.883		
20	7.13E-03	3.97E+01	3.68E+01	7.791		
21	8.81E-03	3.42E+01	3.28E+01	4.265		
22	1.10E-02	2.89E+01	2.93E+01	-1.488		
23	1.41E-02	2.48E+01	2.60E+01	-4.657		

R: 213. X: 0. Y: 213. DL: 426. REQ: 237. CF: 1.0000 TOHZ ARRAY, 23 DATA POINTS, RAMP: 185.0 MICROSEC, DATA: 10-01 NEW SMYRNA BEACH WESTERN WELL FIELD

RMS LOG ERROR: 1.62E-02, ANTILOG YIELDS

3.8013 % ABLACKHAWK GEOSCIENCES, INC. INVERSION TABLE NEW SMYRNA BEACH WESTERN WELL FIELD SITE 10 St. Johns River Management District PROJECT NO: 91034 Table 4-11.1 LATE TIME PARAMETERS

\* Blackhawk Geosciences, Incorporated \*

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PARAMETER RESOLUTION MATRIX:

"F" MEANS FIXED PARAMETER

P 1 0.99

P 2 -0.02 0.15

P 3 0.00 -0.04 0.96

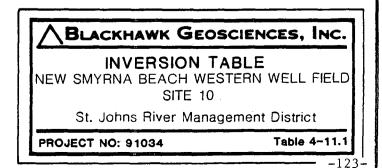
T 1 -0.02 -0.19 0.00 0.93

T 2 0.01 0.13 0.01 0.03 0.98

P 1 P 2 P 3 T 1 T 2
```

## PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

LAYER		MINIMUM	BEST	MAXIMUM	
RHO	1	30.544	34.689	37.583	
	2	107.440	153.693	313.308	
	3	7.316	8.981	10.696	
THICK	1	60.760	80.084	99.551	
	2	186,176	209.110	228.708	
DEPTH	1	60.760	80.084	99.551	
	2	277.223	289.194	304.026	



between 107 ohm-m and 313 ohm-m. For this range of resistivity chloride concentration less than 250 mg/l are expected.

## Summary of Results of TDEM Measurements at the New Smyrna Beach Western Well Field (Site 10)

From the TDEM measurements at New Smyrna Beach Western Well Field, the following information about aquifer characteristics and water quality was derived:

- 1) The depth of occurrence of saline water was interpreted at a depth of 909 ft below msl. The chloride concentration at that depth is expected to exceed 3,420 mg/l.
- 2) On the basis of the resistivity value of about 154 ohm-m measured for the carbonate rock of the Upper Floridan aquifer over the elevation interval between 223 ft to 909 ft below msl, the chloride concentration is expected to be less than 250 mg/l. The depth of occurrence of the 250 mg/l isochlor is expected to approximately coincide with the resistivity boundary measured at an elevation of about 909 ft below msl.

## 4.12 ORANGE COUNTY LANDFILL (SITE 11)

The detailed location map of the sounding is shown in Figure 4-12.1. The measured TDEM data and the inverted geoelectric section is shown in Figure 4-12.2 and it consists of a three layer section.

#### Geologic Interpretation of Geologic Section

A well located approximately 3 miles northeast of the site was completed to a depth of 1,400 ft. The lithologic log of that well (Fig. 4-12.3) shows surficial sands and the Hawthorn Group to extend to approximately 250 ft below surface. The upper layer in the geoelectric section is 288 ft and has a resistivity of 33 ohm-m. If the information in the drilled well 3 miles away can be extrapolated to the location of the TDEM measurement, the upper layer can be expected to represent the surficial sands and the Hawthorn Group. The resistivities of the two layers below 288 ft represent different strata within the carbonate rocks of the Floridan aquifer.

