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TIME DOMAIN ELECTROMAGNETIC SOUNDINGS AND ANALYSIS ST. JOHNS RIVER WATER MANAGEMENT DISTRICT NORTHEAST FLORIDA/SOUTHEAST GEORGIA NOVEMBER 1993

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This is to certify that I, Robert J. Windschauer, have reviewed the figures, tables, and text of the following report.

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SEAL

EXECUTIVE SUMMARY

A time domain electromagnetic (TDEM) survey was performed at 19 sites in the St. Johns River Water Management District and one site in Georgia during the months of April and May, 1993. The TDEM method is a geophysical technique which, through ground surface based measurement, enables description of the vertical distribution (one-dimensional depth layering) of formation electrical resistivity. As such, TDEM soundings provide a gross approximation of an electrical log as performed in a borehole without the significant expense of drilling, completing, and logging such a borehole. In comparing TDEM soundings to electric logs, the minimum thickness of an interval that can be resolved by TDEM is several orders of magnitude larger than what can be resolved by electric logs. The confidence in the conclusions from TDEM findings can be enhanced when water quality information from nearby wells is available. The objective of the TDEM survey was to determine the depths to the 250 mg/L and 5,000 mg/L isochlors.

The determination of the depth to the 5,000 mg/L isochlor was made at 19 of 20 sites. Depths ranged from 413 to 2,427 feet (ft) below land surface (bls). Only one sounding (Site 2--Cumberland Island), failed to detect a low-resistivity basal layer. A forward-modeling sensitivity analysis estimated a minimum depth of 2,200 ft mean sea level (msl) for such a layer. Well data which provided an estimated depth to the 5,000 mg/L isochlor in nearby areas was available at 4 sites. Results from the TDEM soundings reasonably agreed with well data at three of the four sites. The reasons that the results did not agree at the fourth site (Site 8--Union Camp) are not known.

The determination of the depth to the 250 mg/l isochlor was made at 6 of 20 sites. At twelve of the sites (Sites 5, 8, 9, 11, 12, 13, 14, 15, 16, 18, 19, 20) the 250 mg/L isochlor could not be determined because the geoelectric model for the site could not distinguish the Hawthorn Group and surficial sediments from the Floridan aquifer. At one site (Site 2 -Cumberland Island) the depth of investigation was not sufficient to determine the depth to the 250 mg/L isochlor. Accordingly, the assumptions used in the empirical relationships to determine the 250 mg/L isochlor were not valid. At several sites, water quality inferred from TDEM formation resistivities did not agree with water quality results from nearby wells. This discrepancy is likely due to ground water chemistry in these areas not meeting the assumptions of the empirical relationships for the determination of chloride concentration.

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1.0 INTRODUCTION

The St. Johns River Water Management District (SJRWMD) has contracted with Subsurface Detection Investigations, Inc. (SDII) to perform a series of Time Domain Electromagnetic (TDEM) survey measurements in northeast Florida and southeast Georgia during the time period April to May 1993. This latest series of TDEM soundings is a continuation of similar TDEM programs funded by SJRWMD in previous years (Blackhawk, 1990 and CEES, 1992). The TDEM method is a geophysical technique which, through ground surface-based measurement, enables description of the vertical distribution (onedimensional depth layering) of formation electrical resistivity. As such, TDEM soundings provide a gross approximation of an electrical log as performed in a borehole without the significant expense of drilling, completing, and logging such a borehole. In comparing TDEM soundings to electric logs, the minimum thickness of an interval that can be resolved by TDEM is several orders of magnitude larger than what can be resolved by electric logs. As formation resistivity is a direct function of formation lithology, porosity, and pore fluid conductivity, in situ determination of formation resistivity offers a means of inferring the water quality within given formations through empirical relationships between assumed porosity, pore-water chloride concentration, and the measured value of resistivity.

Given this background, SJRWMD has set the objectives of this TDEM survey as:

- 1. determination of the depth to the saltwater interface (water with chloride concentration greater than 5,000 milligrams per liter [mg/L]);
- 2. determination of the depth within the aquifer (above the saltwater interface) at which chloride concentration of pore waters equals 250 mg/L;
- estimation of the chloride content of the saltwater layer assuming values of 25,
 30, and 35 percent for porosity of that layer.

The principal strength of TDEM is the detection and mapping of depths to the top of a conductive layer within an otherwise resistive medium. As such, the first objective (chlorides greater than 5,000 mg/L) is the easiest to accomplish and is the best resolved. Determination of the second and third objectives relies on empirical relationships derived from studies of wells in Seminole County (in east-central Florida) and, therefore, is a less certain and less well-resolved determination.

This report details the field procedures, data quality control and analyses procedures from a total of 20 sites as selected by SJRWMD personnel. Of these, 19 sites are in northeastern and east central Florida, roughly from Orlando to slightly north of Jacksonville. The remaining site was located just offshore of mainland Georgia (Cumberland Island). Figure 1-1 presents the locations for the 20 TDEM sites.



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2.0 HYDROGEOLOGIC SETTING

Ground water is drawn from three principal aquifers within SJRWMD (Figure 2-1). The three principal aquifers are the surficial aquifer system, the intermediate aquifer system and the Floridan Aquifer System (Scott et al., 1991). The surficial aquifer consists primarily of Upper Miocene to Holocene age consolidated to poorly indurated siliclastic sediments (Scott et al., 1991). Permeable interbeds within these sediments are locally significant sources of potable water near coastal areas and within Seminole, western Clay, and Alachua counties (Fernald and Patton, 1985).

The Miocene-age Hawthorn Group separates the surficial aquifer from the Floridan aquifer and creates confining conditions within the Floridan aquifer. The intermediate aquifer system is comprised of high-transmissivity zones within the Hawthorn Group (Figure 2-1). Typically these high-transmissivity zones occur within sandy phosphatic limestone beds. The intermediate aquifer system is a significant source of potable water in southeastern Flagler and eastern Orange counties (Fernald and Patton, 1985). The Hawthorn Group is thin or absent in the area of TDEM sites 1, 15, 16, 17, 18, 19, and 20.

The primary source of potable water throughout the majority of the SJRWMD is the Floridan aquifer. The Floridan aquifer is composed of (from oldest to youngest) the Cedar Keys Formation, Oldsmar Formation, Avon Park Formation, Ocala Limestone (where present), the Suwannee Limestone and the lower formations of the Hawthorn Group (where present; Figure 2-1; Scott et al., 1991). The ages of these formations range from Paleocene to Miocene.

The Floridan aquifer is subdivided into the Upper and Lower Floridan aquifer by a middle semi-confining unit ranging in thickness from nearly 0 to over 1,000 feet (ft). The middle semi-confining unit is leaky and the hydraulic connection between the Upper and Lower Floridan aquifers is variable (Tibbals, 1990). Depth to the division ranges from approximately 300 to 1,200 ft below mean sea level within SJRWMD (Miller, 1986).

The Ocala Limestone is the most productive aquifer within the Floridan aquifer. Along the east coast and southern portion of SJRWMD, the Cedar Keys or Oldsmar Formations typically contain salt water. In the area of TDEM sites 2 and 3, the Fernandina permeable zone is present within the Lower Floridan aquifer in the Oldsmar and Cedar Keys formations at a depth of greater than 1,900 ft below mean sea level. The Fernandina permeable zone is locally cavernous and fractured and is suspected to affect water quality in the Upper Floridan aquifer along the coast of the northeast portion of SJRWMD (Miller, 1986).

Chloride concentrations within the Upper Floridan aquifer are usually less than 50 mg/L in the northern and west central portions of SJRWMD and exceed 250 mg/L in the east central and southern portions of SJRWMD (Fernald and Patton, 1985). Areas of mineralized water in the Floridan aquifer are present within the central and southern portion of SJRWMD. Sources of mineralized water include lateral seawater intrusion, seawater upwelling, and connate water (Scott et al., 1991).

	LITHOSTR UN	ATIGRAPHIC NT	HYDROSTRATIG UNIT	RAPHIC		
F	UNDIFFERENTIATED PLEISTOCENE—HOLOCENE SEDIMENTS ANASTASIA FORMATION CYPRESSHEAD FORMATION NASHUA FORMATION		SURFICIAL AQUIFER SYSTEM			
	HAWTHORN GROUP STATENVILLE FORMATION COOSAWHATCHIE FM. MARKSHEAD FORMATION PENNEY FARMS FM. SUWANNEE LIMESTONE OCALA LIMESTONE AVON PARK FORMATION OLDSMAR FORMATION CEDAR KEYS FORMATION		INTERMEDIATE A SYSTEM (CONFINING (QUIFER DR JNIT		
			FLORIDAN AQUIFER SYSTEM			
A						
C						
UNDIFFEI		RENTIATED	CONFINING UNIT			
		SDII	LITHOS HYDROS	TRATIGRAPH	IIC AND IIC UNITS	<u>,,,,,</u> ,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,

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3.0 FIELD ACQUISITION PARAMETERS, EQUIPMENT, AND DATA PROCESSING

3.1 FIELD ACQUISITION PARAMETERS

Twenty sites were selected by SJRWMD for TDEM soundings. The TDEM method involves the laying of 12 gauge AWG wire in an approximately square or rectangular loop on the ground over a large area (on the order of 10^6 ft² or greater). This is the transmitter, or Tx loop. The Tx loop is energized by a bi-polar electrical current (up to a maximum of 30 amperes). The response of the ground is sensed by a centrally located (midpoint of the Tx loop) search coil (receiver, or Rx coil). The transient response seen by the receiver is recorded digitally by the data-logging module.

To attain the depth of exploration required to determine the depth to the saltwater interface within SJRWMD, Tx loop sizes ranging from 1,000 ft x 1,000 ft up to 1,500 ft x 1,500 ft were employed where possible. Tx loop sizes at individual sites were prescribed by SJRWMD personnel and adjusted in the field to accommodate field logistical constraints such as obvious metal structures, power lines, or limited areas of access. Tx loops were laid out using premarked cables and a compass. Loop dimensions, transmitter currents, and other site-specific information are included in the individual descriptions of the sounding results (Section 5.0).

In addition to the main sounding data set at a given site, SDII also collected quality control (QC) sounding data using an off-center Rx coil location. That is, if there was an obvious, possible source of noise (pipeline or power line, for example) to one side of a Tx loop, then the coupling of the incident pulse from the transmitter with that possible noise source would impart voltage gradients within the loop that would not exist otherwise. In the absence of noise sources, the voltage measured in the loop is very well behaved; it does not vary much with position of the Rx coil. To check for possible interference sources, several soundings are performed 10-15 percent of the Tx loop length away from the initial Rx coil location. It can be shown that the EMF remains relatively flat to about 10 percent L (L being the length of one of the sides of the Tx loop) off center (Blackhawk, 1990). If a shallow noise source is affecting the data quality, it would impose a higher EMF gradient

in one or more directions off center from the Tx loop. In Figure 3-1 examples of TDEM data that are; 1) unaffected by induction noise, 2) affected by induction noise (as from buried metal pipelines), and 3) affected by powerlines are provided. None of the TDEM sites surveyed during the SDII investigation appeared to have been affected by noise sources.

QC measurements were generally performed at four different locations about the loop center. If the data from the off-center Rx location matches the central-loop data, then the data are not noise-affected. If they diverge significantly, the data are noise-affected and should not be used.

The SDII field crew consisted of one senior project geophysicist, Michael Wightman, P.G., assisted by two geophysical field technicians. During the initial phase of the project, Dr. Thomas L. Dobecki, SDII principal geophysicist, complemented the field crew to ensure survey program objectives were being met by reviewing the field procedures, instrument settings, and resulting data. A representative of SJRWMD, Dr. David Toth, was also present in the field. Table 3-1 summarizes the daily field activities.



Table 3-1. Daily log of field activities

Date (1993) Activity

- April 26 Read Daytona Beach Speedway (Site 1) EM37 TDEM sounding.
- April 27 Read Cumberland Island (Site 2) EM37 TDEM sounding.
- April 28 Read Nassau County (Site 3) EM37 TDEM sounding.
- April 29 Read St. Augustine #1 (Site 4) EM37 TDEM sounding.
- April 30 Read St. Augustine #2 (Site 5) EM37 TDEM sounding.
- April 30 Read Picolata (Site 6) EM37 TDEM sounding.
- May 1 Read Green Cove Springs (Site 7) EM37 TDEM sounding.
- May 2 Read Union Camp (Site 8) EM37 TDEM sounding.
- May 3 Read Drayton Island (Site 9) EM37 TDEM sounding.
- May 4 Read Bear Island (Site 10) EM37 TDEM sounding.
- May 5 Read Deseret #1 (Site 11) and Deseret #2 (Site 12) EM37 TDEM sounding.
- May 6 Read University of Central Florida (Site 13) EM37 TDEM sounding.
- May 7 Read Richland Properties (Site 14) EM37 TDEM sounding.
- May 8 Read New Smyrna Beach (Site 15) EM37 TDEM sounding.
- May 9 Read Lake Ashby (Site 16) EM37 TDEM sounding.
- May 10 Read Lake Helen (Site 17) and Deltona (Site 18) EM37 TDEM sounding.
- May 11 Read Blue Springs State Park (Site 19) EM37 TDEM sounding.
- May 12 Read Site 17 (re-done) and De Land (Site 20) EM37 TDEM sounding.

3.2 EQUIPMENT

SDII employed the Geonics EM37 TDEM system to accomplish the tasks of this project. The principal components of the EM37 are:

- Transmitter (Tx) loop (variable length 12 gauge AWG wire, insulated)
- Gasoline power generator/EM37 transmitter box (maximum 30 ampere, bipolar square wave)
- Receiver (Rx) coil (100 square meter effective area)
- EM37 Receiver Module (system control and parameter selection)
- Polycorder digital notebook (data storage)

A block diagram of the field setup of the system is given in Figure 3-2. Once setup is completed, a current waveform as depicted by Figure 3-3 is injected into the Tx loop. The rapid turn-on and turn-off of current in the loop creates a strong EMF which interacts with earth and man-made materials to generate eddy currents within conductive materials. These currents have an associated secondary magnetic field which is detected by the Rx coil as shown on Figure 3-3. Eddy currents close to the Tx coil are induced first and decay below detection limits before deeper currents. Currents in resistive materials also decay faster than currents in conductors. Deeper conductors contribute to responses at later times at the Rx coil than do shallower subsurface features. Thus, by measuring the rate and nature of the decaying magnetic field seen by the Rx coil after Tx shutoff, the distribution of subsurface resistivity can be determined.





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The survey variables that can be selected by the TDEM operator are the size of the Tx coil, Tx coil current (which controls the penetration depth), analog stacking (number of repetitions of summed tests in order to increase signal-to-noise ratio), gain at the receiver, and repetition rate (frequency) of the current cycles. For this investigation SDII used three different frequencies (3 Hz, 7.5 Hz, and 30 Hz) to acquire detailed and overlapping segments of the decay curve which enabled resolution of shallow (30 Hz data) and deeper (3, 7.5 Hz data) portions of the subsurface.

3.3 DATA PROCESSING

Data acquired by the EM37 were recorded by the Polycorder digital notebook logger and downloaded to a portable computer for data editing, processing, and interpretation (inversion). The primary software program used to process the data was TEMIXGL (Interpex, Ltd.). This program accepts raw data from the Polycorder and proceeds through the following general processing steps:

Data Edit - Allows for modification of survey description information, for example, loop size, Tx coil amperage, which may have been entered improperly. Decay curves for all frequencies and gain values taken at a site are displayed; suspect data points can be deleted and the individual curves for different frequencies and gains are averaged and converted to a single, apparent resistivity versus time (after Tx turn-off) field curve (see Figure 3-4, for an example of voltage data and apparent resistivity versus time curves).

The field curve is comprised of 30 data points, where each data point represents an apparent voltage collected at a particular time or time gate. Each frequency has 20 time gates and each frequency overlaps the proceeding or preceding frequency by 10 time gates. Combining data collected at the 30 Hz and 3 Hz frequency produces one sounding curve with 30 time gates, with an overlap between time gates 10 through 20. Data collected at 7.5 Hz provides apparent resistivity values for time gates 5 through 25. An advantage of using 30, 7.5, and 3 Hz frequencies for all the soundings is that different gains can be used for each frequency. Lower gains can be used at a frequency of 30 Hz to avoid saturating early channels, and higher gains can be used at 3 Hz to amplify weaker signals in later channels. The combined data is interpreted as one sounding curve.

<u>Initial Model</u> - Review of the apparent resistivity curve shape allows a trained geophysicist to make an initial guess as to the true resistivity versus depth (layered) model which would produce the observed data set. After such a model is created, a field curve is calculated from the model and compared with the observed data. The degree of agreement between model and field data is measured statistically and expressed as the fitting error. The geophysicist may then, in an interactive mode, adjust the model to obtain a better fit or can modify the starting model.

As part of the modeling procedure early and late time data is commonly discarded. Typically, apparent resistivity values collected at early times are discarded because the data collected at these times is often not representative of geological conditions because of the affect of the Tx coil shut off not being truly instantaneous. In the final modeling of this data, in may appear that the model curve passes through several of these early time points, but not all the points. In such a case, all the early time data points are discarded because it is not good modeling practice to delete data points from the middle portion of a curve and utilize data points preceding them. Often, later time data is also not representative of geological conditions because the primary EMF field strength has been too dissipated to provide a representative apparent resistivity value. Suspect late time data is also discarded. Poorly fitting data points are marked with a "x", utilized data points are marked with a square (Figure 3-4).

Modeled curves quite often demonstrate an upward curvature during early times. This upward curvature is usually due the TDEM response not following theoretical behavior or the affect of the Tx coil shut off not being truly instantaneous. This deviation produces a distortion, however, this distortion has little or no affect on the results from the TDEM survey when the target depth is several hundred feet below land surface.

<u>Automatic Inversion</u> - Based upon the initial model, the program will attempt to create a better fit to the observed data using an iterative, Inman Ridge Regression routine to adjust layer thicknesses and resistivities until a minimum error of fit is realized; our goal was to produce models which fit the observed data within a 5% error of fit. This final model is termed the "best fit" model (see Figure 3-5). Only the data points utilized in the determination of the modeled curve are used in calculation of the fitting error.

Equivalence Analysis - Electrical resistivity methods are, as with other geophysical methods, plagued by the so-called "non-uniqueness" problem. That is, while a best-fit model produces an acceptable fit to field data curves, there are several other models having different thicknesses and resistivities which will also provide a "reasonable" fit to the same data. TEMIXGL will produce a suite of models, using the best-fit model as a start, which would produce a reasonably close fit (see Figure 3-6). If the equivalence model segments (layers and resistivities) are tightly constrained then the layering provided by the best-fit model is very good. Those parts of the equivalence models that scatter quite a bit around the best-fit model show less confidence in the absolute values of layer thickness and resistivity. A poorly constrained equivalence model for a given layer means either there are too few data points in the raw data to adequately describe that layer or the data is just not very sensitive to that specific layer.

It is important to note that the interpretations resulting from the TDEM data are, specifically, one-dimensional models of layer thickness and layer resistivity. That is, if the earth subsurface is not, effectively, a one-dimensional horizontal layer, then the produced model may have inherent error. Also, the depths to levels of chloride concentration and not resistivity rely on empirical relationships between resistivity and chloride concentration. This latter point will be detailed further in Section 4.0.







4.0 TECHNICAL APPROACH TO SATISFYING SURVEY OBJECTIVES

4.1 GENERAL

As stated previously, the final product of the *geophysical* investigation is a best-fit, one-dimensional model of layer resistivity versus depth. To satisfy the requirements of the survey, these models must be correlated with models of chloride concentration versus depth. Specifically, the resistivity structure must be viewed in terms of determining the depth of occurrence of the 250 mg/L isochlor and the depth to salt water as defined by the 5,000 mg/L isochlor. To ensure that the results from the 1993 TDEM survey are directly comparable to and compatible with the results of TDEM surveys performed in previous years (Blackhawk, 1990 and CEES, 1992), SDII will utilize the identical relationships between resistivity and isochlor depths for the Floridan aquifer. These relationships and assumptions are detailed in the following sections. However, it must be realized that correlations of TDEM-derived layer conductivities with specific chloride values are approximate and based on several simplifying assumptions.

4.2 CORRELATION OF INVERTED GEOELECTRICAL (RESISTIVITY) PROFILES TO CHLORIDE CONCENTRATIONS

In previous studies, it was presumed that the depth to salt water was such that this interface was inferred to occur within the Floridan aquifer system. The only noted exceptions to this were soundings in the area of Jacksonville where the great depth (>2,000 ft) and the very low resistivity (< 2 ohm-m) of the deep, low resistivity layer placed the interface below the Lower Floridan aquifer (CEES, 1992). For such deep sites with very low resistivities, the published relationships between resistivity and chloride concentration cannot be used; it is merely presumed that the chloride concentration at these sites exceeds 5,000 mg/L for the saltwater section.

In cases where the electrical response between the Floridan aquifer and overlying sediments are indistinguishable, the two hydrostratigraphic units must be combined into a single geoelectric layer. Similar to the situation where the interface is below the Floridan aquifer, the published relationships between resistivity and chloride concentration are invalid and the chloride concentration in ground water above the saltwater interface cannot be determined.