## Depth of Occurrence of Saline Water

The lowest resistivity encountered within the effective depth of measurement is 10.6 ohm-m, and it is first encountered at an elevation of about 2,304 ft below msl. The resistivity below that depth is 10.6 ohm-m, and at an average porosity of 25%, that value corresponds to a chloride content in excess of 2,880 mg/l.

## Depth of Occurrence of 250 mg/l Isochlor

The resistivity of the interval between elevations of 203 ft below msl and 2,304 ft below msl is 708 ohm-m. At an average porosity of 25%, the chloride content of this interval is expected to be less than 100 mg/l. The 250 mg/l isochlor is expected to approximately correspond with the elevation of the boundary between 708 ohm-m and 10.6 ohm-m at 2,304 ft below msl. A transition zone from ground water with a low chloride content to a chloride content in excess of 2,880 mg/l can be expected. Such transition zones are difficult to map with surface geophysics.

## Accuracy of Measurement and Interpretation

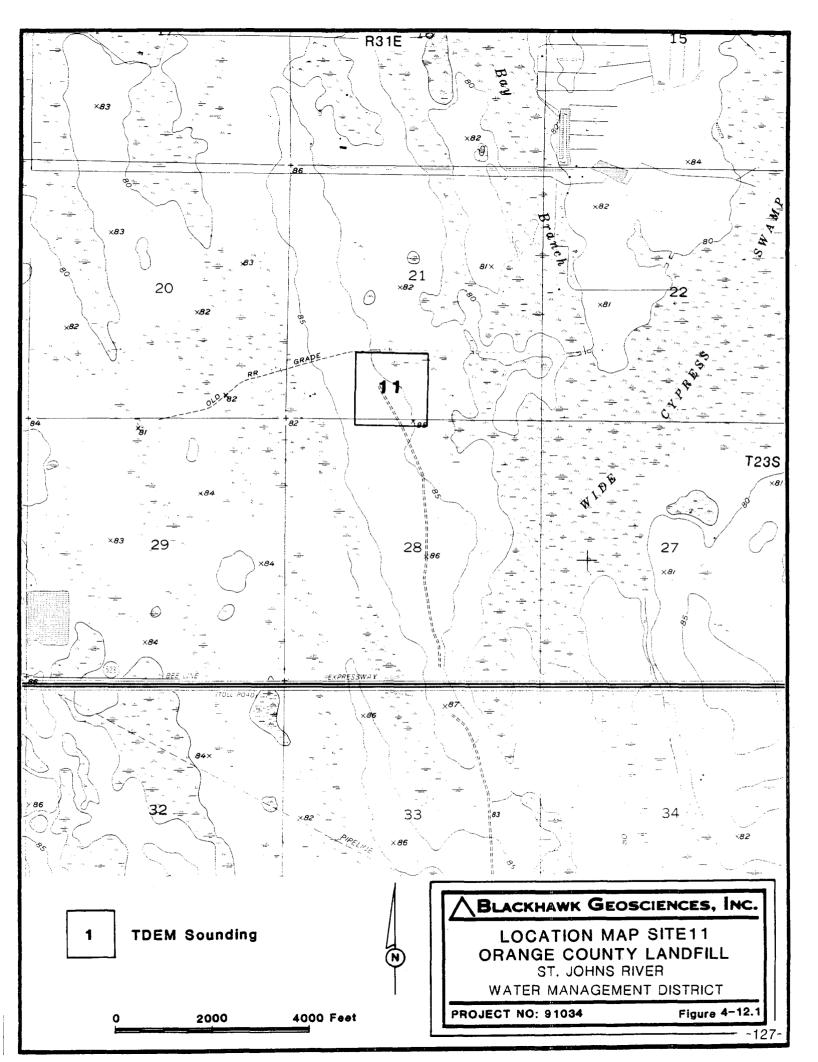
Figure 4-12.4 shows the evaluation of equivalence of the TDEM sounding at this site and the inversion table (Table 4-12.1) lists the upper and lower bounds of the parameters of the geoelectric section. The depth to the saline water is determined with an accuracy of about  $\pm$  5% of total depth.

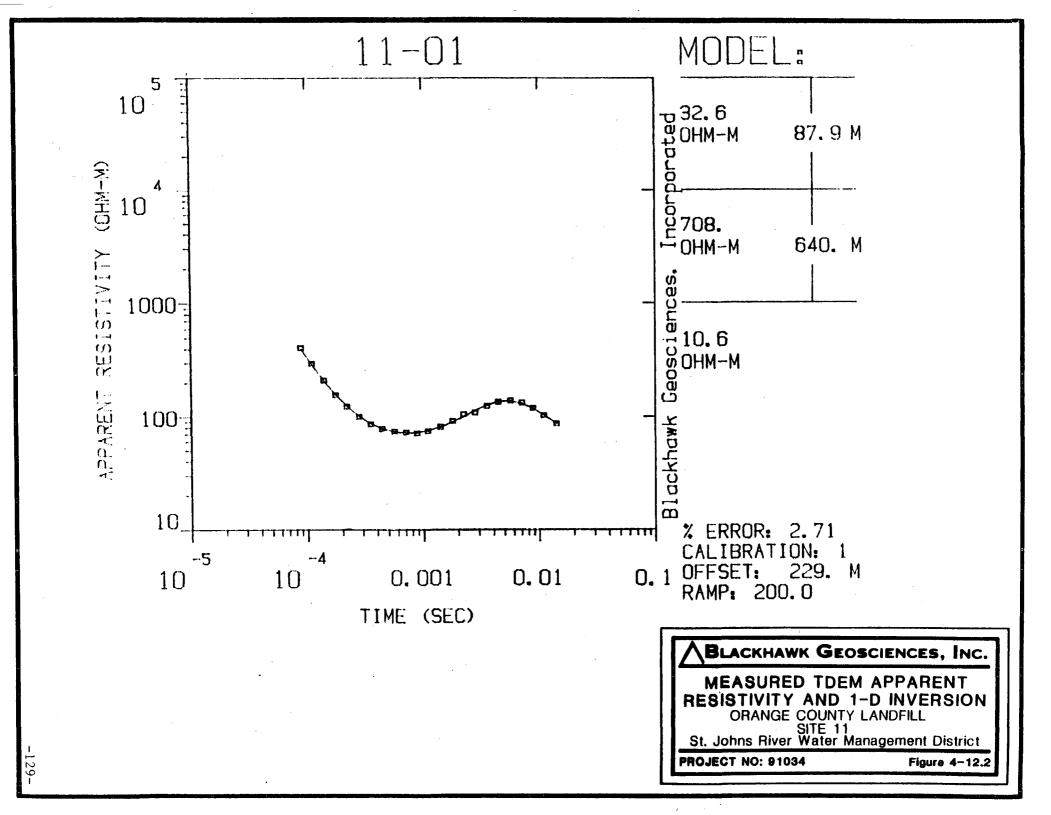
The equivalence in the resistivity of the layer at elevations between 203 ft below msl and 2,304 ft below msl is between 441 ohm-m and 1,303 ohm-m. Although this range of equivalence is large, it does not negate the conclusion that the chloride content over this depth interval is well below 250 mg/l.

## <u>Summary of Results of TDEM Measurements at the Orange County Landfill</u> (Site 11)

From the TDEM soundings the following information about water quality was derived:

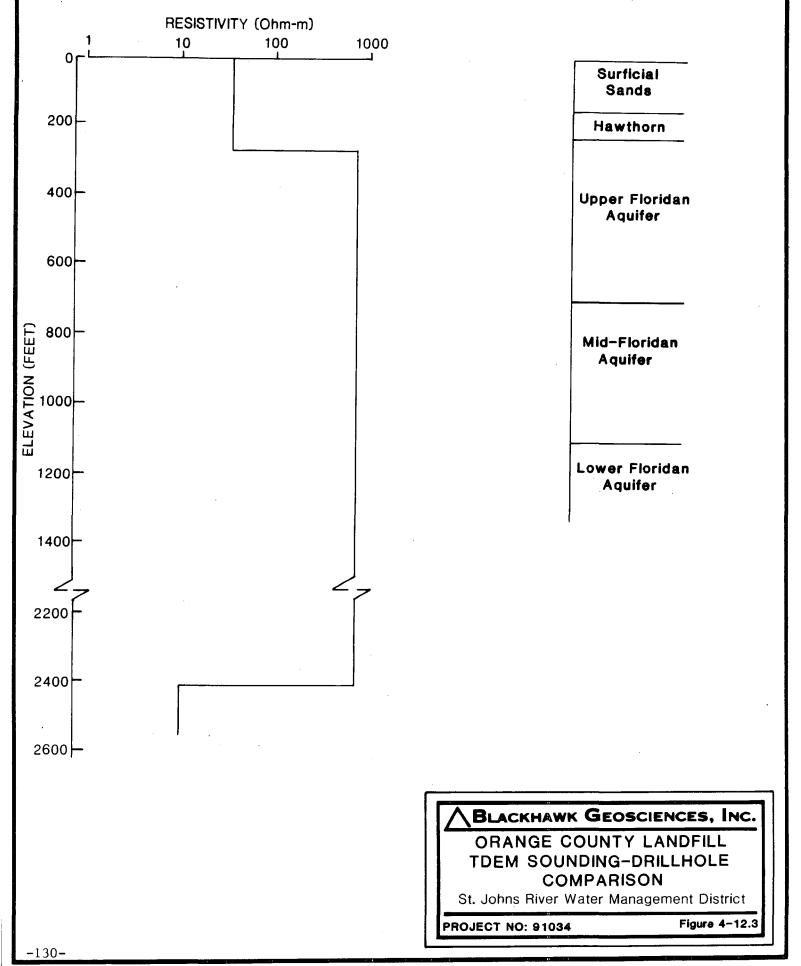
- The depth of occurrence of saline water was interpreted to occur at an elevation of 2,304 ft below msl, and below that elevation chloride concentrations in excess of 2,880 mg/l chloride are expected.
- 2) Below the Hawthorn Group, at an elevation of 203 ft below msl and above the interface with saline water at 2,304 ft below msl, the chloride content of ground water is expected to be less than 250 mg/l.

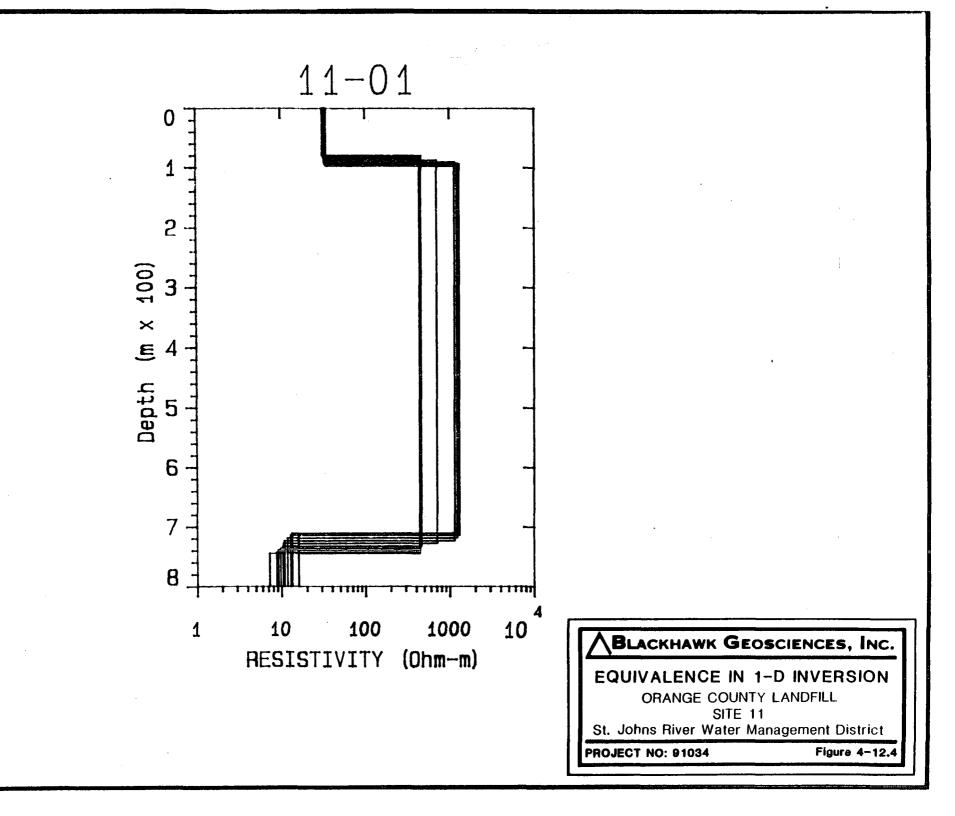




## ORANGE COUNTY LANDFILL SITE 11 TDEM SOUNDING

## LOWER FLORIDAN EXPLORATORY WELL





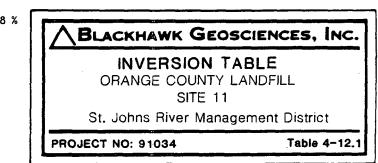
131-

MODEL: 3 LAYERS

RE	SISTIVITY	THICKNESS	ELEVAT	ION	CONDUCTANCE	(S)
	(OHM-M)	(M)	(M)	(FEET)	LAYER	TOTAL