For the majority of soundings conducted previously, the saltwater interface positions were "inferred to occur within the Floridan aquifer system" (Blackhawk, 1990 and CEES, 1992) and, therefore, the published relationships are applicable. When the saltwater interface occurred within the Floridan aquifer, the following procedure was used in both this and previous studies (Blackhawk, 1990 and CEES, 1992).

The carbonate rocks of the Floridan aquifer system (as opposed to the highly variable lithologies of overlying formations) are expected to be uniform and, as such, their resistivities are determined principally by, porosity and specific conductance of pore fluids. The governing empirical "law" relating formation resistivity (Ro), fluid resistivity (Rw) and porosity (ϕ) in a clay-free lithology is Archie's Law:

$$\mathbf{F} = \mathbf{Ro}/\mathbf{Rw} = \mathbf{a}\boldsymbol{\phi}^{\mathbf{m}} \tag{1}$$

where F = "formation factor" and "a" and "m" are empirically derived constants which are specific to a given formation in a given area. Previous TDEM reports have used the values of m = 1.6 and a = 1 from Kwader (1982) as being most appropriate for the Floridan aquifer. These values are from studies of wells completed in the Upper Floridan aquifer in Seminole County, Florida.

Kwader (1982) has also established the following relationship from his study of Seminole County wells:

$$Cl = (3500/Rw) - 153$$
 (2)

where Cl is the equivalent chloride concentration in mg/L and Rw is fluid resistivity in ohmmeters. Extrapolating these expressions by Kwader outside of Seminole County presumes that the relative ionic chemistry (especially a chloride/sulfate ratio of 5:1) remains the same or reasonably close to conditions in that area. Significant chemical variation would cause Equation 2 to be, quite likely, invalid.

Because formation resistivity, Ro, is what the geophysical analysis of TDEM data has produced, a combination of equations (1) and (2) allows for determining a functional relationship between chloride concentration, inferred formation resistivity, and porosity:

$$Cl = (3500\phi^{-1.6}/Ro) - 153$$
 (3)

or, for an assumed 25% porosity for the Upper Floridan aquifer as per previous TDEM reports:

$$Cl = (32,163/Ro) - 153$$
 (4)

Linking this relationship to the cited survey objectives, we would expect that a Floridan aquifer with 25% porosity, similar water chemistry (5:1 chloride to sulfate ratio) to the Kwader study, and a 250 mg/L chloride concentration would yield a measured formation resistivity of 80 ohm-m. Higher resistivities than this would indicate fresher water. Chloride concentrations of 5,000 mg/L would correspond to formations resistivities of 6.2 omh-m; higher concentrations would yield lower resistivities. These values, then, are what we should expect to see for the fresh and saltwater sections of the Floridan aquifer.

One final consideration, besides porosity and similar chemical species/ratios, is made by previous reports (Blackhawk, 1990 and CEES, 1992) and, again, will be adhered to in this 1993 study. The relationships cited are for a clearly defined, carbonate section within the Floridan aquifer (i.e., beneath the Hawthorn Group). If there is a clearly defined thickness of the Hawthorn Group from the electrical sounding results and if that thickness is in agreement with published Hawthorn Group thickness for the area of a specific site, then there is presumed to be no affect of the measured formation resistivity for the Floridan aquifer due to interfingering of clay stringers of the Hawthorn Group. This means that the inversion resistivity results representing the Floridan aquifer layer are valid.

4.3 DETERMINATION OF DEPTH TO 250 mg/L AND 5,000 mg/L ISOCHLORS

The previous discussion of the relationship of formation conductivity to chloride content is particularly applicable to geoelectrical measurements made on a fine, highly resolved scale, such as a borehole electrical log, where an almost continuous measure of resistivity versus depth is available. As known from geophysical logs and water quality studies, the saltwater interface is not a knife-edge interface in the subsurface but is a gradational interface. Within the freshwater section, we would also expect the chloride concentration to follow a gradually increasing-downwards distribution. Therefore, the TDEM sounding, which presents the subsurface as a sequence of a few layers of presumed, uniform resistivity, is not an actual representation of the true subsurface but a low resolution version of it. The saltwater interface (chlorides greater than 5,000 mg/L), which exhibits a much higher gradient of chloride concentration than the overlying fresher water, comes closest to being a true interface. This is why depth to the saltwater interface from TDEM should be close to the low resistivity layer detected.

Actual reported depth to the 5,000 mg/L isochlor in previous reports (CEES, 1992) is determined by the contrast in resistivity of the layers above and below the geoelectrical interface. If the contrast is large (e.g., greater than 80 ohm-m above and less than 20 ohm-m below), then the depth to the 5,000 mg/L isochlor is assumed to be 50 ft below the interface depth determined from geoelectrical inversion. If the contrast is small (e.g., a 20-80 ohm-m layer above and less than 20 ohm-m layer above and less than 20 ohm-m layer below), the depth to the 5,000 mg/L isochlor is taken as equal to the depth to interface determined from the geoelectrical inversion. These adjustments are intended to correct for the existence of the transition zone.

The criterion used to define the depth to the 250 mg/L isochlor in previous TDEM surveys for SJRWMD (Blackhawk, 1990 and CEES, 1992) is also a data-based criterion. That is, the final reported position of this isochlor, relative to the boundary between the Floridan aquifer freshwater geoelectrical layer and the saltwater geoelectrical layer depends upon the layer resistivities above and below the interface as determined by the inversion.
Four data classes have been defined based upon a reference value for resistivity of 80 ohm-m for a portion of the Floridan aquifer. We reproduce the following criteria for positioning the 250 mg/L isochlor (CEES, Table 4-2, 1992).

Summarizing Table 4-2 in CEES (1992), if the Floridan freshwater section is in excess of 80 ohm-m while the underlying layer is less than 20 ohm-m (so-called Class A geoelectrical section), then the 250 mg/L isochlor is placed at a position 50 ft higher than the saltwater interface depth defined from geoelectrical inversion.

If the Floridan freshwater section is in excess of 80 ohm-m while the underlying layer is between 20-40 ohm-m (so-called Class B section), then the 250 mg/L isochlor is placed 25 ft above the saltwater interface depth defined from geoelectrical inversion.

If the Floridan freshwater section is in excess of 80 ohm-m and the underlying layer is between 40-80 ohm-m (Class C), then the 250 mg/L isochlor is placed at the interface.

Finally, if there is no contrast (i.e., a uniform layer of > 80 ohm-m; Class D), then we are not seeing an expected saltwater interface within the depth of exploration of the field sounding. Also, there is no detectable/mappable 250 mg/L isochlor.

In the above determinations for the 250 mg/L isochlor, the "depth" to the saltwater interface referred to is the depth to the low resistivity layer taken directly from the TEMIXGL inversion and not the corrected 5,000 mg/L depth as discussed previously.

5.0 RESULTS AND DISCUSSION

5.1 SUMMARY OF RESULTS

A summary of the 1993 TDEM investigation is presented in this section. The summary includes the resulting geoelectrical inversions, 250 mg/L isochlor depth and the 5000 mg/L isochlor depth. More detailed presentation of the individual site results are contained in the following sections 5.2 through 5.21. Each individual site section will present a site description, site map, apparent resistivity versus time (data) curves, the best-fit geoelectrical section with equivalence analysis, and inferred depths to the 5,000 mg/L (salt water) and 250 mg/L isochlors.

Table 5.1-1 lists the 20 sites with summary information describing site number, name, residing county, latitude, longitude and loop size.

Table 5.1-2 summarizes the results of the TEMIXGL geoelectrical inversion section (number of layers, layer thicknesses and resistivities, and range of equivalence models for each layer parameter).

Table 5.1-3 summarizes the estimated chloride content of the saltwater layer assuming porosities of 25, 30, and 35% for the Floridan Aquifer System.

Table 5.1-4 summarizes the interpreted depths to the 250 mg/L and the 5,000 mg/L isochlors at each site based upon the criteria outlined in Section 4.3 and as utilized in TDEM surveys performed for SJRWMD in previous years (Blackhawk, 1990 and CEES, 1992). As in previous years, these calculations are made assuming a 25% porosity for the Floridan Aquifer System and a 5:1 chloride-to-sulfate ratio for the ground water chemistry. The estimated chloride-to-sulfate ratios at each of the sites is provided in Table 5.1-4.

The effect of a CL/SO₄ ratio less than 5:1 would be for waters with equivalent conductivity to have different CL values. SO₄ is less conductive than CL for an equivalent mass volume. If for example the ratio is less than 5:1, it will take a higher conductivity (lower resistivity) to get a 250 mg/L chloride value. That is, for sites where the 5:1 ratio is 1:1, resistivities would have to be less than 80 ohm-m to reach a chloride content of 250 mg/L.

Table 5.1-1	
SUMMARY OF TDEM SITE SURVEY INFORMATION	

Site Number	Site Name	Residing County	Latitude	Longitude	Loop Size (in feet)
1	Daytona Beach Speedway	Volusia	29°10'37"N	81°04'37"W	1000 x 700
2	Cumberland Island	Camden (GA)	30°48'24"N	81°27'22"W	1750 x 1250
3	Nassau County	Nassau County	30°35'40"N	81°41'03"W	1500 x 1500
4	St. Augustine #1	St. Johns	29°54'49"N	81°24'37"W	1000 x 1000
5	St. Augustine #2	St. Johns	29°53'27"N	81°24'47"W	2000 x 500
6	Picolata	St. Johns	29°55'10"N	81° 34'05 "W	1000 x 1000
7	Green Cove Springs	Clay	29°57'09"N	81°39'52"W	1000 x 1000
8	Union Camp	Putnam	29°21'29"N	81°34'10"W	1220 x 940
9	Drayton Island	Putnam	29°22'49"N	81°38'32"W	1380 x 600
10	Bear Island	Flagler	29°26'22"N	81°28°59"W	1500 x 115
11	Deseret #1	Orange	28°25'17"N	81°05'17"W	1500 x 1000
12	Deseret #2	Orange	28°25'17"N	81°08'27"W	1575 x 1065
13	UCF	Orange	28°36'11"N	81°11'15"W	1500 x 1500
14	Richland Properties	Orange	28°21'43"N	81°23'30"W	1500 x 1500
15	New Smyrna Beach	Volusia	29°03'00"N	80°56'13"W	1000 x 900
16	Lake Ashby	Volusia	28°54'02"N	81°03'53"W	1350 x 650
17	Lake Helen	Volusia	29°00'39"N	81°14 ' 39"W	1000 x 1000
18	Deltona	Volusia 28°55'15"N		81°10'09"W	900 x 500
19	Blue Springs State Park	Volusia	28°57'17"N	81°20'02"W	750 x 250
20	De Land	Volusia	29°06'11"N	81°20'39"W	1130 x 1000

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Site Name	Number of Modeled Layers in Geoelectrical Section		Resistivity p1 (ohm-m))		Thickness b, (meters)	,)		Resistivity p2 (ahm-m)	•		Thickness b ₂ (meters)	·		Resistivity p3 (ohm-m	/)}		Thickness h, (meters)	,	To Dec which as Salt	tal Depth pest Condu h is Interpu Water (m	To actor reted acters)*
		Min	Best	Max	Min	Best	Max	Min	Best	Max	Min	Best	Max	Min	Best	Max	Min	Best	Max	Min	Best	Max
1. Daytona Beach Speedway	3	6.7	10.1	11.9	20.5	32.2	40.0	130	215	464	196	206	213	3.4	3.9	4.8				229	238	245
2. Camberland Island	2	14.4	16.1	17.4	64.8	80.2	95.1	53.3	58	63.6	_	-		-	-		-			Not	Seen	
3. Nassau County	3	22.6	23	23.4	132	132	132	943	6010	60,098	559	608	673	2.8	4.7	7.7	-		-	691	740	805
4. St. Augustine #1	3	14.1	14.6	15.2	73	73	73	33.8	44	63.7	91	108	126	9.4	10.3	11.3		-	-	164	181	199
5. St. Augustine #2	3	8.0	95	10.3	34	45	51	118	169	312	118	129	143	17.2	18.4	19.8		-	-	153	174	190
6. Picolata	3	22.4	22.3	23.8	90	90	90	88.7	97.8	121.5	341	488	583	14.0	27.4	49.3	-		-	431	578	673
7. Green Cove Springs	3	26.6	29.0	31.3	9 3	93	93	79	108	167	363	383	402	1.0	23	4.9	-		_	456	476	495
8. Union Camp	3	45	47	49	246	252	258	32	3.7	42					-		-		-	246	252	258
9. Drayton Island	3	22	23	24	123	127	131	22	2.3	25	-		-	-	-				-	123	127	131
10. Bear Island	3	3,4	5.9	6.9	11	23	28	61	- 98	180	134	141	150	25	2.8	3.1	_			158	164	169
11. Descret #1	3	14.7	18.1	20.6	40	56	74	80	101	140	293	329	357	93	10.7	12.7				357	385	408
12. Descret #2	3	7.3	16.5	18.2	19	46	52	538	957	3027	297	314	329	13.5	15.5	20.1	-	-		326	360	376
13. UCF	3	14.4	23.6	30.3	22	38	50	399	534	955	382	395	409	4.7	5.5	6.4		-		424	433	442
14. Richland Properties	3	8.0	16.5	19.5	18	39	47	1007	2551	7112	523	546	590	23	2.8	7.0	_			562	585	628
15. New Smyrna Beach	2	9.1	9.2	9.4	121	126	130	21	23	25	-		-							121	126	130
16. Lako Ashby	2	40.4	41.8	43.5	226	234	242	3.8	4.6	5.6										226	234	242
17. Lake Heien	3	32.7	39.8	47.2	50	66	84	321	497	994	242	258	275	1.8	2.0	23		_	_	324	324	327
18. Deltona	3	65.2	68.4	71.6	296	301	304	1.1	1.4	1.7	_			-			-	_		296	301	304
19. Blue Springs State Park	3	26.8	27.9	29.2	144	151	156	3.9	4.4	5.0				-						144	151	156
20. De Land	3	40.9	42.3	43.8	269	277	283	2.8	3.3	4.0		_		 		_	_		_	269	277	283

Table 5.1-2Summary of Geoelectrical Sections with Range of Equivalence

1 meter equals 3.28 ft

Table 5.1-3Estimated Depths to Salt Water andEstimated Chloride Concentrations at Three Porosities

Site	Formation Resistivity (ohm-m)	Interpreted Depth of Salt water (ft)	Chloride Conc. (mg/L) $\phi = 25\%$	Chloride Conc. (mg/L) $\phi = 30\%$	Chloride Conc. (mg/L) φ = 35%
1 Daytona Beach Speedway	3.9	780	8,094	6,007	4,661
2 Cumberland Island	Not Present	Beyond System Limit			
3 Nassau County	4.7	2,427	6,690	4,959	3,841
4 St. Augustine #1	10.3	594	2,970	2,180	1,670
5 St. Augustine #2	18.4	571	1,595	1,153	867
6 Picolata	27.4	1,896	1,021	724	532
7 Green Cove Springs	2.3	1,562	13,831	10,293	8,010
8 Union Camp	3.7	826	8,540	6,340	4,921
9 Drayton Island	2.3	416	13,831	10,293	8,010
10 Bear Island	2.8	538	11,334	8,428	6,552
11 Deseret #1	10.7	1,263	2,853	2,092	1,602
12 Deseret #2	15:5	1,181	1,922	1,397	1,058
13 UCF	5.5	1,423	5,695	4,215	3,260
14 Richland Properties	2.8	1,919	11,334	8,427	6,552
15 New Smyrna Beach	2.3	413	13,831	10,293	8,010
16 Lake Ashby	4.6	769	6,839	5,070	3,928
17 Lake Helen	2.0	1,064	15,929	11,860	9,234
18 Deltona	1.4	986	22,821	17,008	13,257
19 Blue Springs State Park	4.4	494	7,157	5,307	4,114
20 De Land	3.3	908	9,594	7,127	5,536

	Site	Estimated Chloride-to-Sulfate ratio ^{1/}	Interpreted Depth 5,000 mg/L Isochlor (ft bls)	Interpreted Depth 250 mg/L Isochlor
1	Daytona Beach Speedway	5:1	830	730
2	Cumberland Island	1:5	Not Present	Cannot be Determined
3	Nassau County	1:1	2,477	2,377
4	St. Augustine #1	1:2	594	Not Present
5	St. Augustine #2	1:1	621	Cannot be Determined
6	Picolata	1:5	1,946	1,846
7	Green Cove Springs	1:1	1,612	1,512
8	Union Camp	2:1	826	Cannot be Determined
.9	Drayton Island	2:1	416	Cannot be Determined
10	Bear Island	2:1	588	488
11	Deseret #1	2:1	1,313	Cannot be Determined
12	Deseret #2	1:1	1,231	Cannot be Determined
13	UCF	1:1	1,473	Cannot be Determined
14	Richland Properties	1:1	1,969	Cannot be Determined
15	New Smyrna Beach	5:1	413	Cannot be Determined
16	Lake Ashby	1:1	769	Cannot be Determined
17	Lake Helen	1:1	1,114	1,014
18	Deitona	10:1	986	Cannot be Determined
19	Blue Springs State Park	10:1	494	Cannot be Determined
20	De Land	10:1	908	Cannot be Determined

Table 5.1-4Depth to 5,000 mg/L and 250 mg/L Isochloras Determined by Time Domain Electromagnetics

1/ Chloride-to-sulfate ratios from all sites except 4, 5, 6, 8, 9, 10, 11, and 12 estimated from Sprinkle, 1981. Chlorideto-sulfate ratios for sites 4, 5, and 6 estimated from Spechler and Hampson, 1984. Chloride-to-sulfate ratios for sites 8, 9, 10, 11 and 12 from SJRWMD.

TDEM SITE 1 - DAYTONA BEACH SPEEDWAY SITE

5.2.1 Location Description and Geoelectrical Section

5.2

This site is located in Daytona Beach, Florida, in a grassy parking area located adjacent to and serving the Daytona Beach Speedway facility (Figure 5.2-1). Because of curvature of the track oval (fence line and embankment), additional steel fence lines, and local highways, the planned 1,000 ft square Tx loop was altered to a 700 ft by 1,000 ft rectangular loop. The Hawthorn Group is not present in this area (Scott et al., 1991) and the Lower Floridan aquifer begins at approximately 800 ft below msl (Miller, 1986).

A monitor well (DB-1) is located at Daytona Beach International Airport (Figure 5.2-1) and is approximately one mile from the TDEM site. Lithologic logs indicate the Floridan aquifer begins at 96 ft below land surface (bls). Measured chloride concentrations show fresh (<250 mg/L) water down to approximately 770 ft followed by a rapidly increasing chloride concentration gradient below 800 ft depth.

The resistivity sounding data and best-fit model inversion are presented on Figure 5.2-2. The interpreted geoelectrical section consists of a three-layer subsurface.

5.2.2 Geological Interpretation of Geoelectrical Model

The model created from the TDEM survey compares well to the nearby well information. The modeled top layer is low resistivity (10 ohm-m) and has a thickness of 105 ft (32.2 m). The top layer is interpreted as the surficial aquifer system and the 105 ft depth is interpreted as the top of the Floridan aquifer. The increased resistivity of the second layer (215 ohm-m) is consistent with freshwater saturation (>80 ohm-m) within the Floridan aquifer; the depth to the interpreted low resistivity (saltwater) layer is 780 ft (238 m). The depth to the saltwater interface occurs at or near top of the the Lower Floridan aquifer. Accordingly, the three-layered geoelectrical model at this site appears to represent the near surface sediments, fresh water within the Floridan aquifer, and saltwater saturation, respectively.

5.2.3 Depth to Occurrence of Salt Water

The bottom (third) layer of the geoelectrical model, with a resistivity of 3.9 ohm-m, is interpreted to represent salt water. It occurs at a depth of 780 ft (-758 ft msl) which is near the top of the Lower Floridan aquifer. Because the resistivity of layer 2 (215 ohm-m) is interpreted to represent fresh water within the Floridan aquifer (i.e., is greater than 80 ohm-m), the interpreted depth to the 5,000 mg/L isochlor is 50 ft below the depth of the geoelectrical interface, or at 830 ft depth (-808 ft msl). The resistivity of layer 3 (3.9 ohm-m) corresponds to a chloride content of 8,090 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2.

5.2.4 Depth of Occurrence of the 250 mg/L Isochlor

The resistivity of layer 2, 215 ohm-m, corresponds to a chloride content of less than 50 mg/L, assuming a 25% porosity and the validity and applicability of equation (4) of Section 4.2. Using the criteria established in Section 4.3 (this is a Class A type resistivity distribution), the position of the 250 mg/L isochlor is placed 50 ft above the depth to the low resistivity interface, or at a depth of 730 ft (-708 ft msl). The depth to the 250 mg/L isochlor (730 ft bls) correlates well to water quality results from nearby monitor well DB-1 which places the 250 mg/L isochlor at 754 ft bls (Figure 5.2-4).

5.2.5 Accuracy of Measurement and Interpretation

Figure 5.2-3 is the equivalence analysis at this site, and the inversion table (Table 5.2-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about \pm 8 m (26 ft), which is 3% of the total depth. The resistivity of this layer has a range of from 3.4 - 4.8 ohm-m. This corresponds to a range in interpreted chloride concentration of from 9,300 mg/L to 6,550 mg/L, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence range of the resistivity of layer 2 is from 130 - 465 ohm-m which corresponds to a range of chloride content of from 200 to less than 50 mg/L. The chloride-to-sulfate ratio at the site is 5:1 (Table 5.1-4). Accordingly, the assumptions implicit within equation (4) are valid.

5.2.6 <u>Summary of TDEM Sounding at Daytona Beach Speedway (Site 1)</u>

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 830 ft (-808 ft msl) which is the top of the Lower Floridan aquifer. The chloride content below that depth is inferred to be 8,094 mg/L.
- The ground water within the Floridan aquifer at this site is interpreted to contain an average chloride concentration of less than 50 mg/L using a porosity value of 25% for the Floridan aquifer. The 250 mg/L isochlor is interpreted to be present within the Floridan aquifer at a depth of 730 ft (-708 ft msl). The estimated depth to 250 mg/L isochlor correlates well to the depth of the 250 mg/L isochlor (754 ft) measured in a nearby well.









CLI LOCAT COU PROJ LOOP S COIL S SOUNDI	ENT: ION: NTY: ECT: IZE: LOC: NG CO	SJWRMD DAYTONA BE VOLUSIA CO SALT WATER 213.000 0.000 ORDINATES:	ACH SPEEDW UNTY, FLOI INTERFACI m by m (X), E:	AY RIDA E DETECTION 305.000 m 0.000 m (Y 0.0000 N:	DATE: 04- SOUNDING: 1 ELEVATION: EQUIPMENT: Geo AZIMUTH:) 0.0000	26-93 7.00 m nics PROTEM
N N		FITTIN	G ERROR:	4.091 PE	RCENT	
L #	RESI (0	STIVITY hm-m)	THICKNES: (meters)	5 ELEVAT (mete:	ION CONDU rs) (Sie	CTANCE mens)
1 2 3	1 21	0.13 5.2 3.94	32.19 205.8	7.0 -25.1 -231.0	0 9 3. 0.	17 956
ALL P.	ARAME	TERS ARE F	REE			
PARA	METER	BOUNDS FR	OM EQUIVAI	LENCE ANALYSI	5	
LAYE	R	MINIMUM	BEST	MAXIMUM		
RHO	1 2 3	6.720 130.497 3.410	10.133 215.256 3.949	3 11.919 5 464.482 9 4.752		
THICK	1 2	20.451 196.383	-0.492	2 39.988 212.832		
DEPTH	1 2	20.451 229.332	32.197 238.025	7 39.988 5 244.503		
CURI FREQUI	RENT: ENCY:	22.00 A 7.50 H	MPS EM-3 z GAIN:	37 COIL AR 6 RAMP TI	EA: 100.00 ME: 212.00 mu	sq m. SEC
No.	Т (:	IME ms)	emi DATA	f (nV/m sqrd) SYNTHE	DIFFE FIC (per	RENCE cent)
1 2 3	0 0 0	.427 .550 .698	40415.0 18264.1 8812.1	31406.3 16902.9 8835.6	22. 7. -0.	29 MASKED 45 265
ST. J Mana	OHNS GEMEN	RIVER WATE	2	SUBSURFACE	TDEM SOUNDI SOUN DAYTONA BEAC	NG DATA TABLE DING 1 CH SPEEDWAY
PALAT	KA, FI	LORIDA		DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: TABLE	93742 5.2-1

No.	TIME	em	f (nV/m sqrd)	DIFFI	ERENCE
	(ms)	DATA	SYNTHE	TIC (pe:	ccent)
4 5 6 7 8 9 10 11 12 13 14 15 16 17 18 19	$\begin{array}{c} 0.869 \\ 1.10 \\ 1.40 \\ 1.75 \\ 2.22 \\ 2.79 \\ 3.42 \\ 4.26 \\ 5.49 \\ 6.96 \\ 8.66 \\ 11.06 \\ 14.00 \\ 17.47 \\ 22.23 \\ 28.10 \end{array}$	4506.5 2324.2 1164.3 645.7 363.6 221.4 152.0 110.8 77.37 50.38 36.48 23.89 14.92 10.47 7.07 4.17	4729.3 2298.8 1143.9 635.5 349.8 227.5 153.8 110.0 72.9 50.4 35.2 23.2 15.0 9.9 6.0 3.7	$ \begin{array}{cccc} -4 \\ 1 \\ 1 \\ 3 \\ -2 \\ -1 \\ 0 \\ 5 \\ 5 \\ 0 \\ -0 \\ 1 \\ 3 \\ 2 \\ 2 \\ 1 \\ -0 \\ 7 \\ 4 \\ 10 \\ \end{array} $.94 .09 .75 .58 .77 .76 .16 .706 .71 .0426 .48 .80 .618 .69 .77 MASKED .30 MASKED
CUR	RENT: 22.00	AMPS EM-	-37 COIL AR	EA: 100.00	sq m.
FREQU	ENCY: 3.00	Hz GAIN:	8 RAMP TI	ME: 212.00 m	uSEC
No.	TIME	en	nf (nV/m sqrd)	DIFF	ERENCE
	(ms)	DATA	A SYNTHE	TIC (pe	rcent)
20 21 22 23 24 25 26 27 28 29 30 31 32 33 7 F" 1 P 1 P 2 P 3 T 1 -	0.857 1.06 1.37 1.74 2.17 2.77 3.50 4.37 5.56 6.98 8.56 10.64 13.70 17.40 METER RESOLUTI INDICATES FIXE 0.95 0.03 0.05 0.02 -0.05 (-0.04 (4590.7 2492.7 1237.1 656.5 380.1 234.3 146.2 102.7 74.60 57.73 32.68 21.89 15.59 11.69	4925.9 2590.0 1225.4 648.1 371.8 233.4 150.4 108.6 73.9 3 52.5 3 38.1 9 26.6 9 17.3 9 11.4	$ \begin{array}{c} -7\\ -3\\ 0\\ 1\\ 2\\ 0\\ -2\\ -5\\ 9\\ 0\\ 3\\ 8\\ 1\\ -16\\ 8\\ -21\\ 4\\ -11\\ 7\\ 1 \end{array} $.30 .90 .941 .28 .18 .364 .91 .75 .810 .99 MASKED .59 MASKED .59 MASKED .20 .85 MASKED
т 2	0.00 0.03 (P1 P2).02 0.00 P3 T1	T 2	TDEM SOUND	
ST. JO	OHNS RIVER WAT	ER	SDII	DAYTONA BEA	NDING 1
MANAO	GEMENT DISTRICT		SUBSURFACE		ACH SPEEDWAY
PALAT	KA, FLORIDA		DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: TABLE	93742 5.2-1

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5.3 TDEM SITE 2 - CUMBERLAND ISLAND SITE

5.3.1 Location Description and Geoelectrical Section

The Cumberland Island Site is located off the southeast coast of Georgia within Cumberland Island National Seashore Park. The crew and equipment were transported to the island on a service boat operated by the National Park Service. The sounding location (Figure 5.3-1) was in an open area in the central part of the island which also serves as an airstrip. The area was clear with the only apparent potential source of cultural noise being a power line laying on the ground within 200 ft of the western leg of the Tx loop. QC soundings were performed 200 ft east and west of the initial location of Rx coil. Results from the QC soundings indicated that the power line had no affect upon the survey values (Figure 5.3-4) except during very-late times. The apparent resistivity values from these very-late times were not used in the development of the geoelectric model.

The Floridan aquifer begins at a depth of approximately 425 ft below msl and is overlain by the Hawthorn Group (Miller, 1986). The Floridan aquifer in this area is approximately 2,400 ft thick and the Fernandina permeable zone occurs at approximately -2,050 ft msl (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.3-2. The interpreted geoelectrical section consists of a two-layer subsurface.

5.3.2 Geological Interpretation of Geoelectrical Model

The two-layer geoelectrical section, Figure 5.3-2, is the only section observed during this study. The deep low resistivity layer (saltwater interface) was not detected by TDEM measurement. The implication is that, in this area, the depth to the saltwater interface is beyond the depth capacity of the measurement system as defined by loop size and current amperage. There is a relatively thin, upper layer (80 m or 262 ft) of 16 ohm-m resistivity which can be interpreted as the Hawthorn Group and surficial sediments. The depth is smaller than values mapped in this region (Miller, 1986), but fixing the depth at 450 ft produced a poor fit to the observations. As this is a fairly unambiguous data set, the inverted value of 80 m was used instead of fixing the depth.

5.3.3 Depth to Occurrence of Salt Water

The saltwater interface is not apparent in this data set. Similar findings in this general area were cited in the CEES (1992) TDEM survey where depths to salt water were cited as being below the base of the Floridan Aquifer System. To model the sensitivity to the existence of a possible saltwater interface, a forward modeling/sensitivity analysis was performed. To do this, the TDEM data which would have been observed if a third layer of low, 3 ohm-m, resistivity had been present was modeled. To do this, the resistivity and thickness values of the upper two layers were fixed as per the original inversion, a 3 ohm-m base layer was added, and then the thickness of layer 2 (depth to layer 3) was varied. The behavior of fit error was viewed (compared to the real data) as a function of layer 2 thickness (layer 3 depth). This sensitivity analysis is expressed on Figure 5.3-5. What is seen is that if the saltwater interface is greater than 2,200 ft deep, it cannot effectively be seen. If the layer was 2,000 ft or shallower, there would have been significant indications of its existence such that a reasonable (less than 10%) fit to the data would not have been possible to produce with a two-layer model. Accordingly, it appears that the depth to the saltwater interface exceeds 2,200 ft below land surface and, quite possibly, is much deeper in this area.

5.3.4 Depth of Occurrence of the 250 mg/L Isochlor

The inversion value of conductivity for layer 2 is 58 ohm-m. It is not possible to determine the chloride concentration in layer 2 because layer 2 contains part of the Hawthorn Group. Accordingly equation (4) may not be valid.

5.3.5 Accuracy of Measurement and Interpretation

Figure 5.3-3 is the equivalence analysis at this site, and the inversion table (Table 5.3-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model. A depth to a low resistivity layer could not be determined. The equivalence range of the resistivity of layer 2 is from 53 - 64 ohm-m. A corresponding chloride concentration cannot be determined because layer 2 is in part comprised of part of the Hawthorn Group. Accordingly, equation (4) may not be valid. Results from a study of the island (McLemore, et al., 1981) indicate that the chloride to sulfate ratio is 1:5 (Table 5.1-4), which varies significantly from the 5:1 ratio used in equation (4). Accordingly, the assumptions implicit in equation (4) are not valid. The same study indicates that the chloride concentration in the Upper Floridan aquifer is 37 mg/L.

5.3.6 <u>Summary of TDEM Sounding at Cumberland Island (Site 2)</u>

- The depth to occurrence of salt water (5,000 mg/L isochlor) is unknown but it seems probable that it must be deeper than 2,200 ft from the modeling/sensitivity analysis.
- The quality of ground water within the Floridan aquifer at this site cannot be interpreted because analysis of the TDEM data does not allow the Hawthorn Group to be distinguished from the Floridan Aquifer System.











CLIENT: SJWR LOCATION: CUMB COUNTY: CAMD PROJECT: SALT LOOP SIZE: 5 COIL LOC: SOUNDING COORDI	MD ERLAND ISLAND EN COUNTY, GEORGI WATER INTERFACE 33.000 m by 0.000 m (X), NATES: E:	SO A ELE DETECTION EQU 81.000 m A 0.000 m (Y) 0.0000 N:	DATE: 27-04-93 UNDING: 1 VATION: 2.00 IPMENT: Geonics ZIMUTH: 0.0000	m PROTEM
	FITTING ERROR:	3.044 PERCE	NT	
L # RESISTIV (ohm-m	ITY THICKNESS) (meters)	ELEVATION (meters)	CONDUCTANC (Siemens)	E
1 16.07 2 57.99	80.19	2.00 -78.19	4.98	
ALL PARAMETERS	ARE FREE			
PARAMETER BOU	NDS FROM EQUIVALE	NCE ANALYSIS		
LAYER MIN	IMUM BEST	MAXIMUM		
RHO 1 1 2 5	4.41816.0793.31257.990	17.385 63.570		
THICK 1 6	4.807 1.000	95.076		
DEPTH 1 6	4.807 80.197	95.076		
CURRENT: 1 FREQUENCY: 3	9.20 AMPS EM-37 0.00 Hz GAIN: 3	COIL AREA: RAMP TIME:	100.00 sq m. 307.00 muSEC	
NO. TIME (ms)	emf DATA	(nV/m sqrd) SYNTHETIC	DIFFERENCE (percent)	
1 0.086 2 0.108 3 0.138 4 0.175 5 0.218 6 0.278 7 0.351	7 60832.5 68162.1 74568.1 79366.6 81405.9 80962.0 52363.2	65593.5 64313.1 62213.8 59155.7 55086.3 48967.6 41631.7	-7.82 MA 5.64 MA 16.56 MA 25.46 MA 32.33 MA 39.51 MA 20.49 MA	SKED SKED SKED SKED SKED SKED SKED
ST. JOHNS RIV MANAGEMENT D	ER WATER DISTRICT	SUBSURFACE	TDEM SOUNDING D SOUNDING CUMBERLAND IS	ATA TABLE 2 SLAND
PALATKA, FLOR	IDA	DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: TABLE	93742 5.3-1

No.	TIME (ms)	em DATA	f (nV/m sqrd) SYNTHE	DIFI FIC (pe	ERENCE ercent)
8 9 10 11 12 13 14 15 16 17 18 19 20	0.438 0.558 0.702 0.858 1.06 1.37 1.74 2.17 2.77 3.50 4.37 5.56 7.03	34285.0 24088.2 17238.4 12408.8 8344.5 4918.4 2773.5 1607.4 857.7 452.2 245.9 125.9 63.57	33844.8 25330.4 18087.6 12827.6 8448.2 4949.1 2827.2 1636.0 864.8 456.0 245.6 121.6 61.0	-5 -4 -3 -1 -0 -1 -1 -0 -0 -0 -0 -0 -0 -0 -0 -0 -0 -0 -0 -0	L.28 5.15 4.92 3.37 L.24 D.625 L.93 L.77 D.827 D.827 D.838 D.118 3.38 4.04
CURI FREQUI	RENT: 19.20 ENCY: 7.50	AMPS EM- Hz GAIN:	37 COIL AR 6 RAMP TI	EA: 100.00 ME: 307.00 1	0 sq m. muSEC
No.	TIME (ms)	en DATA	nf (nV/m sqrd) SYNTHE	DIF: TIC (P	FERENCE ercent)
21 22 23 24 25 26 27 28 29 30 31 32 33 34 35 36 37 38 39 40	0.346 0.427 0.550 0.698 0.869 1.10 1.75 2.22 2.79 3.42 4.26 5.49 6.96 8.66 11.06 14.00 17.47 22.23 28.10	57629.5 37691.9 25410.8 18102.2 12558.9 7158.2 5007.1 2895.0 1596.7 860.4 504.3 280.0 140.5 70.77 37.18 18.76 8.64 4.32 2.74 1.42	42160.5 34799.3 25861.9 18293.1 12570.1 7843.4 4722.8 2817.3 1569.9 870.5 505.4 280.0 138.4 71.9 38.9 19.6 10.0 2.5.3 4.2.6 2.1.3	2 - - - - - - - - - - - - - - - - - - -	6.84 MASKED 7.67 1.77 1.05 0.0888 9.57 MASKED 5.67 2.68 1.67 1.17 0.221 0.0115 1.48 1.60 4.72 4.71 6.24 MASKED 4.20 MASKED 2.03 5.17
ST. JOI	INS RIVER WAT	ER	SDII	TDEM SOUND SOUN CUMBERIA	NG DATA TABLE DING 2 ND ISLAND
MANAGE PALATK	MENT DISTRICT A, FLORIDA		SUBSURFACE DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: TABLE	93742 5.3-1

CURRENT: 19.20 AMPS EM-37 COIL AREA: 100.00 sq m. 3.00 Hz FREQUENCY: GAIN: 7 RAMP TIME: 307.00 muSEC No. TIME emf (nV/m sqrd) DIFFERENCE (ms) DATA SYNTHETIC (percent) 41 0.857 12760.9 12893.7 -1.0442 1.06 8658.2 8484.6 2.00 43 1.37 5081.1 4982.5 1.93 44 2857.4 1.74 2920.3 2.15 45 2.17 1662.2 1663.0 -0.048246 2.77 918.2 888.0 3.28 47 3.50 473.8 475.7 -0.393 48 4.37 257.5 261.8 -1.66 49 5.56 132.8 134.3 -1.07 50 6.98 80.22 71.93 10.32 MASKED 51 8.56 35.83 40.79 -13.83 MASKED 52 10.64 21.14 22.29 -5.44 PARAMETER RESOLUTION MATRIX: "F" INDICATES FIXED PARAMETER 0.98 P 1 P 2 -0.01 0.97 Т 1 -0.04 -0.04 0.92 P1 P2 T1 TDEM SOUNDING DATA TABLE SOUNDING 2 ST. JOHNS RIVER WATER CUMBERLAND ISLAND MANAGEMENT DISTRICT SUBSURFACE DETECTION PALATKA, FLORIDA PROJECT NO .: 93742 INVESTIGATIONS TABLE 5.3-1 INCORPORATED

5.4 TDEM SITE 3 - NASSAU COUNTY SITE

5.4.1 Location Description and Geoelectrical Section

The site is located in northern Nassau County, Florida, in a wooded area that had been clear-cut of timber (Figure 5.4-1). The ground surface was rather rough but accessible; no obvious noise sources could be observed near the site. A Tx loop of 1,500 ft was used. This site is 10-15 miles east, northeast from previously performed TDEM soundings (sites 3 and 4, CEES, 1992) and approximately 14 miles southwest of a USGS well N-32 (discussed in Brown, 1980).

The Floridan aquifer occurs at an approximate depth of 430 ft below msl and is overlain by the Hawthorn Group and surficial sediments (Scott et al., 1991). The top to the lower Floridan aquifer occurs at approximately 1,200 ft below msl (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.4-2. The interpreted geoelectrical section consists of a three-layer subsurface.

5.4.2 Geological Interpretation of Geoelectrical Model

The geoelectrical model for this area is similar to those from previous TDEM investigations and in general agreement with the results from USGS well N-32. A top layer of low resistivity (23 ohm-m) was present. A thickness of 433 ft (132 m) was fixed. The top layer is interpreted as the Hawthorn Group and surficial sediments and the 433 ft depth is interpreted as the top of the Floridan aquifer. The USGS well places this interface at approximately 500 ft. The increased resistivity of the second layer (6,010 ohm-m) is consistent with freshwater saturation (>80 ohm-m) within the Floridan aquifer, and the depth to our interpreted low-resistivity (saltwater) layer at 2,427 ft is deeper than but comparable to indications of salt water (7,800 mg/L at a depth of 2,094 ft) in the USGS well. According to Miller (1986) the base of the Floridan aquifer. The TDEM soundings from previous studies (CEES, 1992) determined that the interface was below the bottom of the Floridan aquifer at sites 10-15 miles west, southwest of this site. The base of the Floridan aquifer occurred at approximately 2,350 ft below msl for site 3 and 2,100 ft below msl for site 4 (CEES, 1992).

5.4.3 Depth to Occurrence of Salt Water

The bottom (third) layer of the geoelectrical model, with a resistivity of 4.7 ohm-m, is interpreted to represent salt water. It occurs at a depth of 2,427 ft (-2,424 ft msl). Because the resistivity of Layer 2 (6,010 ohm-m) is interpreted to represent fresh water within the Floridian aquifer (i.e., is greater that 80 ohm-m), the interpreted depth of the 5,000 mg/L isochlor is taken as 50 ft greater than the depth of the geoelectric interface, or at a depth of 2,477 ft (-2,474 ft msl). The resistivity of Layer 3 (4.7 ohm-m) corresponds to a chloride concentration of 6,690 mg/L assuming a porosity of 25% and the validity and applicability of Equation (4) in Section 4.2.

5.4.4 Depth of Occurrence of the 250 mg/L Isochlor

The resistivity of layer 2 (6,010 ohm-m) corresponds to a chloride content of less than 50 mg/L, assuming a 25% porosity and the validity and applicability of equation (4) in section 4.2. Using the criteria established in Section 4.3 (Class A type resistivity distribution), the position of the 250 mg/L isochlor is placed 50 ft above the depth to the low resistivity interface, or at a depth of 2,377 ft (-2,374 ft msl).

5.4.5 Accuracy of Measurement and Interpretation

Figure 5.4-3 is the equivalence analysis at this site, and the inversion table (Table 5.4-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about \pm 57 m (187 ft) which is 8% of the total depth. The resistivity of this layer has a range of from 2.8 ohm-m to 7.7 ohm-m. This corresponds to a range in interpreted chloride content from 11,334 mg/L to over 4,024 mg/l, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence range of the resistivity of layer 2 is from 943 to over 60,000 ohm-m which corresponds to a chloride concentration of less than 50 mg/L. The chloride-to-sulfate ratio at the site is 1:1 (Table 5.1-4) rather than 5:1. Accordingly, equation (4) may not be valid. A chloride concentration of less than 250 mg/L was mapped in the area of the site by Sprinkle (1981).

5.4.6 <u>Summary of TDEM Sounding at Nassau County (Site 3)</u>

- The depth to occurrence of salt water (5,000 mg/L isochlor) is interpreted to be near the base of the Floridan aquifer at a depth of 2,477 ft (-2,474 ft msl).
- The ground water within the Floridan aquifer at this site is interpreted to contain an average chloride concentration of less than 50 mg/L owing to the very high and poorly constrained resistivity valve for layer 2. The 250 mg/L isochlor is interpreted to occur above the bottom of the base of the Floridan aquifer at 2,377 ft (-2,374 ft msl).
- The results of the TDEM survey agree with water quality results from other studies 10 to 15 miles from the site.