			25.9	85.0		
	32.64	87.9	-62.0	-203.3	2.7	2.7
7	08.29	640.4 -	702.3 -	2304.2	0.9	3.6
	10.56					
	TIMES	DATA	CALC	% ERROR	STD ERR	
1	8.90E-05	4.11E+02	4.00E+02	2.678		
2	1.10E-04	2.97E+02	2.92E+02	1.820		
3	1.40E-04	2.11E+02	2.09E+02	0.954		
4	1.77E-04	1.56E+02	1.57E+02	-0.540		
5	2.20E-04	1.23E+02	1.25E+02	-1.016		
6	2.80E-04	9-98E+01	1.01E+02	-1.382		
7	3.55E-04	8.56E+01	8.65E+01	-1.055		
8	4.43E-04	7.77E+01	7.80E+01	-0.431		
9	5.64E-04	7.34E+01	7.29E+01	0.740		
10	7.13E-04	7.20E+01	7.11E+01	1.199		
11	8.81E-04	7.07E+01	7.18E+01	-1.490		
12	1.10E-03	7.40E+01	7.46E+01	-0.841		
13	1.41E-03	8.09E+01	8.06E+01	0.357		
14	1.80E-03	9.11E+01	8.92E+01	2.101		
15	2.22E-03	1.05E+02	9.93E+01	5.391		
16	2.80E-03	1.08E+02	1.12E+02	-3.620		
17	3.55E-03	1.24E+02	1.26E+02	-1.736		
18	4.43E-03	1.35E+02	1.36E+02	-0.703		
19	5.64E-03	1.39E+02	1.39E+02	0.007		
20	7.13E-03	1.33E+02	1.30E+02	1.584		
21	8.81E-03	1.18E+02	1.18E+02	0.671		
22	1.10E-02	1.02E+02	1.03E+02	-0.488		
23	1.41E-02	8.59E+01	8.64E+01	-0.592		
R:	229. X:	0.Y: 229	. DL: 457.	REQ: 254.	CF: 1.0000	

R: 229. X: 0. Y: 229. DL: 457. REQ: 254. CF: 1.0000 CLHZ ARRAY, 23 DATA POINTS, RAMP: 200.0 MICROSEC, DATA: 11-01 ORANGE COUNTY LAND FILL

RMS LOG ERROR: 1.16E-02, ANTILOG YIELDS 2.7148 %



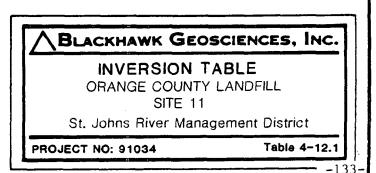
#### LATE TIME PARAMETERS

\* Blackhawk Geosciences, Incorporated \*

PARAMETER RESOLUTION MATRIX: "F" MEANS FIXED PARAMETER P 1 1.00 P 2 0.00 0.86 P 3 0.00 -0.04 0.96 T 1 0.00 -0.02 0.00 1.00 T 2 0.00 0.01 0.00 0.00 1.00 P 1 P 2 P 3 T 1 T 2

PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

LAYER		MINIMUM	BEST	MAXIMUM
RHO	1	31.098	32.637	34.463
	2	441.161	708.289	1303.037
	3	7.175	10.558	15.788
THICK	1	79.843	87.872	97.893
	2	618.089	640.354	659.821
DEPTH	1	79.843	87.872	97.893
	2	711.182	728.226	745.170



## 4.13 DESERET RANCH (SITE 12)

## Location and Geoelectric Section

The detailed location map of this sounding is shown in Figure 4-13.1. The TDEM data and the inverted geoelectric section is shown in Figure 4-13.2 and consists of a three layer section.

## Depth of Occurrence of Saline Water

The lowest resistivity encountered within the effective depth of exploration of the measurement is 9.4 ohm-m, and it is first encountered at an elevation of 1,292 ft below msl. Assuming an average porosity of 25% the chloride concentration below that depth is expected to exceed 3,270 mg/l.

#### Depth of Occurrence of 250 mg/l Isochlor

The upper 55 ft thick layer with a resistivity of 188 ohm-m is assumed to represent surficial sands and the Hawthorn Group, and the second layer with a resistivity of 26 ohm-m between elevations of 8 ft above msl and 1,292 ft below msl is characteristic of the carbonate rocks of the Upper Floridan aquifer, but also may contain portions of the Hawthorn Group, since it may be more than 55 ft thick here (Toth, 1988). However, since the thickness of the 26 ohm-m is 1,300 ft, the 26 ohm-m value is probably not much influenced by thin clay layers in the Hawthorn Group. Assuming a porosity of 25% a resistivity of 26 ohm-m corresponds to a chloride concentration in excess of 250 mg/l for that entire interval.