DATA	SET: SITE3CC	
CLIENT: SJWRMD LOCATION: NASSAU COUNTY, COUNTY: NASSAU COUNTY, PROJECT: SALTWATER INTE LOOP SIZE: 457.000 m b COIL LOC: 0.000 m (SOUNDING COORDINATES: E:	D. SOUND FLORIDA ELEVAT RFACE DETECTION EQUIPM (), 0.000 m (Y) 0.0000 N:	ATE: 04-28-93 ING: 1 ION: 1.00 m ENT: Geonics PROTEM UTH: 0.0000
FITTING ER	ROR: 4.898 PERCENT	
L # RESISTIVITY THIC (ohm-m) (me	CKNESS ELEVATION ters) (meters)	CONDUCTANCE (Siemens)
1 22.96 13 2 6009.8 60 3 4.69	1.00 2.0 * -131.0 7.8 -738.8	5.74 0.101
"*" INDICATES FIXED PARA	IETER	
PARAMETER BOUNDS FROM E	QUIVALENCE ANALYSIS	
LAYER MINIMUM	BEST MAXIMUM	
RHO 1 22.597 2 2 942.927 60 3 2.839	22.964 23.379)9.815 60098.148 4.691 7.749	
THICK 1 132.000 2 559.064	0.000 132.000 1.000 672.902	
DEPTH 1 132.000 1 2 691.064 7	32.000132.00089.884804.902	
CURRENT: 19.50 AMPS FREQUENCY: 7.50 Hz (EM-57 COIL AREA: SAIN: 5 RAMP TIME: 3	100.00 sq m. 17.00 muSEC
No. TIME (ms)	emf (nV/m sqrd) DATA SYNTHETIC	DIFFERENCE (percent)
1 0.698 2289 2 0.869 1341 3 1.10 749	0.6 15072.7 3.7 10136.6 7.1 6118.8	34.15 MASKED 24.43 MASKED 18.38 MASKED
ST. JOHNS RIVER WATER	SUBSURFACE T	DEM SOUNDING DATA TABLE SOUNDING 3 NASSAU COUNTY
PALATKA, FLORIDA	DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: 93742 TABLE 5.4-1

No.	TIME	emf	(nV/m sqrd)	DIFFERENCE
	(ms)	DATA	SYNTHETI	C (percent)
4	1.40	3847.7	3537.6	$ \begin{array}{r} 8.05\\ 1.00\\ -2.75\\ -7.18\\ -5.69\\ -2.42\\ 0.703\\ 2.93\\ 5.01\\ 4.68\\ -5.90\end{array} $
5	1.75	2040.8	2020.3	
6	2.22	1015.6	1043.6	
7	2.79	495.8	531.4	
8	3.42	266.5	281.6	
9	4.26	132.7	135.9	
10	5.49	58.78	58.36	
11	6.96	26.59	25.81	
12	8.66	13.79	13.10	
13	11.06	6.74	6.43	
14	14.00	3.52	3.73	
CURR	ENT: 19.5() AMPS EM-5	7 COIL AREA	1: 100.00 sq m.
FREQUE	NCY: 3.0() Hz GAIN:	7 RAMP TIME	1: 317.00 muSEC
No.	TIME	emf	(nV/m sqrd)	DIFFERENCE
	(ms)	DATA	SYNTHETI	C (percent)
15 16 17 18 19 20 21 22 23 24 25 26 PARAME "F" IN P 1 1 P 2 0 F 1 0 T 2 0	0.857 1.06 1.37 1.74 2.17 2.77 3.50 4.37 5.56 6.98 8.56 10.64 TER RESOLUT DICATES FIXE 00 00 0.00 01 -0.01 0.00 0.00 P 1 P 2	13924.5 8140.8 4149.1 2098.3 1083.6 545.7 253.5 123.3 60.16 28.49 14.09 6.59 (ON MATRIX: ED PARAMETER 0.19 0.00 0.00 0.12 0.00 (F 3 F 1	10413.3 6660.4 3748.2 2052.2 1115.3 544.2 262.1 125.7 56.77 26.36 14.28 7.88	25.21 18.18 9.66 2.19 -2.92 0.27 -3.41 -1.97 5.63 MASKED 7.45 MASKED -1.32 -19.49 MASKED
ST. JOHI MANAGEN PALATKA,	NS RIVER WAT MENT DISTRICT , FLORIDA	ER	SUBSURFACE DETECTION INVESTIGATIONS INCORPORATED	TDEM SOUNDING DATA TABLE SOUNDING 3 NASSAU COUNTY PROJECT NO.: 93742 TABLE 5.4-1

5.5 TDEM SITE 4 - ST. AUGUSTINE #1

5.5.1 Location Description and Geoelectrical Section

The site is located in eastern St. Johns County near St. Augustine, Florida (Figure 5.5-1) and 1.5 miles from Site 5 (Figure 1-1). The site is located within an inactive agricultural field. A possible interference source (a chain link fence) existed 100 ft north of the Tx loop. An underground pipeline is also reported to be present west of the Tx loop. QA soundings were performed 100 ft north, south, east and west of the initial Rx coil location. Results from the QA soundings indicate that the apparent resistivity values were unaffected (Figure 5.5-4) by any interference sources.

The Floridan aquifer occurs at an approximate depth of 200 ft below msl and is overlain by the surficial aquifer system and the Hawthorn Group (Scott et al., 1991). The top of the Lower Floridan aquifer occurs at approximately 900 ft below msl (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.5-2. The interpreted geoelectrical section consists of a three-layer subsurface.

5.5.2 Geological Interpretation of Geoelectrical Model

The three-layered geoelectrical section consists of a low resistivity (15 ohm-m), upper layer which is considered to be the Hawthorn Group and surficial sediments above the Floridan aquifer. The thickness of layer 1 was fixed at a 73 m (239 ft) value based on published information (Scott et al., 1991). The second layer has only intermediate resistivity (44 ohm-m) which, because it is less than 80 ohm-m, suggests the Floridan aquifer at this site contains brackish water. The thickness of the brackish section is 108 m (354 ft), placing the depth to the low resistivity (saltwater) layer at 181 m (594 ft) below ground surface. The resistivity of the saltwater saturated layer is 10.3 ohm-m. Layer 1 is considered to be the Hawthorn Group and surficial sediments, layer 2 to be the Floridan aquifer (brackish) and layer 3 to be the salt water within the Floridan aquifer.

5.5.3 Depth to Occurrence of Salt Water

The bottom (third) layer of the geoelectrical model, with a resistivity of 10.3 ohm-m, is interpreted to represent salt water. It occurs at a depth of 594 ft (-555 ft msl). Because the resistivity of layer 2 (44 ohm-m) is interpreted to represent brackish water within the Floridan aquifer (i.e., is less than 80 ohm-m), the interpreted depth to the 5,000 mg/L isochlor is equal to the depth of the geoelectrical interface, or at 594 ft depth (-555 ft msl). The resistivity of layer 3 (10.3 ohm-m) corresponds to a chloride content of 2,970 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2.

5.5.4 Depth of Occurrence of the 250 mg/L Isochlor

The resistivity of layer 2, 44 ohm-m, corresponds to a chloride content above 250 mg/L, assuming a 25% porosity and the validity and applicability of equation (4) of Section 4.2. As the interpreted chloride content exceeds 250 mg/L, the 250 mg/L isochlor does not occur within the Floridan aquifer at this site.

5.5.5 Accuracy of Measurement and Interpretation

Figure 5.5-3 is the equivalence analysis at this site and the inversion table (Table 5.5-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about \pm 18 m (59 ft) which is 10% of the total depth. The resistivity of this layer has a range from 9.4 -11.3 ohm-m. This corresponds to a range in interpreted chloride content of from 3,269 mg/L to 2,693 mg/L, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence range of the resistivity of layer 2 is from 34 - 64 ohm-m which corresponds to chloride content above 250 mg/L. The chloride-to-sulfate ratio at the site is 1:2 (Table 5.1-4), rather than 5:1. Accordingly, equation (4) may not be valid. A chloride concentration above 250 mg/L was mapped in the Upper Floridan aquifer at this site by Spechler and Hampson (1984).
5.5.6 <u>Summary of TDEM Sounding at St. Augustine #1 (Site 4)</u>

- The depth to occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 594 ft (-555 ft msl) and occur within the Upper Floridan aquifer.
- The ground water within the Floridan aquifer at this site is interpreted to contain an average chloride concentration above 250 mg/L. The 250 mg/L isochlor is not interpreted to be present within the Floridan aquifer.
- Results of the TDEM survey agree with a water quality study in the area of the site which indicates that chloride concentrations in the Floridan aquifer are above 250 mg/L.









		DATA SET: S	SITE4C	
CLIE LOCATI COUN PROJE LOOP SI COIL L SOUNDIN	NT: SJWRMD ON: ST. AUGUS TY: ST. JOHNS CT: SALT-WATE ZE: 305.000 OC: 0.000 IG COORDINATES	TINE SITE 1 COUNTY, FLO R INTERFACE D m by C D m (X), : E:	SC DRIDA ELE DETECTION EQU 305.000 m / 0.000 m (Y) 0.0000 N:	DATE: 29004-93 DUNDING: 2 EVATION: 12.00 m JIPMENT: Geonics PROTEM AZIMUTH: 0.0000
	FITTI	NG ERROR:	2.296 PERCH	ENT
L #	RESISTIVITY (ohm-m)	THICKNESS (meters)	ELEVATION (meters)	N CONDUCTANCE (Siemens)
1 2 3	14.64 44.03 10.31	73.00 108.0	12.00 * -61.00 -169.0	4.98 2.45
"*" IN	DICATES FIXED	PARAMETER		
PARAM	IETER BOUNDS FI	ROM EQUIVALI	ENCE ANALYSIS	
LAYEF	R MINIMUM	BEST	MAXIMUM	
RHO	1 14.118 2 33.753 3 9.449	14.649 44.037 10.313	15.174 63.719 11.287	
THICK	1 73.000 2 91.238	0.000 1.000	73.000 126.498	
DEPTH	1 73.000 2 164.238	73.000 181.011	73.000 199.498	
CURI FREQUI	RENT: 14.00 ENCY: 7.50	AMPS EM-5 Hz GAIN:	7 COIL AREA 6 RAMP TIME	: 100.00 sq m. : 172.00 muSEC
No.	TIME (ms)	emf DATA	(nV/m sqrd) SYNTHETI	DIFFERENCE C (percent)
1 2 3	0.346 0.427 0.550	42574.3 29824.4 18484.0	39953.5 29130.5 18684.3	6.15 2.32 -1.08
ST. JO	HNS RIVER WATE	ER	SUBSUBSIA	TDEM SOUNDING DATA TABLE SOUNDING 4 ST. AUGUSTINE #1
MANAGI PALATK	A, FLORIDA		SUBSURFACE DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: 93742 TABLE 5.5-1

No	ጥፐለፑ	omf	(nV/m_sard)	DIFFERENCE
10.	(ms)	DATA	SYNTHETIC	C (percent)
4	0.698	11420.3	11640.9	-1.93
5	0.869	6984.7	7222.5	-3.40
6	1.10	4184.7	4114.6	1.67
7	1.40	2350.8	2347.5	0.141
8	1.75	1386.6	1375.4	0.809
9	2.22	793.4	784.1	1.16
10	2.79	460.5	462.2	-0.364
11	3.42	293.3	291.1	0.726
12	4.26	182.0	180.2	0.959
13	5.49	103.3	102.8	0.518
14	6.96	60.20	61.63	-2.36
15	8.66	36.73	37.91	-3.19 MASKED
10	14 00	21.30	21.99 19 77	-2.00 MADABU _10 40 Magypd
10 10	14.UU 17 <i>1</i> 7	5 20 TT·30	14.11	-10.40 MACKED -20.91 Macked
10	1/04/ 20.03	2 22	4.23	-31.30 MASKED
20	28,10	1.79	2.35	-31.01 MASKED
20	~~	1 • <i>1 J</i>	ل ل • ي	
CUR	RENT: 14.00	AMPS EM-5	7 COIL AREA	: 100.00 sq m.
FREQU	ENCY: 3.00	Hz GAIN:	8 RAMP TIME	: 172.00 muSEC
				DT 3755 51/27
No.	TIME	emf	(nv/m sqrd)	
	(ms)	DATA	SYNTHETI	c (percent)
9 1	0 257	7157 ว	7454 1	_4.15
∠⊥ 22	1.06	4460.3	4507.1	-1.05
23	1.37	2481.9	2486.9	-0.202
24	1.74	1414.5	1396.3	1.28
25	2.17	830.7	828.1	0.306
26	2.77 490.5		471.5	3.87
27	3.50	279.9	278.1	0.635
28	4.37	169.6	171.9	-1.34
29	5.56	100.9	101.4	-0.440
30	6.98	61.66	62.47	-1.31
31	8.56	38.83	40.02	-3.06
32	10.64	24.21	24.97	-3.13
33	13.70	14.39	14.28	0.762
34	17.40	8.54	8.39	1./9 2.61
35	21.70	5.24	5.05	J.01 50 05 MACVED
30	21.10	1.19	2.0/	-32.23 HADVED
		DADAMPOPD	RESOLUTION MAD	'RTX:
		- аканьтык "F" типтся	TES FIXED DADA	METER
		P 1 0.99	LAUN FARM	
		P 2 0.05	0.51	
		P 3 0.00	-0.04 0.98	
		F1 0.00	0.00 0.00 0	.00
		T 2 -0.01	0.19 0.04 0	0.00 0.89
		P 1	P2 P3	FI TZ
L				
		T T	CUIL	TDEM SOUNDING DATA TABLE
ST. JO	HNS RIVER WA	TER I	SDI	SOUNDING 4
MANAG	EMENT DISTRICT	r I	SUBSURFACE	SI. AUGUSTINE #1
	(A. FLORIDA		DETECTION	PROJECT NO.: 93742
	- y - LONION	l	INVESTIGATIONS	
			INCORPORATED	1-100LE 3,3-1

5.6 TDEM SITE 5 - ST. AUGUSTINE #2

5.6.1 Location Description and Geoelectrical Section

The site is in eastern St. Johns County near St. Augustine, Florida (Figure 5.6-1) and 1.5 miles from Site 4 (Figure 1-1). The site is located within a tree farm. A 16-inch water line runs parallel to the west side of the Tx loop approximately 75 ft away. QA soundings were performed 50 ft north, south, east, and west of the initial Rx coil position. Results from the QA soundings indicate that the apparent resistivity values were unaffected by any interference sources (Figure 5.6-4).

The Floridan aquifer occurs at an approximate depth of 200 ft below msl and is overlain by the surficial aquifer system and the Hawthorn Group (Scott et al., 1991). The top of the Lower Floridan aquifer occurs at approximately 900 ft below msl (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.6-2. The interpreted geoelectrical section consists of a three-layer subsurface.

5.6.2 Geological Interpretation of Geoelectrical Model

There is a sufficient electrical resistivity contrast to distinguish two geological layers above the third saltwater saturated layer. The first layer occurs at a depth of 45 m (148 ft) and not at the hydrostratigraphic contact (239 ft bls) between the Hawthorn Group and the Floridan Aquifer System. The first layer has a low-resistivity value (9.5 ohm-m) and is considered to represent the upper portion of the Hawthorn Group and surficial sediments. The second layer has a highresistivity value (169 ohm-m) and a thickness of 129 m (423 ft). It is considered to represent a combined but indistinguishable (geoelectricaly) layer consisting of a portion of the Hawthorn Group and the upper portion of the Upper Floridan aquifer. The third layer is considered to represent a saltwater saturated Floridan aquifer at a depth of 571 ft.

5.6.3 Depth to Occurrence of Salt Water

The bottom (third) layer of the geoelectrical model, with a resistivity of 18.4 ohm-m, is interpreted to represent salt water. It occurs at a depth of 571 ft (-532 ft msl). Because the resistivity of layer 2 (169 ohm-m) is greater than 80 ohm-m, the interpreted depth to the 5,000 mg/L isochlor is taken as 50 ft greater than the depth of the geoelectrical interface, or at a depth of 621 ft (-582 ft msl). The resistivity of layer 3 (18.4 ohm-m) corresponds to a chloride content of 1,595 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2.

5.6.4 Depth of Occurrence of the 250 mg/L Isochlor

Because of the inability to segregate the Floridan aquifer from the overlying surficial aquifer system and the Hawthorn Group, the effective chloride concentration of Layer 2 cannot be calculated.

5.6.5 Accuracy of Measurement and Interpretation

Figure 5.6-3 is the equivalence analysis at this site and the inversion table (Table 5.6-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about $\pm 19 \text{ m}$ (62 ft) which is 11% of the total depth. The resistivity of this layer has a range of from 17.2 - 19.8 ohm-m. This corresponds to a range in interpreted chloride content of from 1,471 mg/L to 1,717 mg/L, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence range of the resistivity of layer 2 is from 118 - 312 ohm-m. A corresponding chloride concentration cannot be determined because Layer 2 is in part comprised of the Hawthorn Group and surficial sediments. A chloride concentration of approximately 250 mg/L was mapped by Spechler and Hampson (1984) in the area of the site.

5.6.6 <u>Summary of TDEM Sounding at St. Augustine #2 (Site 5)</u>

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 621 ft (-582 ft msl) and occur within the Upper Floridan aquifer.
- The ground water within the Floridan aquifer at this site cannot be interpreted because analysis of the TDEM data does not allow the surficial aquifer system and the Hawthorn Group to be distinguished from the Floridian Aquifer System.









DATA SET: SITE5CC								
CLI LOCAT COU PROJ LOOP S COIL SOUNDI	ENT: S ION: S NTY: S ECT: S IZE: LOC: NG COO	JWRMD T. AUGUST T. JOHNS ALT WATER 145.000 0.000 RDINATES:	INE-SITE 2 COUNTY INTERFACE 1 m by 60 m (X), E:	SC ELH DETECTION EQU 09.600 m A 0.000 m (Y) 0.0000 N:	DATE: 30-04-93 DUNDING: 1 EVATION: 12.00 m JIPMENT: Geonics PROTEM AZIMUTH: 0.0000			
		FITTIN	G ERROR:	1.951 PERCE	ENT			
L # RESISTIVITY THICKNESS ELEVATION CONDUCTANCE (ohm-m) (meters) (meters) (Siemens)								
1 2 3	9 169 18	.46 .0 .40	44.62 129.3	12.00 -32.62 -161.9	4.71 0.765			
ALL P	ARAMET	ERS ARE F	REE					
PARA	METER	BOUNDS FR	OM EQUIVALE	NCE ANALYSIS				
LAYE	R	MINIMUM	BEST	MAXIMUM				
RHO	1 2 3	7.953 117.906 17.158	9.463 169.073 18.406	10.331 312.475 19.821				
THICK	1 2	34.273 117.992	0.540 1.000	51.209 142.644				
DEPTH	1 2	34.273 152.728	44.630 174.000	51.209 189.693				
CUR FREQU	RENT: ENCY:	18.00 A 7.50 H	MPS EM-57 z GAIN: 6	COIL AREA RAMP TIME	: 100.00 sq m. : 286.00 muSEC			
No.	TI (#	ME IS)	emf DATA	(nV/m sqrd) SYNTHETI(DIFFERENCE C (percent)			
1 2 3	0. 0. 0.	346 427 550	62284.0 36900.1 21169.2	52025.2 35263.4 21175.3	16.47 MASKED 4.43 -0.0290			
ST. MAN	ST. JOHNS RIVER WATER ANA SOUNDING DATA TABLE SOUNDING 5 MANAGEMENT DISTRICT SUBSURFACE ST. AUGUSTINE #2							
PAL	PALATKA, FLORIDA DETECTION PROJECT NO.: 93742 INVESTIGATIONS							

No.	TIME	emf	(nV/m sqrd)	DIFFERENCE				
	(ms)	DATA	SYNTHETIC	(percent)				
4 5 6 7 8 9 10 11 12 13 14 15 16 17 18 19 20	0.698 0.869 1.10 1.40 1.75 2.22 2.79 3.42 4.26 5.49 6.96 8.66 11.06 14.00 17.47 22.23 28.10	12503.77442.84418.12335.11319.0718.0396.2243.8145.381.7946.5627.6616.008.775.042.691.27	12662.67648.24236.32350.21330.4727.3409.5248.6146.980.6146.2827.6515.518.775.102.761.50	-1.27 -2.75 4.11 -0.647 -0.864 -1.29 -3.36 -1.99 -1.11 1.45 0.589 0.0325 3.04 -0.0455 -1.21 -2.51 -18.33 MASKED				
CU:	RRENT: 18.00	AMPS EM-57	COIL AREA:	100.00 sq m.				
FREQ	UENCY: 3.00	Hz GAIN: 7	RAMP TIME:	286.00 muSEC				
No.	TIME	emf	(nV/m sqrd)	DIFFERENCE				
	(ms)	DATA	SYNTHETIC	(percent)				
21	0.857	7825.9	7906.3	-1.02				
22	1.06	4687.5	4660.2	0.580				
23	1.37	2521.6	2497.2	0.969				
24	1.74	1364.5	1351.3	0.967				
25	2.17	761.7	771.0	-1.22				
26	2.77	430.2	418.0	2.83				
27	3.50	234.9	236.1	-0.502				
28	4.37	137.9	139.1	-0.856				
29	5.56	79.84	79.09	0.949				
30	6.98	44.04	46.73	-6.09 MASKED				
31	8.56	26.69	29.10	-9.05 MASKED				
PARA	PARAMETER RESOLUTION MATRIX:							
"F"	"F" INDICATES FIXED PARAMETER							
P 1	P 1 0.93							
P 2	P 2 -0.01 0.02							
P 3	P 3 0.03 0.00 0.93							
T 1 T 2	-0.10 -0.07 0. -0.03 0.10 0. P1 P2	.05 0.83 .08 -0.06 0. P 3 T 1	88 T 2					
ST. MAN PAL/	JOHNS RIVER WAT AGEMENT DISTRICT ATKA, FLORIDA	TER	SUBSURFACE DETECTION INVESTIGATIONS INCORPORATED	TDEM SOUNDING DATA TABLE SOUNDING 5 ST. AUGUSTINE #2 PROJECT NO.: 93742 TABLE 5.6-1				

5.7 TDEM SITE 6 - PICOLATA SITE

5.7.1 Location Description and Geoelectrical Section

The site is located in western St. Johns County, Florida (Figure 5-7.1). The site was located within a wooded area with no obvious sources of interference.

The Floridan aquifer begins at a depth of approximately 285 ft below msl and is overlain by the Hawthorn Group and surficial aquifer system (Scott et al., 1991). The top of the Lower Floridan aquifer begins at approximately 800 ft below msl. The base of the Floridan aquifer occurs at approximately 2,100 ft below msl in this area (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.7-2. The interpreted geoelectrical section consists of a three-layer subsurface.

5.7.2 Geological Interpretation of Geoelectrical Model

The three-layer geoelectrical section consists of a low resistivity (23 ohm-m), upper layer which correlates with the Hawthorn Group and surficial sediments above the Floridan aquifer. The thickness of layer 1 was fixed at 90 m (295 ft) based on the information from Scott et al. (1991). The second layer has high resistivity (98 ohm-m) which means that because it is greater than 80 ohm-m the Floridan aquifer at this site contains fresh water. The thickness of the freshwater section is 488 m (1,601 ft) placing the depth to the low resistivity (saltwater) layer at 578 m (1,896 ft) below ground surface. The resistivity of the saltwater layer is 27.4 ohm-m. Layer 1 is considered to be the Hawthorn Group and surficial sediments, layer 2 to be the Floridan aquifer.

5.7.3 Depth to Occurrence of Salt Water

The bottom (third) layer of the geoelectrical model, with a resistivity of 27.4 ohm-m, is interpreted to represent salt water. It occurs at a depth of 1,896 ft (-1,886 ft msl). Because the resistivity of layer 2 (98 ohm-m) is interpreted to represent fresh water within the Floridan aquifer (i.e., is greater than 80 ohm-m), the interpreted depth to the 5,000 mg/L isochlor is taken as 50 ft greater than the depth of the geoelectrical interface, or at 1,946 ft depth (-1,936 ft msl). The resistivity of layer 3 (27.4 ohm-m) corresponds to a chloride content of 1,021 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2. It is presumed that because of the expected high chlorinity gradients, this value is sufficiently close to the 5,000 mg/L isochlor that they represent the same effective depth.

5.7.4 Depth of Occurrence of the 250 mg/L Isochlor

The resistivity of layer 2, 98 ohm-m, corresponds to a chloride content of less than 250 mg/L, assuming a 25% porosity and the validity and applicability of equation (4) of Section 4.2. The 250 mg/L isochlor is placed in the Lower Floridan aquifer at a depth 50 ft above the layer 3 interface or at 1,846 ft (-1,836 ft msl).

5.7.5 Accuracy of Measurement and Interpretation

Figure 5.7-3 is the equivalence analysis at this site and the inversion table (Table 5.7-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about ± 121 m (397 ft) which is 21% of the total depth. The resistivity of this layer has a range of from 14.0 - 49.3 ohm-m. This corresponds to a range in interpreted chloride content of from 2,144 mg/L to 499 mg/L, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence range of the resistivity of layer 2 is from 89 - 122 ohm-m which corresponds to a chloride content of less than 250 mg/L. The chloride-to-sulfate ratio at the site is 1:5 (Table 5.1-4), rather than 5:1. Accordingly, equation (4) may not be valid. A chloride concentration of less than 50 mg/L was determined in the area of the site by Spechler and Hampson (1984).

5.7.6 Summary of TDEM Sounding at Picolata (Site 6)

- The depth to occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 1,946 ft (-1,936 ft msl) and occur within the Lower Floridan aquifer.
- The ground water within the Floridan aquifer at this site is interpreted to contain an average chloride concentration of less than 250 mg/L. The 250 mg/L isochlor is interpreted to be present in the Lower Floridan aquifer at a depth of 1,846 ft (-1,836 ft msl).
- Results from the TDEM survey agree with the results from a water quality study in the area of the site.







DATA SET: SITE6

CLIE LOCATI COUN PROJE LOOP SI COIL I	ENT: S ION: F NTY: S SCT: S IZE: LOC:	JWRMD PICOLATA ST. JOHNS (SALT WATER 305.000 0.000	COUNTY, FL INTERFACE m by m (X), F,	ORIDA DETECTION 305.000 m 0.000 m (0.0000 N:	SOU ELEV EQUI AZ Y)	DATE: 31-04- NDING: 1 ATION: 3. PMENT: Geonic IMUTH:	93 00 m s PROTEM
SOONDII		BIMMINO.		0.0000 M.	PDCPN	m	
		LTTTN	3 ERROR:	2.40/ F	ercen		
L #	RESIS (ol	STIVITY 1m-m)	THICKNESS (meters)	ELEVA (met	TION ers)	CONDUCTA (Siemen	NCE S)
1 2 3	2: 9: 2:	3.26 7.79 7.36	90.00 488.4	3. * -87. -575.	00 00 4	3.86 4.99	
"*" I	NDICA	res fixed 1	PARAMETER				
PARA	METER	BOUNDS FRO	OM EQUIVAI	LENCE ANALYS	IS		
LAYE	R	MINIMUM	BEST	MAXIMUM	E Contraction of the second se		
RHO	1 2 3	22.388 88.691 13.969	23.263 97.799 27.369	23.814 2121.519 249.336			
THICK	1 2	90.000 341.324	0.000	90.000 582.781	•		
DEPTH	1 2	90.000 431.324	90.000 578.472	90.000 2 672.781			·
CUR FREQU	RENT: ENCY:	18.00 A 7.50 H	MPS EM- z GAIN:	57 COIL A 4 RAMP I	REA:	100.00 sq 183.00 muSEC	m. :
No.	Т (IME ms)	em: DATA	f (nV/m sqrd SYNTH	l) IETIC	DIFFEREI (percei	NCE ht)
1 2 3	0 0 1	.698 .869 .10	9362.0 5183.3 2829.5	8683. 5188. 2815.	.9 .0 .6	7.24 -0.093 0.492	LO 2
ST. JOH	NS RI	VER WATER		SUBSUBEACE	-	TDEM SOUNDING SOUNDII PICOL	: DATA TABLE NG 6 ATA
PALATKA	, FLO	RIDA		DETECTION INVESTIGATION INCORPORATE	IS D	PROJECT NO .: TABLE	93742 5.7-1

No.	TIME (ms)	em DATA	f (nV/m sqrd) SYNTHET	DIFFERE IC (perce	NCE nt)				
4 5 7 8 9 10 11 12	1.40 1.75 2.22 2.79 3.42 4.26 5.49 6.96 8.66	1462.7 799.7 415.8 214.4 120.7 63.67 30.39 14.99 8.59	1516.1 824.3 421.7 218.1 119.8 63.95 30.63 16.09 8.95	-3.64 -3.07 -1.42 -1.72 0.77 -0.43 -0.79 -7.32 -4.13	8 6 7 MASKED MASKED				
13	11.06	5.07	4.85	4.27	MASKED				
CURI FREQUI	RENT: 18.00 ENCY: 3.00	AMPS EM- Hz GAIN:	57 COIL ARE 8 RAMP TIM	A: 100.00 sq E: 183.00 muSE	[m. C				
No.	TIME (ms)	em DATA	f (nV/m sqrd) SYNTHET	DIFFERE IC (perce	INCE ent)				
14 15 16 17 18 19 20 21 22 23 24 25 26	0.857 1.06 1.37 1.74 2.17 2.77 3.50 4.37 5.56 6.98 8.56 10.64 13.70	5407.2 3077.7 1581.2 826.9 444.1 234.4 115.4 59.79 29.86 16.30 9.29 5.39 2.99	5367.7 3107.8 1616.3 838.0 450.3 223.1 112.3 59.67 29.87 16.23 9.47 5.55 3.01	$\begin{array}{c} 0.73 \\ -0.97 \\ -2.21 \\ -1.33 \\ -1.39 \\ 4.83 \\ 2.63 \\ 0.20 \\ -0.02 \\ 0.44 \\ -1.86 \\ -2.87 \\ -0.49 \end{array}$	0 7 8 9 9 8 8 8 1 7 9 0				
PARAM "F" I P 1 P 2 P 3 F 1 T 2 -	PARAMETER RESOLUTION MATRIX: "F" INDICATES FIXED PARAMETER P 1 0.99 P 2 0.02 0.92 P 3 0.01 -0.09 0.24 F 1 0.00 0.00 0.00 0.00 T 2 -0.01 0.07 0.20 0.00 0.88 P 1 P 2 P 3 F 1 T 2								
ST. JOH MANAGE	NS RIVER WAT	ER		TDEM SOUNDING SOUND PICOL	G DATA TABLE NG 6 ATA				
PALATKA	, FLORIDA		DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: TABLE	93742 5.7-1				

5.8 TDEM SITE 7 - GREEN COVE SPRINGS SITE

5.8.1 Location Description and Geoelectrical Section

The site is in eastern Clay County, Florida (Figure 5-8.1). The site is a pasture. A flowing well was located near the center of the Tx loop.

The Floridan aquifer begins at a depth of approximately 285 ft below msl and is overlain by the Hawthorn Group and surficial aquifer system (Scott et al., 1991). The top of the Lower Floridan aquifer begins at approximately 700 ft below msl. The base of the Floridan aquifer is approximately 2,000 ft below msl in this area (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.8-2. The interpreted geoelectrical section consists of a three-layer subsurface.

5.8.2 Geological Interpretation of Geoelectrical Model

The three-layered geoelectrical section consists of a low resistivity (29 ohm-m), upper layer considered to be the Hawthorn Group and surficial sediments above the Floridan aquifer. The thickness of layer 1 was fixed at 93 m (305 ft) based on published information (Scott et al. 1991). The second layer has high resistivity (108 ohm-m) which, because it is greater than 80 ohm-m, means the Floridan aquifer at this site contains fresh water. The thickness of the freshwater section is 383 m (1,257 ft) placing the depth to the low resistivity (saltwater) layer at 476 m (1,562 ft) below ground surface. The resistivity of the saltwater layer is 2.3 ohm-m. Layer 1 is considered to be the Hawthorn Group and surficial sediments, layer 2 to be the Floridan aquifer containing fresh water and layer 3 to be the salt water within the Lower Floridan aquifer.

5.8.3 Depth to Occurrence of Salt Water

The bottom (third) layer of the geoelectrical model, with a resistivity of 2.3 ohm-m, is interpreted to represent salt water. It occurs at a depth of 1,562 ft (-1,542 ft msl). Because the resistivity of layer 2 (108 ohm-m) is interpreted to represent fresh water within the Floridan aquifer (i.e., is greater than 80 ohm-m), the interpreted depth to the 5,000 mg/L isochlor is taken as 50 ft greater than the depth of the geoelectrical interface, or at 1,612 ft depth (-1,592 ft msl). The resistivity of layer 3 (2.3 ohm-m) corresponds to a chloride content of 13,831 mg/L, assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2.

5.8.4 Depth of Occurrence of the 250 mg/L Isochlor

The resistivity of layer 2, 108 ohm-m, corresponds to a chloride content of less than 250 mg/L, assuming a 25% porosity and the validity and applicability of equation (4) of Section 4.2. The 250 mg/L isochlor is placed within the lower Floridan aquifer at a depth 50 ft above the layer 3 interface or at 1,512 ft (-1,492 ft msl).

5.8.5 Accuracy of Measurement and Interpretation

Figure 5.8-3 is the equivalence analysis at this site and the inversion table (Table 5.8-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about $\pm 20 \text{ m}$ (66 ft) which is 4% of the total depth. The resistivity of this layer has a range from 1.0 - 4.9 ohm-m. This corresponds to a range in interpreted chloride content of from 32,011 mg/L to 6,411 mg/L. The chloride-to-sulfate ratio at the site is 1:1 (Table 5.1-4), rather than 5:1. Accordingly, equation (4) may not be valid. A chloride concentration of less than 250 mg/L was determined in the area of the site by Sprinkle (1981).

5.8.6 <u>Summary of TDEM Sounding at Green Cove Springs (Site 7)</u>

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 1,612 ft (-1,592 ft msl) and occur within the Lower Floridan aquifer.
- The ground water within the Floridan aquifer at this site is interpreted to contain an average chloride concentration of less than 250 mg/L. The 250 mg/L isochlor is interpreted to be present within the Lower Floridan aquifer at a depth of 1,512 ft (-1,492 ft msl).
- Results of the TDEM survey agree with the results of other water quality studies in the area of the site.



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		DATA SET:	SITE7	
CLII LOCAT COU PROJI LOOP S COIL I SOUNDII	ENT: SJWRMD ION: GREEN COV NTY: CLAY COUN ECT: SALTWATEN IZE: 305.00 LOC: 0.00 NG COORDINATES	VE SPRINGS NTY, FLORIDA R INTERFACE 1 00 m by 00 m (X), 5: E:	S EL DETECTION EQ 305.000 m 0.000 m (Y) 0.0000 N:	DATE: 01-05-93 OUNDING: 1 EVATION: 6.00 m UIPMENT: Geonics PROTEM AZIMUTH: 0.0000
	FITT]	ING ERROR:	5.711 PERC	ENT
L #	RESISTIVITY (ohm-m)	THICKNESS (meters)	ELEVATIO (meters	N CONDUCTANCE) (Siemens)
1 2 3	28.96 108.4 2.30	93.00 383.2	6.00 * -87.00 -470.2	3.21 3.53
"*" I	NDICATES FIXE	D PARAMETER		
PARA	METER BOUNDS I	ROM EQUIVAL	ENCE ANALYSIS	
LAYE	R MINIMUM	BEST	MAXIMUM	
RHO	1 26.57 2 79.104 3 1.03	L 28.969 H 108.451 L 2.305	31.279 167.219 4.877	
THICK	1 93.000 2 362.998	0.000 3 1.000	93.000 401.831	
DEPTH	1 93.000 2 455.990	93.000 8 476.255	93.000 494.831	
CUR FREQU	RENT: 16.00 ENCY: 7.50	AMPS EM-5 Hz GAIN:	7 COIL AREA 3 RAMP TIME	: 100.00 sq m. : 202.00 muSEC
No.	TIME (ms)	emf DATA	(nV/m sqrd) SYNTHETI	DIFFERENCE C (percent)
1 2 3	0.869 1.10 1.40	3682.8 1815.9 839.7	3170.3 1689.3 875.7	13.91 MASKED 6.97 -4.29
ST. JO MANAG	HNS RIVER WAT EMENT DISTRICT	ER	SUBSURFACE	TDEM SOUNDING DATA TABLE SOUNDING 7 GREEN COVE SPRINGS
PALATK	A, FLORIDA		DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: 93742 TABLE 5.2-1

No.	TIME (ms)	D.	emf (n ATA	V/m sqrd) SYNTHE	FIC	DIFFERE (perce	NCE nt)
4 5 6	1.75 2.22 2.79	433 214 104	.3 .4 .9	459.0 223.9 109.7		-5.92 -4.45 -4.55	MASKED
CURI FREQUI	RENT: 1 ENCY:	6.00 AMPS 3.00 Hz GA	EM-57 IN: 8	COIL ARI RAMP TII	EA: ME: 20	100.00 sq 2.00 muSE	m. C
No.	TIME (ms)	D.	emf (n ATA	V/m sqrd) SYNTHE	FIC	DIFFERE (perce	NCE nt)
7 8 9 10 11 12 13 14 15 16 17 18 19	0.857 1.06 1.37 1.74 2.17 2.77 3.50 4.37 5.56 6.98 8.56 10.64 13.70	3838 1993 916 452 233 118 59 32 17 10 8 55	.7 .9 .2 .8 .5 .7 .51 .98 .99 .49 .09 .99 .09	3284.0 1874.9 938.4 468.1 241.4 113.4 57.1 31.0 18.0 11.8 8.6 6.4 4.6	2 6 4 7 2 0	14.45 5.96 -2.42 -3.38 -3.38 4.48 4.01 5.82 -0.26 -12.99 -7.17 -7.12 9.76	MASKED 7 MASKED
PARAM "F" I P 1 P 2 P 3 F 1 T 2	ETER RESO NDICATES 0.99 0.05 0.7 0.03 -0.2 0.00 0.0 0.00 0.0 P 1 P	LUTION MATRI FIXED PARAME 5 4 0.32 0 0.00 0.0 0 -0.03 0.0 2 P 3 F	X: TER 0 0.99 1 T	2			
ST. JOHN MANAGEM	IS RIVER I IENT DISTR	WATER NCT	SUBS DETE	DII SURFACE	TDEN GF	A SOUNDING SOUNDIA REEN COVE	DATA TAE NG 7 SPRINGS 93742
· · · · · · · · · · · · · · · · · · ·	. 2011071		INVES INCO	STIGATIONS RPORATED	TABI	£	5.2-1

5.9 TDEM SITE 8 - UNION CAMP SITE

5.9.1 Location Description and Geoelectrical Section

The site is located in a wooded area of Putnam County, Florida (Figure 5.9-1). QA soundings were performed 70 ft to the south and 60 ft to the west of the initial Rx coil position. Results from the QA soundings indicate that the apparent resistivity values were not affected by any interference sources (Figure 5.9-4).

The Floridan aquifer begins at an approximate depth of 65 ft below msl and is overlain by the Hawthorn Group and the surficial aquifer system (Scott et al., 1991). The Lower Floridan aquifer begins at an approximate depth of 700 ft below msl (Miller, 1986). An observation well P-0735 is present approximately 1/4 mile from the site (Figure 5.9-1). Based on information provided by SJRWMD the depth to the 250 mg/L isochlor in the observation well occurs at approximately 340 - 360 ft bls and the depth to the 5,000 mg/L isochlor occurs at approximately 450 ft bls.

The resistivity sounding data and best-fit model inversion are presented on Figure 5.9-2. The interpreted geoelectrical section consists of a two-layer subsurface.

5.9.2 Geological Interpretation of Geoelectrical Model

There is insufficient electrical resistivity contrast between the surficial aquifer system layer, the Hawthorn Group and the underlying Floridan aquifer to distinguish the three. Fixing the thickness of the upper layer does not resolve this dilemma; therefore it can be interpreted that there exists a two-layer geoelectrical section with a relatively thick (251.9 m = 826 ft) surface layer of intermediate resistivity (47.2 ohm-m) overlying a low resistivity layer (3.7 ohm-m). It can be interpreted that the surficial aquifer system, the Hawthorn Group, and the upper part of the Floridan aquifer system exist as a combined but indistinguishable (geoelectrical) layer, overlying a saltwater saturated Lower Floridan aquifer at a depth of 826 ft bls.

5.9.3 Depth to Occurrence of Salt Water

The bottom (second) layer of the geoelectrical model, with a resistivity of 3.7 ohm-m, is interpreted to represent salt water. It occurs at a depth of 826 ft (-810 ft msl). Because the resistivity of layer 1 (47.2 ohm-m) is less than 80 ohm-m, the interpreted depth to the 5,000 mg/L isochlor is taken at the depth of the geoelectrical interface, or at 826 ft depth (-810 ft msl). The resistivity of layer 2 (3.7 ohm-m) corresponds to a chloride content of 8,540 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2. The depth to the 5,000 mg/L isochlor (826 ft bls) does not correlate well to water quality results from nearby observation well P-0735 which places the 5,000 mg/L isochlor at 450 ft bls (Figure 5.10-5).

5.9.4 Depth of Occurrence of the 250 mg/L Isochlor

Because of the inability to segregate the Floridan aquifer from the overlying Hawthorn Group and surficial aquifer system, the effective chloride concentration of layer 1 cannot be calculated.

5.9.5 Accuracy of Measurement and Interpretation

Figure 5.9-3 is the equivalence analysis at this site and the inversion table (Table 5.9-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model. The range of equivalence in determining the depth to the low resistivity layer is about \pm 6 m (20 ft) which is 2% of the total depth. The estimated depth to the 5,000 mg/L isochlor from the TDEM study (826 ft bls) is not in agreement with the data from monitor well, P-0735.

The resistivity of this layer has a range of from 3.2 - 4.2 ohm-m. This corresponds to a range in interpreted chloride content of from 9,898 mg/L to 7,505 mg/L, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence range of the resistivity of layer 1 is from 45 - 49 ohm-m. A corresponding chloride concentration cannot be determined because layer 1 is in part comprised of the Hawthorn Group and surficial sediments. Accordingly, equation (4) may not be valid.