#### Accuracy of Measurement and Interpretation

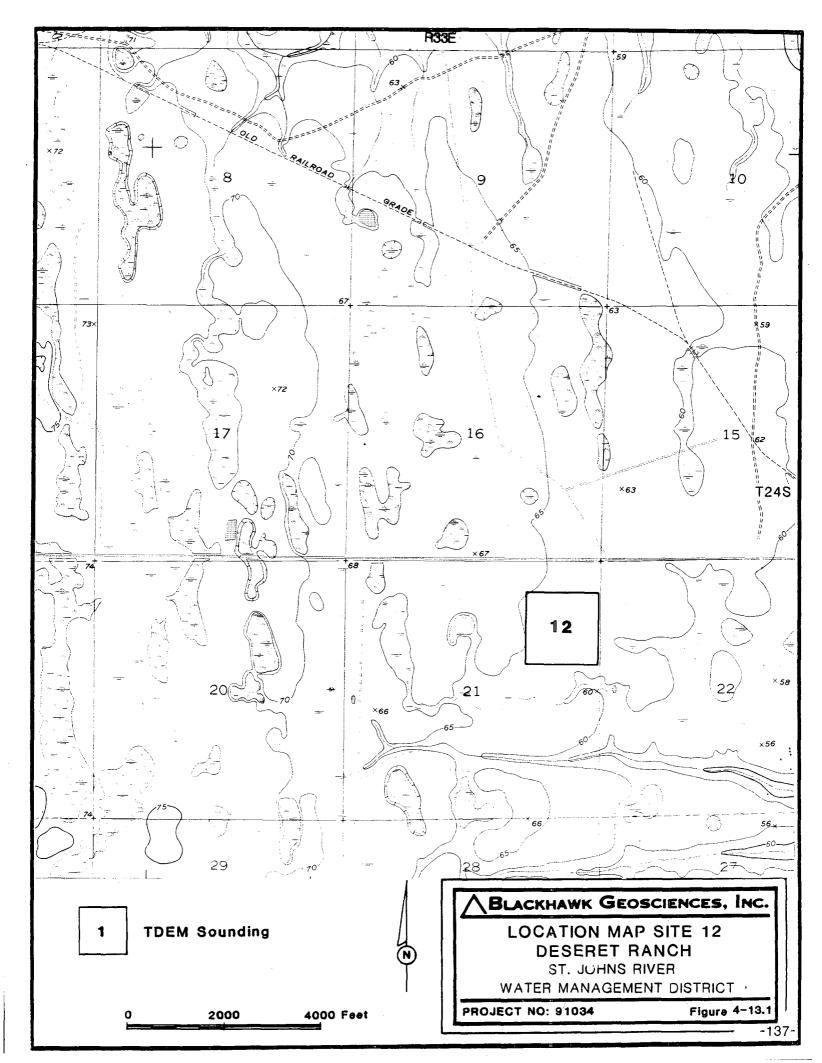
Figure 4-13.3 shows the evaluation of equivalence of the TDEM sounding at this site, and the inversion table (Table 4-13.1) lists the upper and lower bounds of the parameters of the geoelectric section.

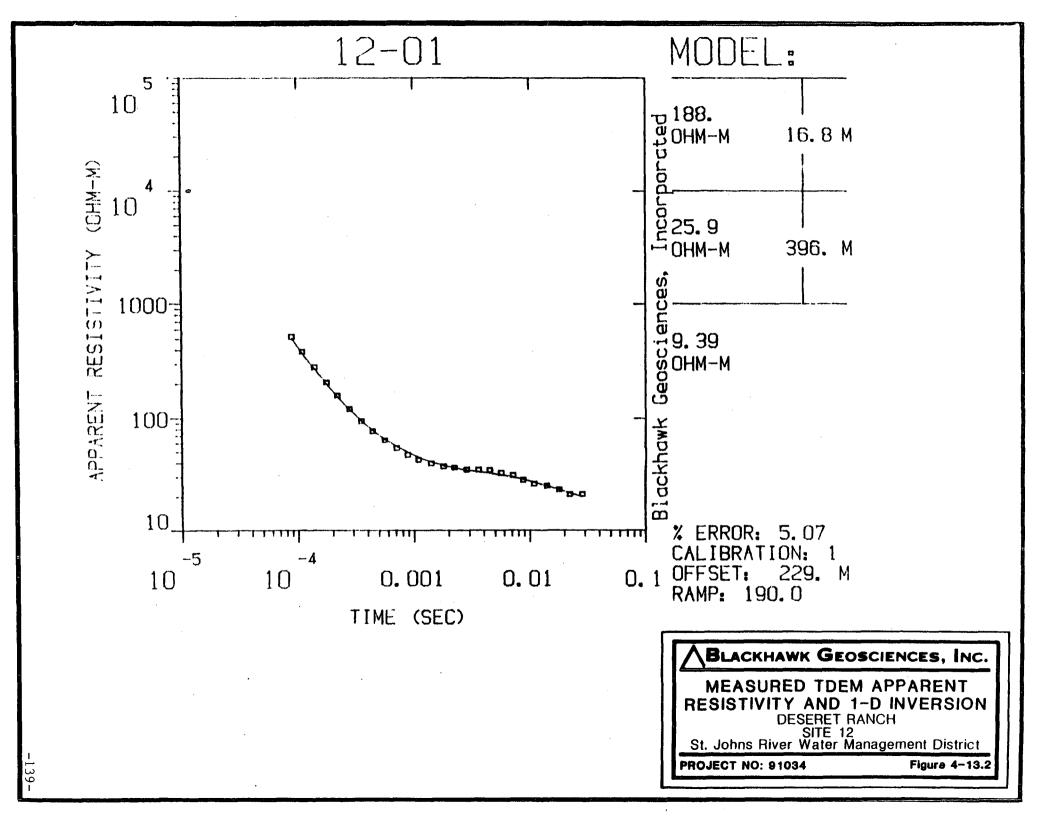
The range of equivalence in determining the depth to saline layer (9.6 ohm-m) is relatively large,  $\pm$  15% of total depth or about  $\pm$  200 ft for this geoelectric section. The range of equivalence for the resistivity of the second layer, the main section of the Floridan aquifer, is between 25 ohm-m and 27 ohm-m, and for this range of resistivity values chloride concentration in excess of 250 mg/l are predicted for a porosity of 25%.

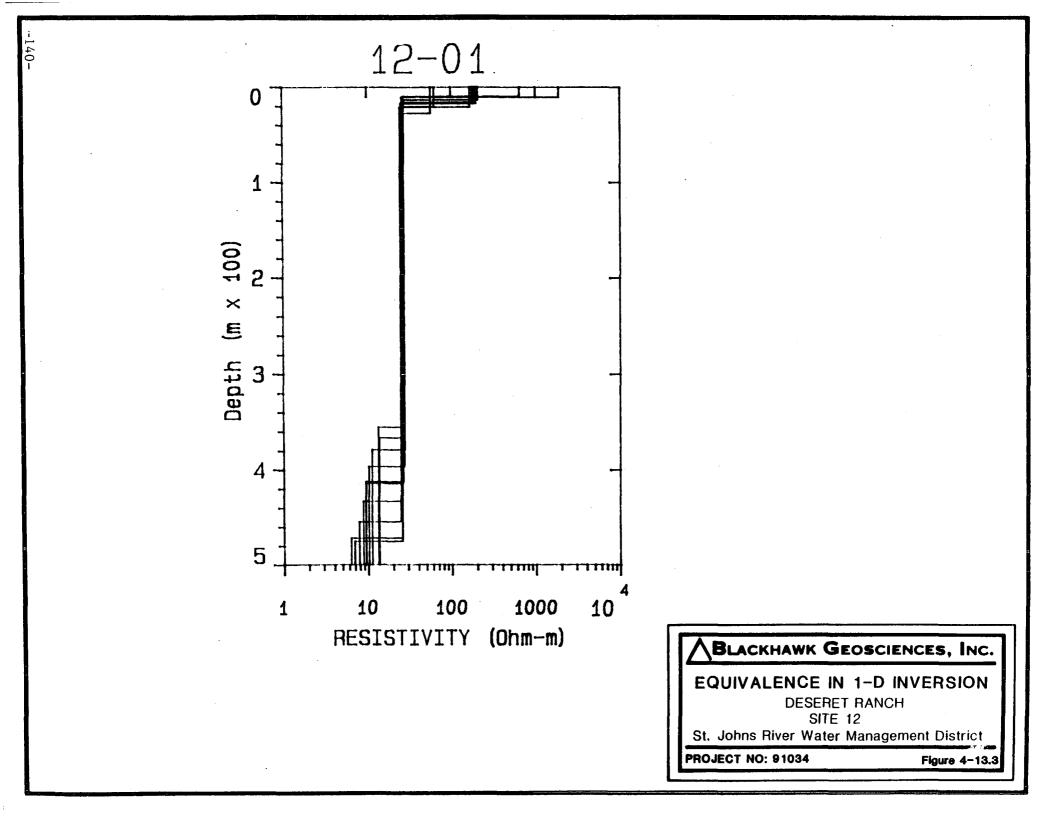
## Summary of Results of TDEM Measurements at Deseret Ranch (Site 12)

From the TDEM measurements at the Deseret Ranch, the following information about water quality was derived:

- The depth of occurrence of saline water was interpreted to occur at an elevation of 1,292 ft below msl, and below that elevation chloride concentrations in excess of 3,270 mg/l are expected. The range of equivalence in determining depth to that layer was relatively large (<u>+</u> 200 ft).
- 2) In the Floridan aquifer above the interface with saline water at 1,292 ft below msl and below the Hawthorn Group, the chloride concentration of ground water is predicted to be in excess of 250 mg/l.







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MODEL:

**3** LAYERS

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TOTAL

R	ESISTIVITY	THICKNESS	ELEVA	TION	CONDUCTANCE	(\$)
	(OHM-M)	(M)	(M)	(FEET)	LAYER	TOTA
			19.2	63.0		
	187.68	16.8	2.4	8.0	0.1	0.1
	25.95	396.3	-393.9	1292.2	15.3	15.4
	9.39					
	TIMES	DATA	CALC	% ERROR	STD ERR	
1	8.90E-05	5.24E+02	2 5.07E+02	2 3.352		
2						
3						
4			2.00E+02			
5	2.20E-04	1.58E+02	2 1.54E+02	2 2.483		
6	2.80E-04	1.20E+02	1.19E+02	2 1.073		
7	3.55E-04	9.41E+01	9.45E+01	-0.470		
8	4.43E-04	7.69E+01	7.82E+01	-1.627		
9	5.64E-04	6.38E+01	6.53E+01	-2.301		
10	7.13E-04	5.41E+01	5.61E+01	-3.534		
11	8.81E-04	4.73E+01	4.99E+01	-5.333		
12	1.10E-03	4.28E+01	4.51E+01	-5.091		
13	1.41E-03	3.95E+01	4.09E+01	-3.272		
14	1.80E-03	3.72E+01	3.79E+01	-1.826		
15	2.22E-03	3.63E+01	3.61E+01	0.572		
16	2.85E-03	3.48E+01	3.45E+01	l 1.031		
17	3.60E-03	3.49E+01	3.34E+01	4,500		
18	4.49E-03	3.45E+01	3.24E+01	6,465		
19	5.67E-03	3.26E+01	3.13E+01	3.887		
20	7.16E-03	3.11E+01	3.00E+01	3.686		
21	8.81E-03	2.83E+01	2.86E+01	I -1.146		
22	1.10E-02	2.62E+01	2.70E+0	-3,107		
23	1.41E-02	2.50E+01	2.51E+01	-0.007		
24	1.80E-02	2,32E+01	2.32E+01	0.025		
25	2.22E-02	2.11E+01	2.17E+0	-2.455		
26	2.85E-02	2.12E+01	2.01E+0	5.583		

R: 229. X: 0. Y: 229. DL: 457. REQ: 254. CF: 1.0000 CLHZ ARRAY, 26 DATA POINTS, RAMP: 190.0 MICROSEC, DATA: 12-01W



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DESERT RANCH

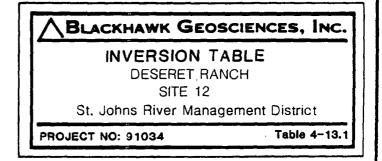
RMS LOG ERROR: 2.15E-02, ANTILOG YIELDS 5.0655 % LATE TIME PARAMETERS

\* Blackhawk Geosciences, Incorporated \*

PARAMETER RESOLUTION MATRIX: "F" MEANS FIXED PARAMETER P 1 0.05 P 2 0.00 1.00 P 3 0.00 0.00 0.89 T 1 0.19 0.01 0.01 0.92 T 2 -0.01 0.00 0.04 0.00 0.98 P 1 P 2 P 3 T 1 T 2

# PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

LAYER		MINIMUM	BEST	MAXIMUM	
RHO	1	57.660	187.682	1876.823	
	2	24.535	25.946	27.277	
	3	6.241	9.388	13.529	
THICK	1	9.882	16.773	27.735	
	2	339.539	396.305	458.229	
DEPTH	1	9.882	16.77 <b>3</b>	27.735	
	2	355.784	413.078	475.514	



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## 4.14 LAKE MARY (SITE 13)

## Location and Geoelectric Section

The detailed location map of this sounding is shown in Figure 4-14.1. The TDEM data and the inverted geoelectric section is shown in Figure 4-14.2, and consists of a three layer section.

#### Depth of Occurrence of Saline Water

The lowest resistivity encountered within the effective depth of exploration of the measurement is 5.5 ohm-m, and it is first encountered at an elevation of 1,203 ft below msl. Assuming an average porosity of 25% the chloride concentration below that depth is expected to exceed 5,700 mg/l.

## Depth of Occurrence of 250 mg/l Isochlor

The upper 123 ft thick layer with a resistivity of 60 ohm-m is assumed to represent surficial sands and the Hawthorn Group, and the second layer with a resistivity of 293 ohm-m between elevations of 78 ft below msl and 1,203 ft below msl is characteristic of the carbonate rocks of the Upper Floridan aquifer. Assuming a porosity of 25% a resistivity of 293 ohm-m corresponds to a chloride concentration less than 250 mg/l for that entire interval. Chloride concentrations less than 250 mg/l for this area are also reported by Tibbals (1990). The 250 mg/l isochlor is expected to coincide with the resistivity boundary between 293 ohm-m and 5.5 ohm-m observed at an elevation of 1,203 ft below msl. A transition zone from ground water with chloride concentrations less than 250 mg/l to in excess of 5,700 mg/l can be expected, so that the isochlor may be some 100 ft above the mapped interface.

#### Accuracy of Measurement and Interpretation

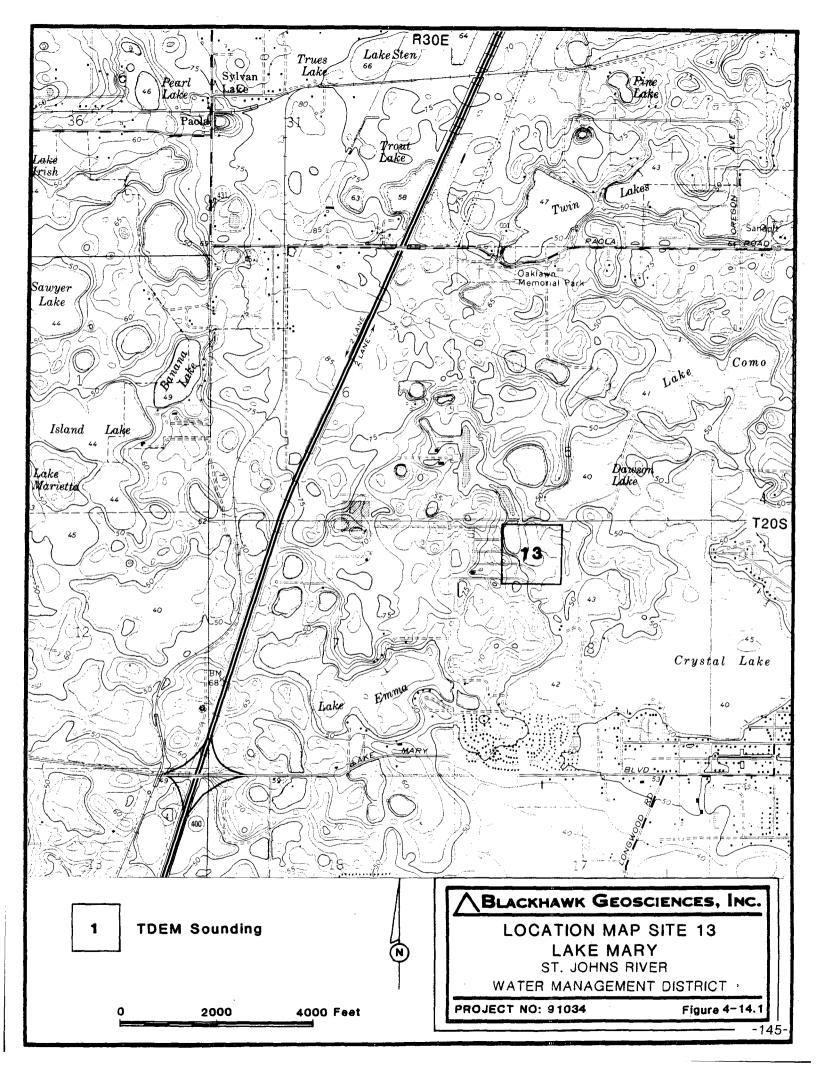