5.9.6 <u>Summary of TDEM Sounding at Union Camp (Site 8)</u>

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 826 ft (-810 ft msl) and occur within the Lower Floridan aquifer. The estimated depth to the 5,000 mg/L isochlor from the TDEM study does not agree with the estimated depth from a nearby monitor well. No interference sources appear to be present near the Tx coil that might have affected the quality of survey data. The chloride content below that depth is inferred to be 8,540 mg/L.
- The quality of ground water within the Floridan aquifer at this site cannot be interpreted because the analysis of the TDEM data does not allow the Hawthorn Group and surficial aquifer system to be distinguished from the Floridan Aquifer System.












CLIE LOCATI COUN PROJE LOOP SI COIL I SOUNDIN	ENT: SJWRM LON: UNION NTY: PUTNA ECT: SALTW LZE: 35 LOC: NG COORDIN	D I CAMP M COUNTY, FLOR NATER INTERFACE 0.999 m by 0.000 m (X), NATES: E:	S IDA EL DETECTION EQ 350.999 m 0.000 m (Y) 0.0000 N:	DATE: 02-05- OUNDING: 1 EVATION: 5 UIPMENT: Geonic AZIMUTH: 0.0000	-93 .00 m 28 PROTEM
	F	ITTING ERROR:	3.713 PERC	ENT	
L #	RESISTIVI (ohm-m)	TY THICKNES (meters)	S ELEVATIC (meters	N CONDUCTA	ANCE 15)
1 2	47.17	251.9	5.00 -246.9	5.34	
ALL PA	ARAMETERS	ARE FREE			
PARAM	METER BOUN	IDS FROM EQUIVA	LENCE ANALYSIS		
LAYEI	R MINI	MUM BEST	MAXIMUM		
RHO	1 45 2 3	5.48347.173.2353.68	0 49.010 9 4.212		
THICK	1 245	5.972 1.00	0 257.518		
DEPTH	1 245	5.972 251.94	9 257.518		
CURI FREQUI	RENT: 19 ENCY: 7	9.20 AMPS EM- 7.50 Hz GAIN:	57 COIL AREA 6 RAMP TIME	A: 100.00 sq S: 252.00 muSE	m. C
No.	TIME (ms)	en DATA	f (nV/m sqrd) SYNTHET:	DIFFERE [C (perce	NCE nt)
1	0.698	3969.7	4092.6	-3.09	MASKED
3	1.10	1462.2	1452.9	-7.68	8
4 5	$1.40 \\ 1.75$	918.7 611.0	885.0 579.3	3.66 5.18	
6 7	2.22	402.8	388.1 273.1	3.64	3
·					•
		4			
		,			
ST. JOH	INS RIVER	WATER	SDII	TDEM SOUNDIN SOUND UNION	G DATA TABLE ING 8 CAMP
	MENT DIST A. FLORIDA	RICI	SUBSURFACE DETECTION	PROJECT NO .:	93742
	, , , , , , , , , , , , , , , , , , , ,		INVESTIGATIONS INCORPORATED	TABLE	5.9-1

NO. TII (ms	1E e 3) DAT	mf (nV/m sqrd) A SYNTHET	DIFFERENC IC (percent	CE 2)
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	202.8 148.7 6 102.8 7 72.27 8 51.42 0 34.46 9 22.93 8 15.28	-2.58 -4.13 -3.59 -3.28 -1.66 2.91 3.18 3.14	
17 28.3	LO 5.9	6 5.90	0.905	
CURRENT: FREQUENCY:	19.20 AMPS EM 3.00 Hz GAIN	I-57 COIL ARE I: 8 RAMP TIM	A: 100.00 sq m E: 252.00 muSEC	n
No. TII (ma	1E e 5) DAI	mf (nV/m sqrd) A SYNTHET	DIFFEREN(IC (percent	CE t)
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	2586.3 1586.0 934.0 590.4 407.4 280.9 200.8 147.8 105.0 8 75.78 8 55.91 8 39.94 9 26.60 9 17.74	-6.94 -0.669 4.11 6.36 3.52 3.22 -1.70 -4.22 -3.60 -7.550E-04 -2.44 -0.897 -0.7901 -0.2881	MASKED MASKED
PARAMETER RI "F" INDICAT P 1 0.99 P 2 -0.01 T 1 0.00 P 1	ESOLUTION MATRIX: ES FIXED PARAMETE 0.93 0.01 1.00 P 2 T 1	R		
ST. JOHNS RIVE	R WATER		TDEM SOUNDING SOUNDIN UNION C	DATA TABLE G 8 AMP
PALATKA, FLORI	DA	DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: TABLE	93742 5.91

5.10 TDEM SITE 9 - DRAYTON ISLAND SITE

5.10.1 Location Description and Geoelectrical Section

The site is located on Drayton Island which is within the St. Johns River in Putnam County, Florida (Figure 5-10.1). Powerlines were present to the south of the Tx Loop. QA soundings were performed 80 ft to the south and 95 ft to the north of the initial Rx coil position. Results from the QA soundings indicated that the apparent resistivity values were not affected by any interference sources (Figure 5.10-4).

The Floridan aquifer begins at an approximate depth of 50 ft below msl and is overlain by the Hawthorn Group and the surficial aquifer system (Scott et al., 1991). The Lower Floridan aquifer begins at an approximate depth of 700 ft below msl (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.10-2. The interpreted geoelectrical section consists of a three-layer subsurface.

5.10.2 Geological Interpretation of Geoelectrical Model

There is insufficient electrical resistivity contrast between the upper surficial aquifer system layer, the Hawthorn Group and the underlying Floridan aquifer to distinguish the three. Fixing the thickness of the upper layer does not resolve this dilemma; therefore it can be interpreted that there exists a two-layer geoelectrical section with a relatively thick (126.8 m = 416 ft) surface layer of intermediate to low resistivity (22.7 ohm-m) overlying a low resistivity layer (2.3 ohm-m). It can be interpreted that the upper surficial aquifer system, the Hawthorn Group and part of the Floridan Aquifer System exist as a combined but indistinguishable (geoelectrically) layer, overlying a saltwater saturated Floridan aquifer at a depth of 416 ft bls.

5.10.3 Depth to Occurrence of Salt Water

The bottom (second) layer of the geoelectrical model, with a resistivity of 2.3 ohm-m, is interpreted to represent salt water. It occurs at a depth of 416 ft (-406 ft msl). Because the resistivity of layer 1 (22.7 ohm-m) is less than 80 ohm-m, the interpreted depth to the 5,000 mg/L isochlor is equal to the depth of the geoelectrical interface, or at 416 ft depth (-406 ft msl). The resistivity of layer 2 (2.3 ohm-m) corresponds to a chloride content in excess of 13,831 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2.

5.10.4 Depth of Occurrence of the 250 mg/L Isochlor

Because of the inability to segregate the Floridan aquifer from the overlying surficial aquifer system and the Hawthorn Group, the effective chloride concentration of layer 1 cannot be determined.

5.10.5 Accuracy of Measurement and Interpretation

Figure 5.10-3 is the equivalence analysis at this site and the inversion table (Table 5.10-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model. The range of equivalence in determining the depth to the low resistivity layer is about $\pm 4 \text{ m}$ (13 ft) which is 3% of the total depth. The resistivity of this layer has a range of from 2.2 - 2.5 ohm-m. This corresponds to a range in interpreted chloride content of from 14,467 mg/L to 12,712 mg/L, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence range of the resistivity of layer 1 is from 21.8 - 23.6 ohm-m. A corresponding chloride concentration cannot be determined because layer 1 is in part comprised of the Hawthorn group and surficial sediments. Accordingly, equation (4) may not be valid. Chloride concentrations of less than 250 mg/L have been mapped in the area of the site by SJRWMD.

5.10.6 <u>Summary of TDEM Sounding at Drayton Island (Site 9)</u>

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 416 ft (-406 ft msl) and occur within the Floridan aquifer. The chloride content below that depth is inferred to be in excess of 10,000 mg/L.
- The quality of ground water within the Floridan aquifer at this site cannot be interpreted because analysis of the TDEM data does not allow the surficial aquifer system and the Hawthorn Group to be distinguished from the Floridan Aquifer System.



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CLIE LOCATI COUN PROJE LOOP SI COIL I SOUNDIN	ENT: (ION: 1 NTY: 1 SCT: (IZE: LOC: NG COO	SJWRMD DRAYTON IS PUTNAM COU SALTWATER 452.000 0.000 ORDINATES	SLAND JNTY, FLORII INTERFACE I) m by J) m (X), : E:	SOU DA ELEX DETECTION EQUI 182.800 m A2 0.000 m (Y) 0.0000 N:	DATE: 03-05-93 JNDING: 1 VATION: 3.00 m IPMENT: Geonics PROTEM ZIMUTH: 0.0000
		FITTI	NG ERROR:	3.407 PERCEN	ЯТ
L #	RESI: (O	STIVITY hm-m)	THICKNESS (meters)	ELEVATION (meters)	CONDUCTANCE (Siemens)
1 2	2	2.66 2.34	126.8	3.00 -123.8	5.59
ALL PA	ARAME	TERS ARE I	FREE		
PARAN	ÆTER	BOUNDS FI	ROM EQUIVALE	INCE ANALYSIS	
LAYER	ર	MINIMUM	BEST	MAXIMUM	
RHO	1 2	21.767 2.177	22.663 2.341	23.637 2.523	
THICK	1	123.209	1.000	130.573	
DEPTH	1	123.209	126.884	130.573	
CURI FREQUE	RENT: ENCY:	16.10 2 30.00 1	AMPS EM-57 Hz GAIN: 3	COIL AREA: RAMP TIME:	100.00 sq m. 192.00 muSEC
No.	Т: (1	IME ms)	emf DATA	(nV/m sqrd) SYNTHETIC	DIFFERENCE (percent)
1 2 3 4 5 6 7	0 0 0 0 1 1	.351 .438 .558 .702 .858 .06 .37	31561.9 15864.9 8685.2 5332.2 3733.2 2651.1 1813.8	21426.1 14053.4 8732.9 5583.5 3841.5 2644.8 1770.5	32.11 MASKED 11.41 MASKED -0.549 -4.71 -2.90 0.238MASKED 2.38 MASKED
OT 1011			<u> </u>	SDII	TDEM SOUNDING DATA TAN

ST. JOHNS RIVER WATER MANAGEMENT DISTRICT PALATKA, FLORIDA



TDEM SOUNDING DATA TABLE SOUNDING 9 DRAYTON ISLAND PROJECT NO.: 93742 TABLE 5.10-1

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No.	TIM	Ē		emf	(nV/m sqr	:d)	DIFFER	ENCE
	(ms)	1	DATA	SYNT	HETIC	(perce	ent)
8	17	А	125	0 5	1000	2	2 1	C NA CERED
0	2 1	7	125	2.2	1232		2.10	D MASKED
9	2.1	7	90.	2.1	889	.4	1.4	U MASKED
10	2.1	/	62	b.7	613	• 1	2.1	7 MASKED
11	3.5	0	42	0.4	425	.0	-1.0	9 MASKED
CUR FREQU	RENT: ENCY:	16.10 A 7.50 H	MPS Iz G	EM-57 AIN: 3	COIL RAMP	AREA: TIME:	100.00 so 192.00 muS	q m. EC
N -				-	· ·	-		
NO.	TIM (ms	E)	1	emf DATA	(nV/m sqr SYNI	d) HETIC	DIFFER (perce	ENCE ent)
12	с Л З	46	3407	3 0	22140	5	25.20	A NACVED
12	0.5	20	1776	5.0	22140		35.3	9 MASKED
13	0.4	<i>21</i> 50	L//D	2.8	14897	• 2	10.0	B MASKED
14	0.5	50	931	9.5	9116	.4	2.1	7
15	0.6	98	567	4.1	5772		-1.7	3
16	0.8	69	380	7.2	3878	.3	-1.8	6
17	1.1	0	268	8.2	2600	1.1	3.2	7
18	1.4	0	188	3.0	1818	3.2	3.4	4
19	1.7	5	137	1.6	1327	1.1	3.2	4
20	2.2	2	97.	9.5	956	5.6	2.3	3
21	2.7	9	69	6.1	694	.3	0.2	57
22	3.4	2	51	4.0	520	.4	-1.2	5
23	4.2	6	36	8.2	376	.7	-2.2	9
24	5.4	9	24	6.8	254	.8	-3.2	6
25	6.9	6	16	5.2	173	.8	-5.2	2
26	8.6	6	11.	3.9	119	. 8	-5.1	4
27	11.0	6	7	5.91	77	46	-2.0	4
28	14.0	Ō	4	9.14	49	. 62	-0.9	62
29	17.4	7	3	2.46	31	95	-0.5	7
30	22.2	3	2	0.43	19	25	5.7	6
31	28.1	0	1	2.37	11	.44	7.4	9
CUR FREQU	RENT: ENCY:	16.10 A 3.00 H	AMPS Iz G	EM-57 AIN: 8	COIL RAMP	AREA: TIME:	100.00 so 192.00 muS	q m. EC
No.	TIM	E		emf	(nV/m sqr	d)	DIFFER	ENCE
	(ms)]	DATA	SYNT	HETIC	(perce	ent)
32	0.8	57	394	9.2	3983	.1	-0.8	59
32 33	0.8	57 6	394 284	9.2 2.9	3983 2773	8.1 8.4	-0.8 2.4	59 4
ST. JOHN MANAGEM	IS RIVER	WATER TRICT		S	SDII UBSURFACE		TDEM SOUNDIN SOUND DRAYTON	g data table Ing 9 Island
PALATKA.	FLORID	م		D	ETECTION		PROJECT NO .:	93742
· ··· - ·,				l li	VESTIGATION	IS	TARI F	5 10-1
				11	NCORPORATE	D	IADLE	5.10-1

No.	TIME (ms)	er DAT?	nf (nV/m sqrd) A SYNTHI) ETIC	DIFFEREN (percen	ICE It)
34 35 36 37 38 39 40 41 42 43 44 45 46 47 48 49	1.37 1.74 2.17 2.77 3.50 4.37 5.56 6.98 8.56 10.64 13.70 17.40 21.70 27.70 35.00 43.70	$1969.4 \\ 1407.2 \\ 1013.6 \\ 731.3 \\ 510.3 \\ 363.1 \\ 250.9 \\ 176.0 \\ 121.0 \\ 85.12 \\ 55.18 \\ 35.73 \\ 24.81 \\ 14.84 \\ 8.54 \\ 4.64 \end{bmatrix}$	1892.9 1346.9 996.7 710.7 511.8 370.3 257.2 179.8 128.5 36.2 36.2 14.4 8.8 5.4	5 9 1 1 8 3 2 3 5 7 6 5 4 2 5 1 4 2 5 1 4 1	$\begin{array}{r} 3.90 \\ 4.29 \\ 1.72 \\ 2.90 \\ -0.284 \\ -1.97 \\ -2.49 \\ -2.18 \\ -6.19 \\ -4.28 \\ -2.47 \\ -1.44 \\ 4.83 \\ 2.80 \\ -3.07 \\ -16.45 \end{array}$	MASKED MASKED
PARAM "F" I P 1 P 2 T 1	ETER RESOLUTIO NDICATES FIXED 1.00 0.00 0.98 0.00 0.00 1. P 1 P 2	ON MATRIX: D PARAMETER 00 T 1	2			
ST. JOHN MANAGEM	S RIVER WATER ENT DISTRICT		SUBSURFACE	TDEM	SOUNDING SOUNDING DRAYTON IS	DATA TABLE 9 LAND
PALATKA,	FLORIDA		DETECTION INVESTIGATIONS INCORPORATED	PROJEC TABLE	CT NO.:	93742 5.10-1

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5.11 TDEM SITE 10 - BEAR ISLAND SITE

5.11.1 Location Description and Geoelectrical Section

The site is located on an island within Crescent Lake, east of the St. Johns River in Flagler County, Florida (Figure 5.11-1). No apparent sources of interference were visible.

The Floridan aquifer begins at an approximate depth of 75 ft below msl and is overlain by the Hawthorn Group and surficial sediments (Scott et al., 1991). The Lower Floridan aquifer begins at an approximate depth of 800 ft below msl (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.11-2. The interpreted geoelectrical section consists of a three-layer subsurface.

5.11.2 <u>Geological Interpretation of Geoelectrical Model</u>

The best-fit model shows a three-layer subsurface with the uppermost layer (presumed Hawthorn Group and surficial sediments; 6 ohm-m) overlying a second layer (Floridan aquifer; 98 ohm-m). Layer 1 is interpreted to represent the Hawthorn Group and surficial sediments. Layer 1 is 22.6 m (74 ft) thick which corresponds well to published thickness of the Hawthorn Group and surficial sediments (78 ft) in this area (Scott et al., 1991). Layer 2 is presumed to represent the freshwater saturated Floridan aquifer. These two layers are underlain by a low resistivity layer (2.8 ohm-m) at a depth of 164 m (538 ft) which is presumed to represent the saturated Floridan aquifer.

5.11.3 Depth to Occurrence of Salt Water

The bottom (third) layer of the geoelectrical model, with a resistivity of 2.8 ohm-m, is interpreted to represent salt water. It occurs at a depth of 538 ft (-535 ft msl). Because the resistivity of layer 2 (98 ohm-m) is interpreted to represent fresh water within the Floridan aquifer (i.e., is greater than 80 ohm-m), the interpreted depth to the 5,000 mg/L isochlor is taken as 50 ft greater than the depth of the geoelectrical interface, or at 588 ft depth (-585 ft msl). The resistivity of layer 3 (2.8 ohm-m) corresponds to a chloride content of 11,334 mg/L, assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2.

5.11.4 Depth of Occurrence of the 250 mg/L Isochlor

The resistivity of layer 2, 98 ohm-m, corresponds to a chloride content less than 250 mg/L, assuming a 25% porosity and the validity and applicability of equation (4) of Section 4.2. The 250 mg/L isochlor is placed in the Floridan aquifer at a depth 50 ft above the layer 3 interface or at 488 ft bls (-485 ft msl).

5.11.5 Accuracy of Measurement and Interpretation

Figure 5.11-3 is the equivalence analysis at this site and the inversion table (Table 5.11-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about \pm 6 m (20 ft) which is 4% of the total depth. The resistivity of this layer has a range of from 2.5 - 3.1 ohm-m. This corresponds to a range in interpreted chloride content of from 12,712 mg/L to 10,222 mg/L, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence range of the resistivity of layer 2 is from 61 - 180 ohm-m which corresponds to a chloride content ranging from less than 250 mg/L to greater than 250 mg/L. The chloride-to-sulfate ratio at the site is 2:1 (Table 5.1-4), rather than 5:1. Accordingly, equation (4) may not be valid. Chloride concentrations are not known on the island. Chloride concentrations have been mapped above 250 mg/L east of the island and less than 250 mg/L west of the island (SJRWMD, personnel communication).

5.11.6 Summary of TDEM Sounding at Bear Island (Site 10)

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 588 ft (-535 ft msl) and occur within the Floridan aquifer. The chloride content below that depth is inferred to be in excess of 10,000 mg/L.
- The ground water within the Floridan aquifer at this site is interpreted to contain an average chloride concentration below 250 mg/L. The 250 mg/L isochlor is interpreted to be present within the Floridan aquifer at a depth of 488 ft bls. Water quality on the island is unknown.