Figure 4-14.3 shows the evaluation of equivalence of the TDEM sounding at this site, and the inversion table (Table 4-13.1) lists the upper and lower bounds of the parameters of the geoelectric section.

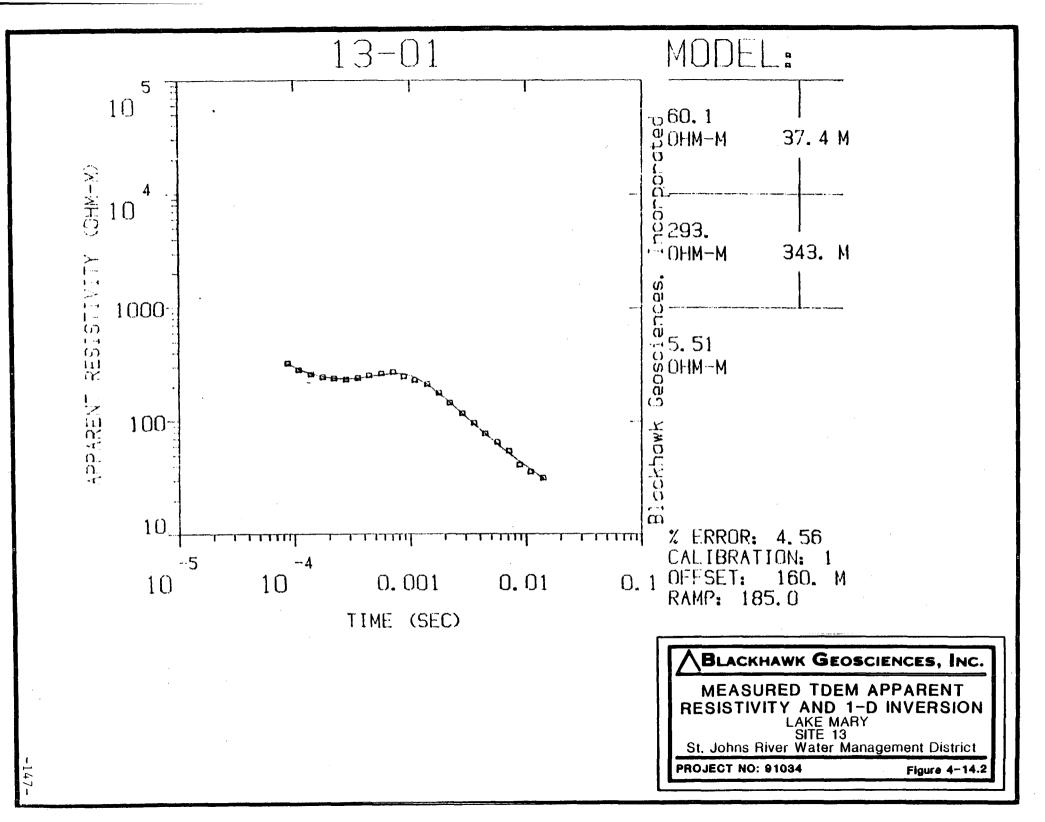
The range of equivalence in determining the depth to the saline layer (5.5 ohm-m) is less than  $\pm$  5% of total depth, or about  $\pm$  60 ft. The equivalence in the resistivity of the second layer, representing the main section of the Upper Floridan aquifer, is between 244 ohm-m and 387 ohm-m. Assuming a porosity of 25% chloride concentrations less than 250 mg/l are expected for that range of resistivity.

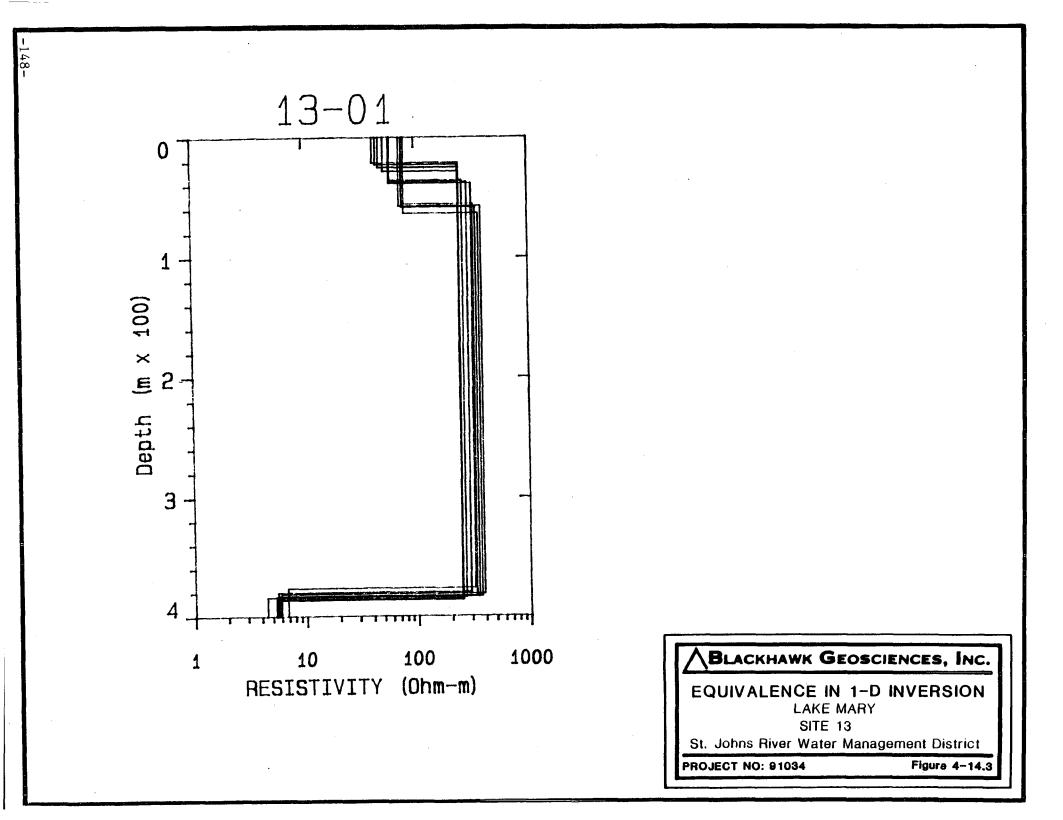
#### Summary of Results of TDEM Measurements at Lake Mary (Site 13)

From the TDEM measurements at Lake Mary, the following information about water quality was derived:

- The depth of occurrence of saline water was interpreted to occur at an elevation of 1,203 ft below msl, and below that elevation chloride concentrations in excess of 5,700 mg/l are expected.
- 2) Above the interface with saline water at 1,203 ft below msl and below the Hawthorn Group (78 ft below msl), the chloride concentration of ground water is predicted to be less than 250 mg/l for that entire interval. The depth of occurrence of the 250 mg/l isochlor is expected to approximately coincide with the interface mapped at an elevation of 1,203 ft below msl.







MODEL: 3 LAYERS

R	ESISAIVITY	THICKNESS - jp	ELEVAS	ON	CONDUCTANCE	(S)
	(OHM-M)	(M)	(H)	(FEET)	LAYER	TOTAL
			13.7	45.0		
	60.06	37.4	-23.7	-77.7	0.6	0.6
	292.61	342.9 -	·366.5 - '	1202.6	1.2	1.8
	5.51					
	TIMES	DATA	CALC	% ERROR	STD ERR	
1	8.90E-05	3.23E+02	3.19E+02	1.084		
2	1.10E-04	2.84E+02	2.87E+02	-1.014		
3	1.40E-04	2.58E+02	2.61E+02	-0.955		
4	1.77E-04	2.43E+02	2.45E+02	-0.930		
5	2.20E-04	2.37E+02	2.37E+02	-0.017		
6	2.80E-04	2.33E+02	2.35E+02	-0.709		
7	3.55E-04	2.40E+02	2.39E+02	0.537		
8	4.43E-04	2.53E+02	2.47E+02	2.460		
9	5.64E-04	2.64E+02	2.58E+02	2.670		
10	7.13E-04	2.71E+02	2.64E+02	2.610		
11	8.81E-04	2.48E+02	2.61E+02	-5.042		
12	1.10E-03	2.30E+02	2.44E+02	-5.810		
13	1.41E-03	2.13E+02	2.11E+02	0.653		
14	1.78E-03	1.77E+02	1.75E+02	0.978		
15	2.21E-03	1.44E+02	1.45E+02	-0.628		
16	2.83E-03	1.17E+02	1.16E+02	0.834		
17	3.57E-03	9.64E+01	9.41E+01	2.514		
18	4.46E-03	7.79E+01	7.74E+01	0.660		
19	5.67E-03	6.53E+01	6.30E+01	3.708		
20	7.16E-03	5.48E+01	5.20E+01	5.259		
21	8.81E-03	4.14E+01	4.42E+01	-6.422		
22	1.10E-02	3.58E+01	3.76E+01	-4.551		
23	1.41E-02	3.19E+01	3.14E+01	1.397		

R: 160. X: 0. Y: 160. DL: 320. REQ: 178. CF: 1.0000 CLHZ ARRAY, 23 DATA POINTS, RAMP: 185.0 MICROSEC, DATA: 13-01 LAKE MARY

RMS LOG ERROR: 1.94E-02, ANTILOG YIELDS 4.5585

ABLACKHAWK GEO	SCIENCES, IN
INVERSION T LAKE MAR SITE 13	
St. Johns River Mana	gement District
PROJECT NO: 91034	Table 4-14

LATE TIME PARAMETERS

\* Blackhawk Geosciences, Incorporated \*

```
PARAMETER RESOLUTION MATRIX:

"F" MEANS FIXED PARAMETER

P 1 0.78

P 2 -0.11 0.85

P 3 0.00 -0.04 0.92

T 1 -0.36 -0.24 -0.01 0.37

T 2 0.04 0.03 0.01 0.07 0.99

P 1 P 2 P 3 T 1 T 2
```

PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

LAYER		MINIMUM	BEST	MAXIMUM
RHO	1	42.682	60.060	80.696
	2	244.228	292.606	386.935
	3	4.404	5.514	6.755
THICK	1	21.378	37.386	63.826
	2	319.637	342.869	365.184
DEPTH	1	21.378	37.386	63.826
	2	376.715	380.255	386.562
and a second				



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## 4.15 WEKIVA SPRINGS ROAD (SITE 14)

## Location and Geoelectric Section

The detailed location map of this sounding is shown in Figure 4-15.1. The TDEM data and the inverted geoelectric section is shown in Figure 4-15.2 and consists of a three layer section.

#### Depth of Occurrence of Saline Water

The lowest resistivity encountered within the effective depth of exploration of the measurement is 8.7 ohm-m, and it is first encountered at an elevation of 927 ft below msl. Assuming an average porosity of 25% the chloride concentration below that depth is expected to exceed 3,540 mg/l.

## Depth of Occurrence of 250 mg/l Isochlor

The upper 142 ft thick layer with a resistivity of 23 ohm-m is assumed to represent surficial sands and the Hawthorn Group, but must also include a portion of the upper carbonate rocks, since Toth et al (1989) reported the top of the Floridan in this area to occur at about 75 ft below mean sea level. The second layer with a resistivity of 49 ohm-m between elevations 108 ft below msl and 927 ft below msl is characteristic of the carbonate rocks of the Upper Floridan aquifer at this site. Assuming a porosity of 25% a resistivity of 49 ohm-m corresponds to a chloride concentration in excess of 250 mg/l for that entire interval. This is not consistent with chloride contents between 100 mg/l and 250 mg/l reported by Toth et al (1989) for this general area.

Reasons for this apparent discrepancy might be:

- in this area it may not be valid to extrapolate regional information to a local measurement, because of lateral variation; also, changes might have occurred over the five years since the sampling by Toth (1989);
- ii) the average porosity of the carbonate rocks may be different than the 25% value assumed;
- iii) the chemical composition of ground water may be different from that assumed.

#### Accuracy of Measurement and Interpretation

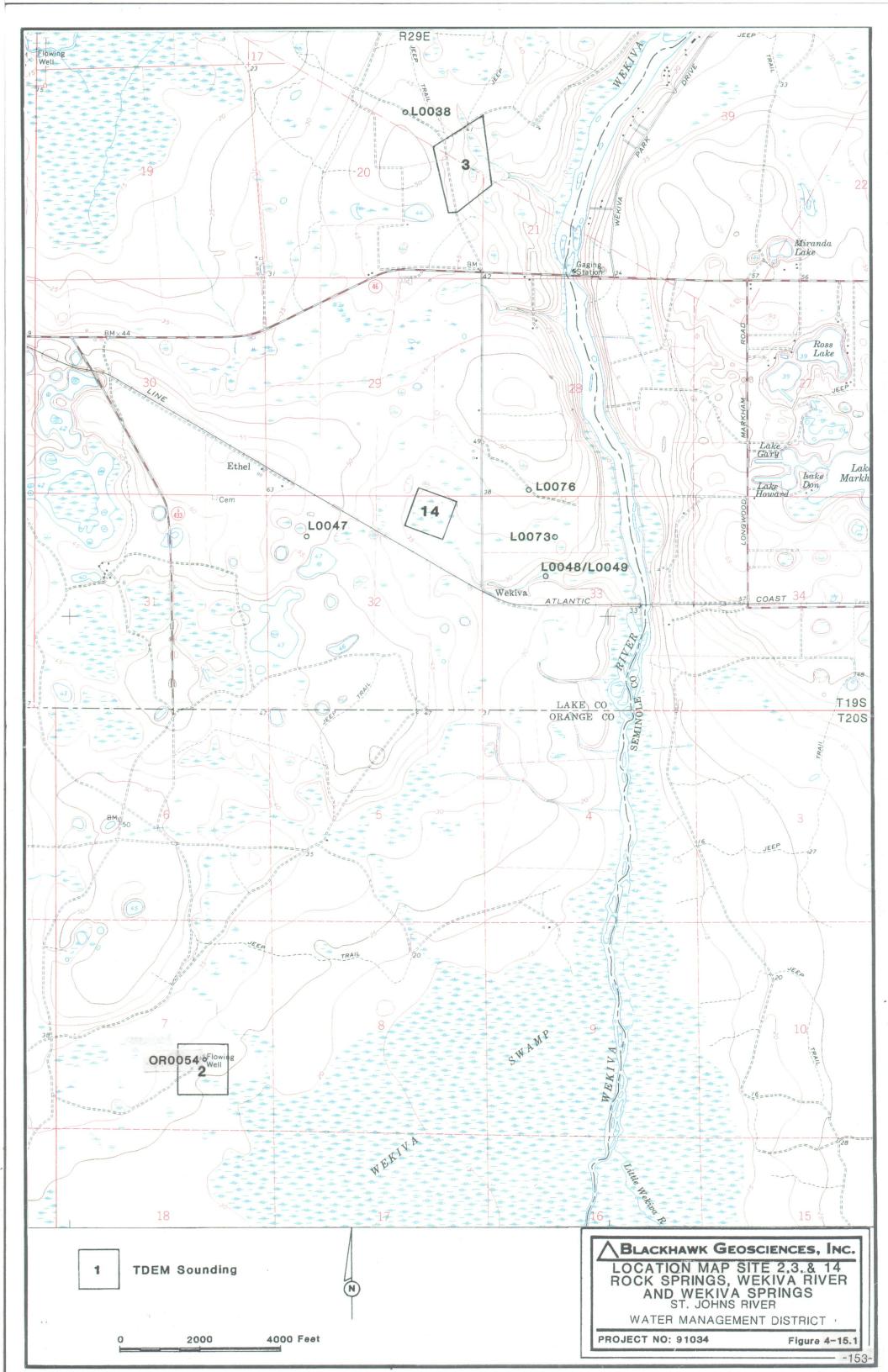
Figure 4-15.3 shows the evaluation of equivalence of the TDEM sounding at this site, and the inversion table (Table 4-14.1) lists the upper and lower bounds of the parameters of the geoelectric section.

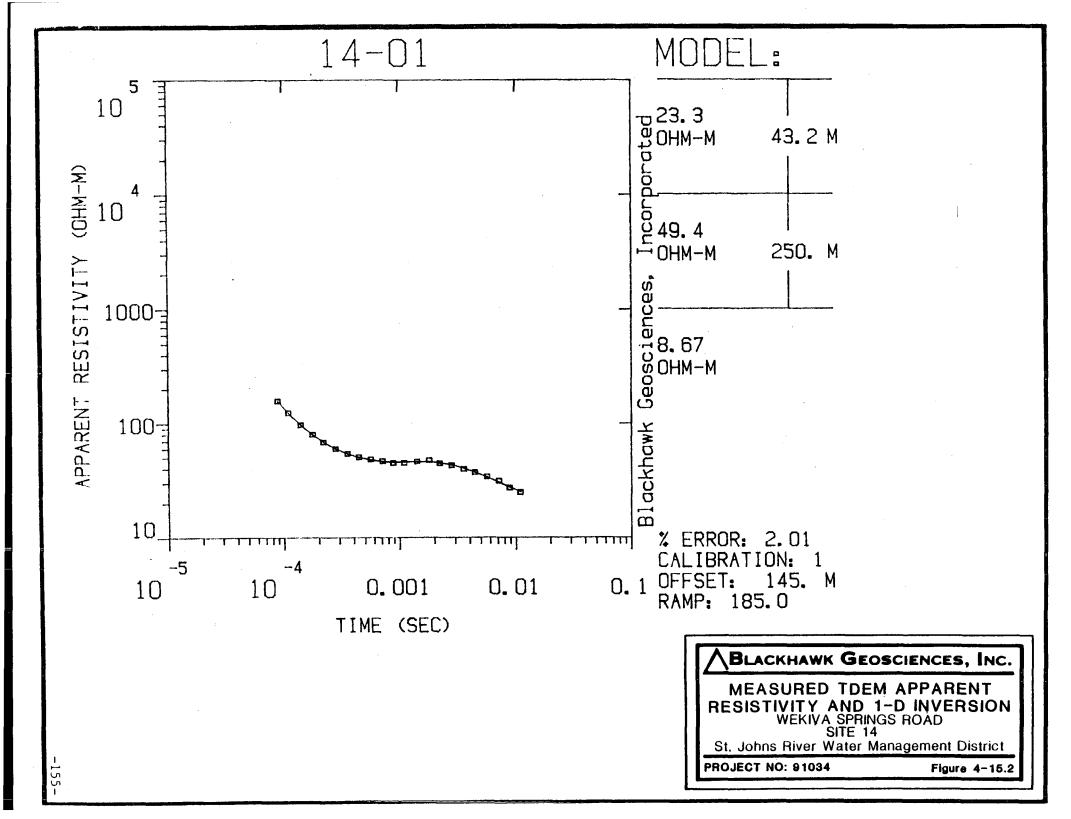
The range of equivalence in determining the depth to the saline layer is  $\pm$  5% of total depth, or about  $\pm$  50 ft. The equivalence in the resistivity of the second layer, representing the main section of the Upper Floridan aquifer, is between 45 ohm-m and 55 ohm-m. Assuming a porosity of 25% chloride concentrations greater than 250 mg/l are expected for that range of resistivity.

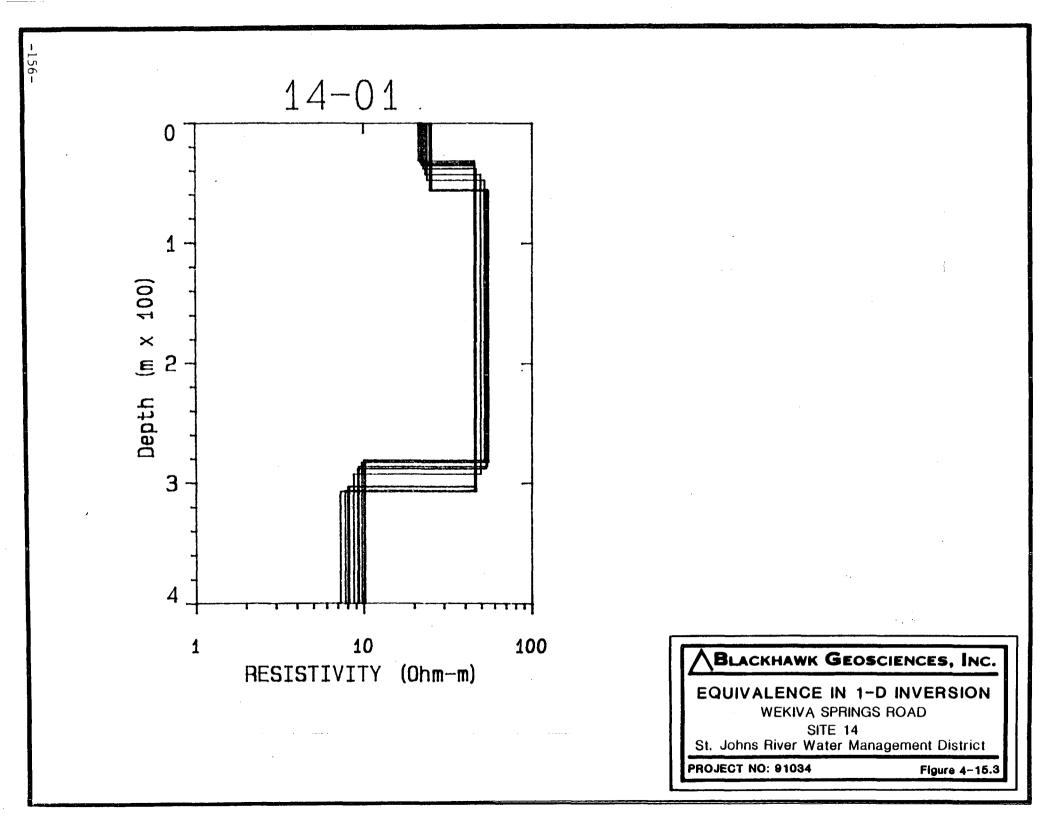
## Summary of Results of TDEM Measurements at Wekiva Springs Road (Site 14)

From the TDEM measurements at Wekiva Springs Road, the following information about water quality was derived:

 The depth of occurrence of saline water was interpreted to occur at an elevation of 927 ft below msl, and below that elevation chloride concentrations in excess of 3,540 mg/l are expected.





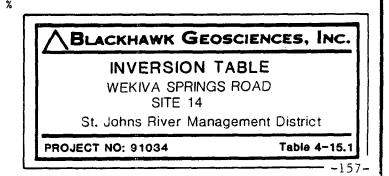


MODEL: 3 LAYERS

RE	RESISTIVITY THICKNESS		ELEVATION		CONDUCTANCE	(S)
	(OHM-M)	(M)	(M)	(FEET)	LAYER	TOTAL
			10.1	33.0		
	23.35	43.2	-33.1	-108.6	1.8	1.8
	49.41	249.6	-282.6	-927.3	5.1	6.9
	8.67					
	TIMES	DATA	CALC	% ERROR	STD ERR	
1	8.90E-05	1.57E+02	1.57E+02	-0.043		
2	1.10E-04	1.24E+02	1.24E+02	0.532		
3	1.40E-04	9.72E+01	9.72E+01	0.006		
4	1.77E-04	7.98E+01	7.98E+01	-0.003		•
5	2.20E-04	6.82E+01	6.86E+01	-0.517		
6	2.80E-04	5.96E+01	5.99E+01	-0.377		
7	3.55E-04	5.40E+01	5.40E+01	-0.023		
8	4.43E-04	5.04E+01	5.02E+01	0.350		
9	5.64E-04	4.81E+01	4.76E+01	1.004		
10	7.13E-04	4.66E+01	4.61E+01	0.939		
11	8.81E-04	4.48E+01	4.56E+01	-1.846		
12	1.10E-03	4.48E+01	4.55E+01	-1.629		
13	1.41E-03	4.60E+01	4.57E+01	0.542		
14	1.80E-03	4.72E+01	4.56E+01	3.425		
15	2.21E-03	4.45E+01	4.49E+01	-0.908		
16	2.80E-03	4.26E+01	4.30E+01	-1.023		
17	3.55E-03	3.95E+01	4.02E+01	-1.720		
18	4.43E-03	3.69E+01	3.71E+01	-0.421		
19	5.64E-03	3.38E+01	3.35E+01	0.927		
20	7.13E-03	3.12E+01	3.02E+01	3.158		
21	8.81E-03	2.72E+01	2.75E+01	-1.241		
22	1.10E-02	2.48E+01	2.50E+01	-0.971		

R: 145. X: 0. Y: 145. DL: 289. REQ: 161. CF: 1.0000 CLHZ ARRAY, 22 DATA POINTS, RAMP: 185.0 MICROSEC, DATA: 14-01

RMS LOG ERROR: 8.63E-03, ANTILOG YIELDS 2.0067 % LATE TIME PARAMETERS



\* Blackhawk Geosciences, Incorporated \*

PI	ARA	METER	RESOLU	JTION M	ATRIX:	
պ	F 11	MEANS	FIXED	PARAME	TER	
Ρ	1	1.00				
P	2	0.00	1.00			
Ρ	3	0.00	0.00	1.00		
T	1	0.00	0.00	0.00	0.99	
T	2	0.00	0.00	0.00	0.00	1.00
		Р 1	Р 2	Р 3	т 1	τ2

## PARAMETER BOUNDS FROM EQUIVALENCE ANALYSIS

2

LAYER		MINIMUM	BEST	MAXIMUM	
RHO	1	21.315	23.349	25.339	
	2	45.471	49.407	55.172	
	3	7.224	8.669	10.183	
THICK	1	31.788	43.156	56.815	
	2	225.310	249.552	271.430	
		,			
DEPTH	1	31.788	43.156	56.815	
	2	280.972	292.708	307.303	



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2) In the Floridan aquifer above the interface with saline water at 927 ft below msl and below the Hawthorn Group, the chloride concentration of ground water is predicted to be in excess of 250 mg/l. This is not consistent with regional water quality reported by Toth et al (1989) for that area.