CLII LOCAT COUN PROJI LOOP SJ COIL I SOUNDIN	ENT: S ION: E NTY: F SCT: S IZE: LOC: NG COC	JWRMD EAR ISLAN UTNAM/FLA ALTWATER 457.200 0.000 RDINATES:	D GLER COUNT INTERFACE m by m (X), E:	SC Y, FL. ELH DETECTION EQU 39.600 m A 0.000 m (Y) 0.0000 N:	DATE: 05-04-93 DUNDING: 1 SVATION: 1.00 JIPMENT: Geonics AZIMUTH: 0.0000	m PROTEM
		FITTIN	G ERROR:	3.387 PERCI	ENT	
L #	RESIE (or	STIVITY 1m-m)	THICKNESS (meters)	ELEVATIO	N CONDUCTANC) (Siemens)	E
1 2 3	98	5.88 3.44 2.76	22.62 141.3	1.00 -21.62 -162.9	3.84 1.43	
ALL P.	ARAME	TERS ARE F	REE			
PARA	METER	BOUNDS FR	OM EQUIVAL	ENCE ANALYSIS		
LAYE	R	MINIMUM	BEST	MAXIMUM	· .	
RHO	1 2 3	3.396 61.457 2.513	5.885 98.441 2.770	6.886 180.317 3.123		
THICK	1 2	10.774 134.149	-1.040 1.000) 27.960) 150.312		
DEPTH	1 2	10.774 158.029	22.623 163.933	27.960 3 169.473		
CUR FREQU	RENT: ENCY:	18.50 A 7.50 H	MPS EM-5 Iz GAIN:	57 COIL AREA 4 RAMP TIME	: 100.00 sq m. : 182.00 muSEC	,
No.	Т (1	IME ms)	emi DATA	f (nV/m sqrd) SYNTHETI	C (percent)	3)
1 2 3	0 0 0	.346 .427 .550	26244.6 14731.2 7400.6	24219.6 14247.7 7374.5	7.71 3.28 0.351	
ST. JO		IVER WATER	٤		TDEM SOUNDING SOUNDING BEAR ISLA	DATA TABLE 10 ND
PALATK	A, FLC	RIDA		DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: TABLE	93742 5.11-1

No.	TIME (ms)	em DATA	f (nV/m sqrd) SYNTHET	DIFFERE CIC (perce	NCE nt)
4 5 6 7 8 9 10 11 12 13 14 15 16 17 18 19 20	0.698 0.869 1.10 1.40 1.75 2.22 2.79 3.42 4.26 5.49 6.96 8.66 11.06 14.00 17.47 22.23 28.10	3763.7 1980.3 1049.4 550.8 327.3 199.6 127.9 90.27 63.17 41.38 27.49 18.49 12.14 7.84 5.24 3.34 1.99	3856.9 2081.2 1061.6 554.9 327.7 193.1 128.8 89.53 63.53 41.79 28.44 19.52 12.63 8.12 5.26 3.18 1.90	$\begin{array}{c} -2.47\\ -5.09\\ -1.16\\ -0.74\\ -0.11\\ 3.25\\ -0.67\\ 0.81\\ -0.56\\ -0.99\\ -3.45\\ -5.58\\ -4.01\\ -3.56\\ -0.26\\ -0.26\\ -0.26\\ -0.26\\ -0.26\\ -0.56\\ -0.26\\ -0.58\\ -0.56\\ -0.58\\ -0.56\\ -0.58\\ -0.56\\ -0.58\\ -0.56\\ -0.58\\ -0.56\\ -0.56\\ -0.58\\ -0.56\\ -0.5$	9 0 9 6 3 7
CUR FREQU	RENT: 18.50 ENCY: 3.00	AMPS EM- Hz GAIN:	57 COIL ARE 8 RAMP TIN	EA: 100.00 sq ME: 182.00 muSE	[m. C
No.	TIME (ms)	em DATA	f (nV/m sqrd) SYNTHEN	DIFFERE FIC (perce	INCE ent)
21 22 23 24 25 26 27 28 29 30 31 32 33 34 35 PARAM "F" J	0.857 1.06 1.37 1.74 2.17 2.77 3.50 4.37 5.56 6.98 8.56 10.64 13.70 17.40 21.70 METER RESOLUTION INDICATES FIXED	2057.8 1154.6 597.3 339.6 209.9 137.6 90.28 62.91 42.61 29.74 19.76 14.04 9.14 6.04 4.12	2163.9 1187.7 590.8 333.6 204.0 131.8 87.4 62.30 42.20 42.20 29.40 20.90 14.5 9.20 5.90 3.90	$ \begin{array}{r} -5.15\\ -2.86\\ 1.07\\ 1.76\\ 2.81\\ 4.19\\ 7 3.11\\ 6 0.87\\ 0 0.96\\ 6 0.92\\ 8 -6.16\\ 1 -3.32\\ 9 -1.56\\ 9 0.82\\ 2 4.75\\ \end{array} $	5 7 5 1 7 3 4 8 3 7 5 2 6 5 6 9
P 1 P 2 P 3 T 1 - T 2	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	93 02 0.88 02 0.01 P 3 T 1	0.99 T 2		
ST. JOHN MANAGEM	S RIVER WATER		SUBSURFACE	TDEM SOUNDING SOUNDIN BEAR ISI	G DATA TABLE G 10 _AND
PALATKA,	FLORIDA		INVESTIGATIONS INCORPORATED	PROJECT NO.: TABLE	93742 5.11-1

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5.12 TDEM SITE 11 - DESERET #1 SITE

5.12.1 Location Description and Geoelectrical Section

The site is located in central Orange County, Florida (Figure 5.12-1) and is 3.3 miles east of TDEM site 12. The site is within a pasture with no obvious signs of interference. The site is approximately 1,200 ft from public supply wells Cocoa 1 and Cocoa 7 (Figure 5.12-1) with depths of 374 ft and 379 ft, respectively. This site is approximately 3.5 miles east of Cocoa C (a nested set of monitor wells) with a maximum depth of 1,357 ft. Water quality results from Cocoa C show a measured chloride concentration of less than 250 mg/L to a depth of at least 1,224 ft and a value of 2,700 mg/L at a depth of 1,357 ft (Figure 5.12-4).

The Floridan aquifer begins at a depth of approximately 170 ft below msl and is overlain by the Hawthorn Group and surficial sediments (Scott et al., 1991). The top of the Lower Floridan begins at an approximate depth of 1100 ft below msl. The base of the Floridan aquifer is approximately 2,600 ft below msl in this area (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.12-2. The interpreted geoelectrical section consists of a three-layer subsurface.

5.12.2 Geological Interpretation of Geoelectrical Model

There is a sufficient electrical resistivity contrast to distinguish two geological layers above the third saltwater saturated layer. The first layer occurs at a depth of 56 m (184 ft) and not at the hydrostratigraphic contact (239 ft bls) between the Hawthorn Group and the Floridan aquifer System. The first layer has a low-resistivity value (18 ohm-m) and is considered to represent the upper portion of the Hawthorn Group and surficial sediments. The second layer has a highresistivity value (101 ohm-m) and a thickness of 329 m (1,079 ft). It is considered to represent a combined but indistinguishable (geoelectrical) layer consisting of a portion of the Hawthorn Group and part of the Floridan aquifer. The third layer is considered to represent a saltwater saturated Floridan aquifer at a depth of 1,263 ft. The depth to the interpreted low resistivity (saltwater) layer at 1,263 ft is in good agreement with the position of the higher chloride value seen in the Cocoa C well.

5.12.3 Depth to Occurrence of Salt Water

The bottom (third) layer of the geoelectrical model, with a resistivity of 10.7 ohm-m, is interpreted to represent moderately salty water. It occurs at a depth of 1,263 ft (-1,194 ft msl). This is in good agreement with the water quality results from the Cocoa C well, 3.5 miles west of the site. Because the resistivity of layer 2 (101 ohm-m) is interpreted to represent fresh water within the Floridan aquifer (i.e., is greater tha 80 ohm-m), the interpreted depth to the 5,000 mg/L isochlor is 50 ft below the depth of the geoelectrical interface, or at 1,313 ft depth (-1,244 ft msl). The resistivity of layer 3 (10.7 ohm-m) corresponds to a chloride concentration of 2,853 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2. It is presumed that because of expected high chlorinity gradients, this value is sufficiently close to 5,000 mg/L that they represent the same effective depth.

5.12.4 Depth of Occurrence of the 250 mg/L Isochlor

Because of the inability to segregate the Floridan aquifer from part of the overlying Hawthorn Group, the effective chloride concentration of Layer 2 cannot be calculated. Water quality results from Cocoa C well places the 250 mg/L isochlor at approximatly 1,250 ft bls (Figure 5.12-4)

5.12.5 Accuracy of Measurement and Interpretation

Figure 5.12-3 is the equivalence analysis at this site, and the inversion table (Table 5.12-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model. Both the estimated depth to the 5,000 mg/L and 250 mg/L isochlor and chloride concentration from the TDEM study are in agreement with the results from nearby wells.

The range of equivalence in determining the depth to the low resistivity layer is about $\pm 26 \text{ m}$ (85 ft) which is 7% of the total depth. The resistivity of this layer has a range of from 9.3 - 12.7 ohm-m. This corresponds to a range in interpreted chloride content of from 3,305 mg/L to 2,380 mg/L, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence of the resisitivity of Layer 2 is from 80 - 140 ohm-m. A corresponding chloride concentration cannot be determined because Layer 2 is in part comprised of the Hawthorn Group.

5.12.6 Summary of TDEM Sounding at Deseret #1 (Site 11)

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 1,313 ft (-1,244 ft msl) in the Lower Floridan aquifer. The estimated depth agrees with water quality results from a nearby well. The measured layer resistivity at this level yields a chloride concentration of 2,853 mg/L. It is assumed that this is sufficiently close to 5,000 mg/L that the derived depth is applicable to the saltwater interface depth.
- The ground water within the Floridan aquifer at this site cannot be interpreted because analysis of the TDEM data does not allow part of the Hawthorn Group to be distinguished from the Floridian Aquifer System.









CLI LOCAT COU PROJ LOOP S COIL SOUNDI	ENT: S ION: D NTY: C ECT: S IZE: LOC: NG COC	GJWRMD DESERET #1 DRANGE COU GALT WATER 305.000 0.000 DRDINATES:	NTY, FLORID INTERFACE m by 4 m (X), E:	SO A ELE DETECTION EQU 57.000 m A 0.000 m (Y) 0.0000 N:	DATE: 05-05-93 UNDING: 1 VATION: 21.00 m IPMENT: Geonics PROTEM ZIMUTH: 0.0000
		FITTIN	G ERROR:	4.192 PERCE	NT
L #	RESIS (oh	STIVITY Dm-m)	THICKNESS (meters)	ELEVATION (meters)	CONDUCTANCE (Siemens)
1 2 3	18 101 10	8.11 .1 .69	55.83 328.9	21.00 -34.83 -363.8	3.08 3.25
ALL P	ARAMEI	ERS ARE F	REE		
PARA	METER	BOUNDS FR	OM EQUIVALE	NCE ANALYSIS	. *
LAYE	R	MINIMUM	BEST	MAXIMUM	
RHO	1 2 3	14.708 80.136 9.336	18.115 101.110 10.691	20.610 140.288 12.683	
THICK	1 2	39.917 292.928	-1.146 1.000	73.873 357.231	
DEPTH	1 2	39.917 356.714	55.840 384.818	73.873 408.038	
CUR FREQU	RENT: ENCY:	17.20 A 7.50 H	MPS EM-57 z GAIN:3	COIL AREA: RAMP TIME:	100.00 sq m. 242.00 muSEC
No.	T] (1	lme as)	emf DATA	(nV/m sqrd) SYNTHETIC	DIFFERENCE (percent)
1 2 3	0 . 0 . 0 .	.346 .427 .550	64179.4 39473.3 20695.3	38094.6 26920.8 16616.4	40.64 MASKED 31.80 MASKED 19.70 MASKED
ST.	JOHNS	RIVER WAT	ER	SUBSUBEACE	TDEM SOUNDING DATA TABLE SOUNDING 11 DESERET #1
MAN PAL	AGEMEN ATKA, F	LORIDA		DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: 93742 TABLE 5.12-1

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No.	TIME	emf	(nV/m sqrd)	DIFF	ERENCE
	(ms)	DATA	SYNTHE	TIC (pe	ercent)
4	0.698	10900.0	9911.9	9	.06
5	0.869	5854.1	5907.2	0	.908
6	1.10	3119.7	3174.0	1	.73
7	1.40	1585.9	1672.4	-5	.45
8	1.75	869.5	905.2	-4	.10
9	2.22	464.6	459.1		.18
10	2.79	251.1	246.5		.85
11	3.42	151.3	144.0		.79
12	4.26	88.32	84.52	2 4	.30
13	5.49	48.73	48.44	8 0	.507
14	6.96	28.86	29.72	2 -2	.96
15	8.66	18.59	19.44	8 -4	.80
16	11.06	11.45	12.0	7 -5	.44
17	14.00	7.27	7.6	0 –4	.57
18	17.47	4.68	4.8	4 –3	3.31
19	22.23	3.09	2.9	1 5	5.90
20	28.10	1.82	1.7	3 5	5.02
CURR	ENT: 17.20	AMPS EM-5	7 COIL ARI	EA: 100.00) sq m.
FREQUE	ENCY: 3.00	Hz GAIN:	7 RAMP TII	ME: 242.00 n	NUSEC
No.	TIME	emf	(nV/m sqrd)	DIFE	FERENCE
	(ms)	DATA	SYNTHE	TIC (Pe	ercent)
21	0.857	6062.5	6112.8	$ \begin{array}{c} -0 \\ -3 \\ -4 \\ -2 \\ -4 \\ -2 \\ -2 \\ -4 \\ -2 \\ 0 \\ -2 \\ -2 \\ 0 \\ -2 \\ -2 \\ 0 \\ -2 \\ -2 \\ -2 \\ -2 \\ -2 \\ -2 \\ -2 \\ -2$	0.829
22	1.06	3390.4	3517.9		3.76
23	1.37	1706.4	1786.4		4.68
24	1.74	895.2	921.3		2.91
25	2.17	491.8	491.1		0.132
26	2.77	272.3	252.6		7.24 MASKED
27	3.50	145.6	137.0		5.92
28	4.37	83.59	80.9		3.16
29	5.56	48.17	48.3		0.295
30	6.98	29.80	30.6		2.69
31	8.56	17.99	20.9		5.16 MASKED
32	10.64	11.99	13.8		5.63 MASKED
33	13.70	7.19	8.6		0.68 MASKED
34	17.40	4.49	5.5		2.42 MASKED
PARAMI "F" II	ETER RESOLUTI NDICATES FIXE	ON MATRIX: D PARAMETER			:
P 1 P 2 - (P 3 T 1 - (T 2	0.95 0.04 0.53 0.01 -0.10 (0.08 -0.19 -(0.01 0.11 (Pl P2	0.84 0.01 0.82 0.06 0.05 P 3 T 1	0.96 T 2		, , ,
ST. JOH MANAGE	NS RIVER WAT	ER	SUBSURFACE	TDEM SOUNDI SOUND DESER	NG DATA TABLE ING 11 ET #1
PALATKA	, FLORIDA		DETECTION INVESTIGATIONS	PROJECT NO .: TABLE	93742 5.12-1

5.13 TDEM SITE 12 - DESERET #2 SITE

5.13.1 Location Description and Geoelectrical Section

The site is located in central Orange County, Florida (Figure 5.13-1) and is 3.3 miles west of TDEM site 11. The site was within an abandoned orchard; no obvious signs of interference sources were present. The site is approximately 1,000 ft away from public supply wells Cocoa 14 and Cocoa 15 and 1,200 ft away from Cocoa C (Figure 5.13-1). The maximum depth of each of the wells is 761, 702 and 1,351 ft, respectively. Water quality results from the Cocoa C well show a measured chloride concentration of less than 250 mg/L to a depth of at least 1,224 ft and a value of 2,700 mg/L at a depth of 1,357 ft.

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The Floridan aquifer begins at a depth of approximately 170 ft below msl and is overlain by the Hawthorn Group and surficial sediments (Scott et al., 1991). The top of the Lower Floridan begins at an approximate depth of 1,100 ft below msl. The base of the Floridan aquifer occurs at approximately 2,600 ft below msl in this area (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.13-2. The interpreted geoelectrical section consists of a three-layer subsurface, similar to Site 11.

5.13.2 Geological Interpretation of Geoelectrical Model

There is a sufficient electrical resistivity contrast to distinguish two geological layers above the third saltwater saturated layer. The first layer occurs at a depth of 46.2 m (152 ft) and not at the hydrostratigraphic contact (239 ft bls) between the Hawthorn Group and the Floridan aquifer System. The first layer has a low-resistivity value (16.5 ohm-m) and is considered to represent the upper portion of the Hawthorn Group and surficial sediments. The second layer has a high-resistivity value (957 ohm-m) and a thickness of 314 m (1,030 ft) and is considered to represent a combined but indistinguishable (geoelectrical) layer consisting of a portion of the Hawthorn Group and part of the Floridan aquifer. The third layer is considered to represent a saltwater saturated Floridan aquifer at a depth of 1,181 ft. The depth to the interpreted low resistivity (saltwater) layer at 1,181 ft is in agreement with the position of a higher chloride value in the Cocoa C well and with the 1,263 ft depth at nearby Site 11.

5.13.3 Depth to Occurrence of Salt Water

The bottom (third) layer of the geoelectrical model, with a resistivity of 15.5 ohm-m, is interpreted to represent moderately salty water. It occurs at a depth of 1,181 ft (-1,112 ft msl). This is in agreement with water quality results from the Cocoa C well, 1,200 ft away. Because the resistivity of layer 2 (957 ohm-m) is interpreted to represent freshwater within the Floridan aquifer (i.e., is greater than 80 ohm-m), the interpreted depth to the 5,000 mg/L isochlor is 50 ft below the depth of the geoelectrical interface, or at 1,231 ft depth (-1,162 ft msl). The resistivity of layer 3 (15.5 ohm-m) corresponds to a chloride concentration of 1,922 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2. It is presumed that because of expected high chlorinity gradients, this value is sufficiently close to 5,000 mg/L that they represent the same effective depth.

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5.13.4 Depth of Occurrence of the 250 mg/L Isochlor

Because of the inability to segregate the Floridan aquifer from part of the overlying Hawthorn Group, the effective chloride concentration of Layer 2 cannot be calculated. Water quality results from the Cocoa C well places the 250 mg/L isochlor at approximately 1,250 ft bls (Figure 5.13-4).

5.13.5 Accuracy of Measurement and Interpretation

Figure 5.13-3 is the equivalence analysis at this site, and the inversion table (Table 5.13-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about $\pm 25 \text{ m}$ (82 ft) which is 7% of the total depth. The resistivity of this layer has a range of from 13.5 - 20.1 ohm-m. This corresponds to a range in interpreted chloride content of from 2,229 mg/L to 1,447 mg/L, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence range of the resistivity of layer 2 is from 538 - 3,027 ohm-m. A corresponding chloride concentration cannot be determined because Layer 2 is in part comprised of the Hawthorn Group.

5.13.6 Summary of TDEM Sounding at Deseret #2 (Site 12)

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 1,231 ft (-1,162 ft msl) and occur in the Lower Floridan aquifer. The estimated depth correlates well to the estimated depth based on water quality in a nearby well. The measured layer resistivity at this level yields a chloride concentration of only 1,922 mg/L, but we assume that this is sufficiently close to 5,000 mg/L that the derived depth is applicable to the saltwater interface depth.
- The ground water within the Floridan aquifer at this site cannot be interpreted because analysis of the TDEM data does not allow part of the overlying Hawthorn Group to be distinguished from the Floridian Aquifer System.








CLI LOCAI COU PROJ LOOP S COIL SOUNDI	ENT: ION: NTY: ECT: IZE: LOC: NG CC	SJWRMD DESERET #2 ORANGE COU SALTWATER 480.000 0.000 0RDINATES:	NTY, FLOF INTERFACE m by m (X), E:	RIDA 1 2 DETECTION 1 325.000 m 0.000 m (Y 0.0000 N:	DATE: 05-0 SOUNDING: 1 ELEVATION: 2 EQUIPMENT: Geon AZIMUTH:) 0.0000	5-93 1.00 m ics PROTEM
		FITTIN	G ERROR:	3.854 PE	RCENT	
L #	RESI (C	STIVITY Dhm-m)	THICKNES (meters)	S ELEVAT: (mete)	ION CONDUC rs) (Siem	TANCE ens)
1 2 3	1 95 1	6.47 7.2 5.51	46.21 313.7	21.00 -25.23 -338.9	0 1 2.8 0.3	0 27
ALL P	ARAME	TERS ARE F	REE			
PARA	METER	BOUNDS FR	OM EQUIVA	LENCE ANALYSIS	S	
LAYE	R	MINIMUM	BEST	MAXIMUM		
RHO	1 2 3	7.308 538.329 13.544	16.47 957.29 15.51	1 18.232 9 3027.245 4 20.061		
THICK	1 2	18.633 297.173	0.26 1.00	9 51.768 0 329.000		
DEPTH	1 2	18.633 326.210	46.21 359.96	2 51.768 9 375.805		
CUR FREQU	RENT: ENCY:	16.75 A 7.50 H	MPS EM- z GAIN:	57 COIL ARE 3 RAMP TIN	EA: 100.00 s ME: 252.00 muS	q m. EC
No.	т ('IME ms)	em DATA	f (nV/m sqrd) SYNTHEN	DIFFER FIC (perc	ENCE ent)
1 2	0	.698 .869	8775.1 4340.6	7665.3	12.6	4
3	1	.10	2123.8	2123.0	0.0	0 375
ST. JOH	INS R	IVER WATER		SUBSUBEACE	TDEM SOUNDI SOUNI DESE	NG DATA TABLE DING 12 RET #2
PALATKA	, FLC	RIDA		DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: TABLE	93742 5.13–1

	No.	TIME (ms)	er DAT	nf (nV/m sq A SYN	Ird) ITHETIC	DIFFERE (perce	INCE ent)
	4 5 6	1.40 1.75 2.22	990.6 515.8 269.9	104 53 27	4.3 9.0 8.4	-5.42 -4.50 -3.15	
	7 8	2.79 3.42	148.9 95.5	14 5 9	9.4	-0.34 -1.42	7
	9 10 11	4.26 5.49 6.96	60.73 37.93	8 5 8 3	9.02	2.88	
	12 13	8.66 11.06	14.9	9 1 9	.5.69 9.66	-4.64	
	14 15	14.00 17.47	6.14 3.74	1 1	6.18 3.79	-0.53 -1.22	8
	16	22.23	2.24	1	2.30	-2.65	
	CURI FREQUI	RENT: 16.79 ENCY: 3.00	5 AMPS EM 0 Hz GAIN	-57 COIL : 8 RAMP	AREA: TIME:	100.00 sq 252.00 muSE	[m. C
	No.	TIME (ms)	er DAT/	nf (nV/m sg A SYN	rd) THETIC	DIFFERE (perce	INCE Int)
	17 18	0.857 1.06	4599.7 2369.9	441 237	0.3	4.11	. 8
	19	1.37	1093.8	112	5.5	-2.89	
	20 21	2.17	287.6	54 29	7.6	-2.12	
	22	2.77	162.5	15	3.1	5.76	MASKED
	23	3.50	92.4	7 9	3.14	-0.73	0
1	24	5.56	37.1	ר ו ג ו	7.07	1.98	3
ļ	26	6.98	24.49	2	4.04	1.80)
-	27	8.56	14.3	9 1	6.73	-16.24	MASKED
	PARAMI "F" II P 1 (P 2 (P 3 (T 1 -(T 2 (ETER RESOLUT: NDICATES FIXI 0.96 0.02 0.01 0.01 -0.03 (0.05 -0.01 (0.00 0.03 (P 1 P 2	ION MATRIX: ED PARAMETE 0.84 0.03 0.94 0.05 0.00 P 3 T 1	0.98 T 2			
	ST. JOHN	IS RIVER WATE	R	SUBSURFACE	- T	DEM SOUNDIN SOUNDI DESERE	G DATA TABLE NG 12 T #2
	PALATKA,	FLORIDA		DETECTION		PROJECT NO .:	93742
				INCORPORAT	ED	TABLE	5.13-1

5.14 TDEM SITE 13 - UNIVERSITY OF CENTRAL FLORIDA (UCF) SITE

5.14.1 Location Description and Geoelectrical Section

The site is located in Orange County, Florida (Figure 5.14-1) on the University of Central Florida campus. The site is located in a wooded area with no obvious sources of interference and approximately 4 miles from a previous TDEM sounding (Site 10; CEES, 1992).