## 5.0 SUMMARY AND CONCLUSIONS

A time domain electromagnetic (TDEM) survey was performed at 14 sites in the St. Johns River Water Management District during the month of August 1991. Several aspects of the survey are summarized below.

From a TDEM survey the resistivity layering (geoelectric section) of the subsurface is derived. To correlate the resistivity values measured to water quality certain manipulations and assumptions were made, such as:

- (i) two factors mainly influence the formation resistivity of the Floridan aquifer - concentration of dissolved solids and porosity. These two factors cannot be separated from surface electrical measurements. To derive water quality a realistic range of porosity for the Upper Floridan aquifer needed to be assumed. A porosity of 25% was consistently used for the carbonate rocks of the Upper Floridan aquifer at all 14 sites (NW Florida Water Management District, 1983).
- (ii) the relation between formational resistivity and chloride content also is influenced by chemical composition of ground water. In correlating resistivity to chloride content the relationship between these two parameters observed by Kwader (1986) in Seminole County, was assumed to be valid throughout the SJRWMD.

Resistivity values not only depend on ionic concentration and porosity, but also on lithology, - particularly clay content. The lithology of the carbonate rocks of the Upper Floridan are expected to be consistent, but major variation in clay content for the Hawthorn Group and younger sediments overlaying the Floridan aquifer may exist. For that reason water quality was inferred for the carbonate rocks below the Hawthorn Group and younger sediments.

At some sites no resistivity boundary is obtained at the interface between the Hawthorn Group and the carbonate rocks of the Floridan aquifer, so that the upper resistivity layer derived includes surficial sediments, the Hawthorn Group, and parts of the rocks of the Upper Floridan. In those situations it is difficult to evaluate the influence of clay layers on the resistivity value measured, and at some sites no information about chloride concentration can be inferred.

## Determining the Depth of the Interface Between Fresh Water and Ground Water of High Chloride Concentration (greater than 1,000 mg/1

Ground water with a chloride content greater than 1,000 mg/l is characterized in the Floridan aquifer by resistivities less than 28 ohm-m when the aquifer has a porosity of about 25%. A layer with a resistivity less than 28 ohm-m was detected at 13 of the 14 sites at depths varying from about 200 ft to 2,400 ft. It was not detected at the site in the Wekiva Springs State Park (Site 1). For this site computer modeling studies showed that the minimum depth a layer with resistivities characteristic of chloride concentrations greater than 1,000 mg/l (28 ohm-m) could occur was about 1,800 ft.

At one of the 14 sites, Astronaut High School, a nearby well penetrated the interface of high chloride concentrations. The interpretation of the depth to the interface derived from the TDEM data and that observed in the well closely correspond. At other sites the depth to the interface was consistent with regional information, such as that published by Tibbals (1990).

# Water Quality in the Upper Floridan Aquifer and Depth of Occurrence of 250 mg/l Isochlor

Deriving water quality in the Upper Floridan aquifer above the interface with ground water of high salinity (greater than 1,000 mg/l chloride) at several sites is complicated by the possibility that the average resistivity measured may be influenced by clay stringers in the Hawthorn Group, and that influence must be evaluated for each site. This complication arises from the fact that sometimes no resistivity contrast is observed at the interface between the Hawthorn Group and Upper Floridan aquifer.

Assuming a porosity of 25% and validity of the relation between chloride concentration and formation resistivity, ground water with chloride concentrations less than 250 mg/l are expected at formation resistivities greater than 80 ohm-m. At some sites formation resistivities in the carbonate rocks of the Upper Floridan aquifer were below 80 ohm-m for the entire depth interval above the interface with highly saline water. In those cases chloride concentrations greater than 250 mg/l are expected for the entire Upper Floridan aquifer. At other sites resistivities greater than 80 ohm-m were determined for the upper portion of the aquifer. Often a boundary was observed between resistivities greater than 80 ohm-m and much less than 80 ohm-m at some depths. In those situations the depth (or elevation) of the resistivity boundary was taken to correspond with the depth of occurrence of the 250 mg/l isochlor. It is common in aquifers to have a transition zone over which water quality gradually increases with depth. Such transition zones are not mapped with TDEM, and the 250 mg/l isochlor likely occurs some distance above the resistivity boundary.

## 6.0 REFERENCES

Kwader, Thomas 1986. The use of geophysical logs for determining formation water quality, Ground Water, V. 24, No. 1, pg. 11-15.

Miller, J.A., 1990. Ground water atlas of the United States, Segment 6, Alabama, Florida, Georgia, and South Carolina, USGS Atlas 730-G.

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Tibbals, C.H., 1990. Hydrology of the Floridan aquifer system in east-central Florida: USGS Professional Paper 1403-E, 98 pg.

Toth, D.J., 1988. Salt water intrusion in coastal areas of Volusia, Brevard and Indian River Counties: St. Johns River Water Management District Technical Publications, SJ 88-1.

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