The Floridan aquifer begins at an approximate depth of 115 ft below msl and is overlain by the Hawthorn Group and surficial sediments (Scott et al., 1991). The top of the Lower Floridan aquifer begins at a depth of approximately 1,000 ft below msl. The base of the Floridan aquifer approximately occurs at 2,400 ft below msl in this area (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.14-2. The interpreted geoelectrical section consists of a three-layer subsurface similar to nearby Sites 11 and 12.

5.14.2 Geological Interpretation of Geoelectrical Model

There is a sufficient electrical resistivity contrast to distinguish two geological layers above the third saltwater saturated layer. The first layer occurs at a depth of 38 m (125 ft) and not at the hydrostratigraphic contact (181 ft bls) between the Hawthorn Group and the Floridan aquifer System. The first layer has a low-resistivity value (24 ohm-m) and a thickness of 38 m (125 ft). It is considered to represent the upper portion of the Hawthorn Group and surficial sediments. The interpreted thickness and resistivity compares well with that obtained at the site 10 (CEES, 1992) sounding (26 ohm-m and 124 ft). The second layer has a high-resistivity value (534 ohm-m) and a thickness of 396 m (1,299 ft). It is considered to represent a combined but indistinguishable (geoelectrical) layer consisting of a portion of the Hawthorn Group and part of the Floridan aquifer. The third layer is considered to represent a saltwater saturated Floridan aquifer at a depth of 1,423 ft. The depth to the interpreted low resistivity (saltwater) layer at 1,423 ft is in agreement with the position of this same layer from the 1992 sounding (1,556 ft).

5.14.3 Depth to Occurrence of Salt Water

The bottom (third) layer of the geoelectrical model, with a resistivity of 5.5 ohm-m, is interpreted to represent salt water. It occurs at a depth of 1,423 ft (-1,357 ft msl). Because the resistivity of layer 2 (534 ohm-m) is interpreted to represent freshwater within the Floridan aquifer (i.e., is greater than 80 ohm-m), the interpreted depth to the 5,000 mg/L isochlor is 50 ft below the depth of the geoelectrical interface, or at 1,473 ft depth (-1,407 ft msl). The resistivity of layer 3 (5.5 ohm-m) corresponds to a chloride content of 5,695 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2.

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5.14.4 Depth of Occurrence of the 250 mg/L Isochlor

Because of the inability to segregate the Floridan aquifer from part of the overlying Hawthorn Group, the effective chloride concentration of Layer 2 cannot be calculated. A chloride concentration of less than 250 mg/L was mapped in the area of the site by Sprinkle (1981).

5.14.5 Accuracy of Measurement and Interpretation

Figure 5.14-3 is the equivalence analysis at this site, and the inversion table (Table 5.14-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about 10 m (33 ft) which is 2% of the total depth. The resistivity of this layer has a well-constrained range of from 4.7 - 6.4 ohm-m. This corresponds to a range in interpreted chloride content of from 6,690 mg/L to 4,872 mg/L, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence range of the resistivity of layer 2 is from 399 - 955 ohm-m. A corresponding chloride concentration cannot be determined because Layer 2 is in part comprised of the Hawthorn Group.

5.14.6 Summary of TDEM Sounding at UCF (Site 13)

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 1,473 ft (-1,407 ft msl) and occur within the Lower Floridan aquifer. The measured layer resistivity at this level yields a chloride concentration of 5,695 mg/L. The results of the TDEM survey are in agreement with other water quality studies conducted in the area of the site.
- The ground water within the Floridan aquifer at this site cannot be interpreted because analysis of the TDEM data does not allow part of the overlying Hawthorn Group to be distinguished from the Floridian Aquifer System.







CLI LOCAT COU PROJ LOOP S COIL SOUNDI	ENT: S ION: U NTY: C ECT: S IZE: LOC: NG COC	SJWRMD JNIVERSITY DRANGE COU SALT WATER 381.000 0.000 ORDINATES: FITTIN	OF CENTRA NTY, FLORI INTERFACE m by m (X), E: G ERROR:	L FLORIDA S DA EI DETECTION EQ 457.000 m 0.000 m (Y) 0.0000 N: 5.673 PERC	DATE: 06- SOUNDING: 1 LEVATION: QUIPMENT: Geo AZIMUTH: 0.0000 CENT	05-93 20.00 m nics PROTEM
L #	RESIS (Oł	STIVITY um-m)	THICKNESS (meters)	ELEVATIO (meters	ON CONDU	CTANCE mens)
1 2 3	23 534 5	3.60 4.0 5.49	37.80 395.6	20.00 -17.80 -413.4	1. 0.	60 740
ALL P. Para	ARAME ¹ METTER	ERS ARE F	REE	FNCF ANALVETS		
LAYE	R	MINIMUM	BEST	MAXIMUM		
RHO	1 2 3	14.370 399.364 4.728	23.605 534.065 5.494	30.255 955.093 6.384		
THICK	1 2	21.939 381.717	-2.938 1.000	49.622 409.431		
DEPTH	1 2	21.939 423.732	37.803 433.496	49.622 441.995		
CUR FREQU	RENT: ENCY:	16.00 A 7.50 H	MPS EM-5 z GAIN:	7 COIL AREA 3 RAMP TIM	A: 100.00 E: 272.00 mu	sq m. SEC
No.	Т: (1	IME NS)	emf DATA	(nV/m sqrd) SYNTHET	DIFFE IC (per	RENCE cent)
1 2 3	0 0 0	.346 .427 .550	48482.4 21861.0 8422.0	16647.8 10219.7 5312.2	65. 53. 36.	66 MASKED 25 MASKED 92 MASKED
ST. J MANA PALAT	IOHNS GEMEN IKA, Fl	RIVER WATE T DISTRICT _ORIDA	R	SUBSURFACE DETECTION INVESTIGATIONS INCORPORATED	TDEM SOUND SOUN UNIVERSITY OF (PROJECT NO.: TABLE	NG DATA TABLE DING 13 CENTRAL FLORIDA 93742 5.14–1

No.	TIME	em	f (nV/m sqrd)	DIFFERE	NCE
	(ms)	DATA	SYNTHEI	CIC (perce	ent)
4 5 6 7 8 9 10 11 12 13 14 15 16 17 18 19 20	0.698 0.869 1.10 1.40 1.75 2.22 2.79 3.42 4.26 5.49 6.96 8.66 11.06 14.00 17.47 22.23 28.10	3393.4 1489.0 667.5 302.9 166.8 100.3 67.62 50.74 38.05 27.65 19.46 14.36 10.45 7.19 5.09 3.68 2.21	2717.0 1423.6 664.1 332.8 176.1 103.3 66.31 48.99 36.46 25.88 18.87 13.89 9.71 6.80 4.68 3.08 1.97	$ \begin{array}{r} 19.93\\ 4.39\\ 0.50\\ -9.87\\ -5.60\\ -2.89\\ 1.94\\ 3.44\\ 4.17\\ 6.39\\ 3.04\\ 3.29\\ 7.00\\ 5.47\\ 7.96\\ 16.31\\ 10.74\\ \end{array} $	MASKED 4 MASKED MASKED MASKED
CURI	RENT: 16.00	AMPS EM-	57 COIL ARE	A: 100.00 sq	m.
FREQUI	ENCY: 3.00	Hz GAIN:	8 RAMP TIM	IE: 272.00 muSE	C
No.	TIME	em:	f (nV/m sqrd)	DIFFERE	NCE
	(ms)	DATA	SYNTHET	CIC (perce	nt)
21 22 23 24 25 26 27 28 29 30 31 32 33 34 PARAME "F" IN P 1 P 2 P 3 T 1 - T 2	0.857 1.06 1.37 1.74 2.17 2.77 3.50 4.37 5.56 6.98 8.56 10.64 13.70 17.40 STER RESOLUTION NDICATES FIXED 0.88 0.02 0.14 0.03 -0.07 (0) 0.13 -0.08 (0) 0.01 0.02 (0) P 1 P 2	1587.9 761.7 341.1 175.4 104.7 70.72 48.73 35.80 25.36 18.43 13.79 11.84 7.87 5.09 ON MATRIX: D PARAMETER	1487.5 751.2 357.6 180.8 109.6 68.91 49.18 36.79 26.99 20.20 15.40 11.53 8.07 5.61	$ \begin{array}{r} 6.32\\ 1.37\\ -4.82\\ -3.09\\ -4.66\\ 2.56\\ -0.92\\ -2.78\\ -6.40\\ -9.59\\ -11.69\\ 2.60\\ -2.51\\ -10.10\\ \end{array} $	MASKED MASKED 5 MASKED
ST. JO MANAG PALATK	HNS RIVER WAT EMENT DISTRICT (A, FLORIDA	ER	SUBSURFACE DETECTION INVESTIGATIONS INCORPORATED	TDEM SOUNDING SOUNDIN UNIVERSITY OF CEN PROJECT NO.: TABLE	DATA TABLE G 13 TRAL FLORIDA 93742 5.14-1

5.15 TDEM SITE 14 - RICHLAND PROPERTIES SITE

5.15.1 Location Description and Geoelectrical Section

The site is located within a pasture in Orange County, Florida (Figure 5.15-1). Power lines were present approximately three-quarters of a mile to the east and west of the site. Quality control soundings made during the survey indicate that the power lines did not have a significant affect on survey results (Figure 5.15-4). Noticeable scatter was evident in late-time data, however, data from these late-times were not used in the modeling of the geoelectric section.

The Floridan aquifer begins at an approximate depth of 135 ft below msl and is overlain by the Hawthorn Group and surficial sediments (Scott et al., 1991). The top of the Lower Floridan aquifer begins at approximately 1,050 ft below msl. The base of the Floridan aquifer occurs at approximately 2,400 ft below msl (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.15-2. The interpreted geoelectrical section consists of a three-layer subsurface.

5.15.2 Geological Interpretation of Geoelectrical Model

There is a sufficient electrical resistivity contrast to distinguish two geological layers above the third saltwater saturated layer. The first layer occurs at a depth of 39 m (128 ft) and not at the hydrostratigraphic contact (220 ft bls) between the Hawthorn Group and the Floridan aquifer System. The first layer has a low-resistivity value (16.5 ohm-m) and is considered to represent the upper portion of the Hawthorn Group and surficial sediments. The second layer has a highresistivity value (2,551 ohm-m) and a thickness of 546 m (1,791 ft). It is considered to represent a combined but indistinguishable (geoelectrical) layer consisting of a portion of the Hawthorn Group and part of the Floridan aquifer. The third layer is considered to represent a saltwater saturated Floridan aquifer at a depth of 1,919 ft.

5.15.3 Depth to Occurrence of Salt Water

The bottom (third) layer of the geoelectrical model, with a resistivity of 2.8 ohm-m, is interpreted to represent salt water. It occurs at a depth of 1,919 ft (-1,834 ft msl). Because the resistivity of layer 2 (2,551 ohm-m) is interpreted to represent fresh water within the Floridan aquifer (i.e., is greater than 80 ohm-m), the interpreted depth to the 5,000 mg/L isochlor is 50 ft below the depth of the geoelectrical interface, or at 1,969 ft depth (-1,884 ft msl). The resistivity of layer 3 (2.8 ohm-m) corresponds to a chloride content of 11,334 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2.

5.15.4 Depth of Occurrence of the 250 mg/L Isochlor

Because of the inability to segregate the Floridan aquifer from part of the overlying Hawthorn Group, the effective chloride concentration of Layer 2 cannot be calculated. A chloride concentration of less then 250 mg/L was mapped in the area of the site by Sprinkle (1981).

5.15.5 Accuracy of Measurement and Interpretation

Figure 5.15-3 is the equivalence analysis at this site, and the inversion table (Table 5.15-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about \pm 33 m (108 ft), which is 6% of the total depth. The resistivity of this layer has a poorly-constrained range of from 2.3 - 7.0 ohm-m. This corresponds to a range in interpreted chloride content of from 13,831 mg/L to 4,442 mg/L, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence range of the resistivity of layer 2 is from 1,007 - 7,112 ohm-m. A corresponding chloride concentration cannot be determined because Layer 2 is in part comprised of the Hawthorn Group.

5.15.6 Summary of TDEM Sounding at Richland Properties (Site 14)

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 1,969 ft (-1,884 ft msl) and occurs within the Lower Floridan aquifer. The measured layer resistivity at this level suggests a chloride concentration of 11,334 mg/L.
- The ground water within the Floridan aquifer at this site cannot be interpreted because analysis of the TDEM data does not allow part of the overlying Hawthorn Group to be distinguished from the Floridian Aquifer System.









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		FITTIN	G ERROR:	4.431 PERC	ENT	
L #	RESI (C	STIVITY hm-m)	THICKNESS (meters)	ELEVATIO (meters	N CONDUCTANCE) (Siemens)	
1 2 3	1 255	16.46 51.4 2.84	39.37 545.8	26.00 -13.37 -559.2	2.39 0.213	
ALL P.	ARAMI	TERS ARE F	REE			
PARA	Metef	R BOUNDS FR	OM EQUIVAL	ENCE ANALYSIS		
LAYE	R	MINIMUM	BEST	MAXIMUM		
RHO	1 2 3	8.037 1007.085 2.269	16.466 2551.418 2.840	19.484 7112.259 6.975		
THICK	1 2	18.155 522.762	-0.151 1.000	47.217 589.838		
DEPTH	1 2	18.155 562.394	39.375 585.271	47.217 627.732		
CUR FREQU	RENT : ENCY :	15.00 A 30.00 H	MPS EM-5 z GAIN:	7 COIL AREA 2 RAMP TIME	: 100.00 sq m. : 262.00 muSEC	
No.	5 (CIME (ms)	emf DATA	(nV/m sqrd) SYNTHETI	DIFFERENCE C (percent)	
1 2	().438	29918.4 14757.2	18902.3 11088.9	36.82 MAS	KED
3	ĺ	0.702	6967.7	6248.7	10.31 MAS	KED
ST. JO	HNS	RIVER WATER	२	SDII	TDEM SOUNDING D SOUNDING RICHLAND PROPI	ATA TABLE 14 ERTIES
MANAG PALATK	EMEN (A, Fl	ORIDA		DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: TABLE	93742 5.15-1

s. Ż

No.	TIME (ms)	emf () DATA	nV/m sqrd) SYNTHETIC	DIFFERENC (percent	E)
4 5	0.858 1.06	3670.4 1812.2	3566.9 1841.5	2.82 -1.61	
ь 7 8	1.37 1.74 2.17	783.7 336.2 154 7	828.5 350.9	-5.71 -4.35	
9 10	2.77 3.50	70.07	65.64 28.63	-5.35 6.33 6.10	
11 12 12	4.37 5.56	15.59 8.39	15.72 8.11	-0.862 3.39	
15	7.03	5.29	5.42	-2.45	
PARAME "F" IN	TER RESOLUTIO DICATES FIXE	ON MATRIX: D PARAMETER			
P 1 0 P 2 0	.98 .00 0.01				
P 3 0 T 1 -0	.01 - 0.03 0 .02 - 0.02 0	.24 .03 0.98			
T 2 0	.00 0.01 -0. P1 P2	.07 0.00 0.98 P3 T1 T	3 2		· ·
			TTT TD	EM SOUNDING DA	TA TABLE
T. JOHNS	S RIVER WATER			SOUNDING RICHLAND PROPE	I 4 RTIES
ALATKA,	FLORIDA	DETE	ECTION PI	ROJECT NO .:	93742
		INCO	RPORATED	ABLE	5.15-1

5.16 TDEM SITE 15 - NEW SMYRNA BEACH SITE

5.16.1 Location Description and Geoelectrical Section

The site is situated in eastern Volusia County, Florida (Figure 5.16-1). The site was located in a grassy area within the New Smyrna Beach Municipal Airport. A buried steel pipe was present approximately 100 ft east of the Tx loop. QA soundings were performed 100 ft north, south, east and 80 ft west of the initial Rx coil location. Results from the QA soundings indicated that apparent resistivity values were not affected by any interference sources (Figure 5.16-4).

Together, Sites 1, 15, 16, 17, 18, 19 and 20 each lie within the area where the Hawthorn Group is either thin or absent. The top of the Floridan aquifer begins at approximately 80 ft below msl (Scott et al., 1991). The top of the Lower Floridan aquifer is approximately 750 ft below msl (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.16-2. The interpreted geoelectrical section consists of a two-layer subsurface.

5.16.2 Geological Interpretation of Geoelectrical Model

There is insufficient electrical resistivity contrast between the surficial aquifer system layer, the Hawthorn Group and the underlying Floridan aquifer to distinguish the three. Fixing the thickness of the upper layer does not resolve this dilemma; therefore, it can be interpreted that there exists a two-layer geoelectrical section with a relatively thick (126 m = 413 ft) surface layer of intermediate to low (9.2 ohm-m) resistivity overlying a very low resistivity layer (2.3 ohm-m). It can be interpreted that the surficial aquifer system, the Hawthorn Group and part of the Floridan aquifer exist as a combined but indistinguishable (geoelectrical) layer, overlying a saltwater saturated Floridan aquifer at a depth of 413 ft.

5.16.3 Depth to Occurrence of Salt Water

The bottom (second) layer of the geoelectrical model, with a resistivity of 2.3 ohm-m, is interpreted to represent salt water. It occurs at a depth of 413 ft (-407 ft msl). Because the resistivity of layer 1 (9.2 ohm-m) is less than 80 ohm-m, the interpreted depth to the 5,000 mg/L isochlor is taken at the depth of the geoelectrical interface, or at 413 ft depth (-407 ft msl). The resistivity of layer 2 (2.3 ohm-m) corresponds to a chloride content of 13,831 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2.

5.16.4 Depth of Occurrence of the 250 mg/L Isochlor

Because of the inability to segregate the Floridan aquifer from the overlying Hawthorn Group and surficial aquifer system, the effective chloride concentration of layer 1 cannot be calculated.

5.16.5 Accuracy of Measurement and Interpretation

Figure 5.16-3 is the equivalence analysis for this site, and the inversion table (Table 5.16-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about $\pm 5 \text{ m}$ (15 ft), which is 4% of the total depth. The resistivity of this layer has an equivalence range of from 2.1 - 2.5 ohm-m. This corresponds to a range in interpreted chloride content of from 15,163 mg/L to 12,712 mg/L, again subject to the same assumptions of porosity and validity of equation (4). The chloride-to-sulfate ratio at the site is 5:1 (Table 5.1-4). Accordingly, equation (4) is valid.

The equivalence range of the resistivity of layer 1 is from 9.1 - 9.4 ohm-m. A corresponding chloride concentration cannot be determined because layer 1 is in part comprised of the Hawthorn Group and surficial sediments. Accordingly, equation (4) may not be valid. Rutledge (1985) mapped a chloride concentration of greater than 250 mg/L in the area of the site.

5.16.6 Summary of TDEM Sounding at New Smyrna Beach (Site 15)

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 413 ft (-407 ft msl) and occur within the Upper Floridan aquifer. The measured layer resistivity at this level yields a chloride concentration of 13,831 mg/L.
- The groundwater within the Floridan aquifer at this site cannot be determined because analysis of the TDEM data does not allow for the surficial aquifer system and the Hawthorn Group to be distinguished from the Floridan Aquifer System.









CLI LOCAT COU PROJ LOOP S COIL SOUNDI	ENT: S ION: NI NTY: V ECT: S IZE: LOC: NG COO	JWRMD EW SMYRNA OLUSIA CO ALT WATER 304.000 0.000 RDINATES: FITTIN	BEACH UNTY, FLO INTERFAC m by m (X), E: G ERROR:	RIDA E DETECTION 274.000 m 0.000 m (Y 0.0000 N: 2.256 PE	DATE: SOUNDING: ELEVATION: EQUIPMENT: AZIMUTH:) 0.00 RCENT	08-05-93 1 2.00 Geonics 1	m PROTEM
L #	RESIS (ohi	TIVITY m-m)	THICKNES (meters)	S ELEVAT (mete	ION CO rs)	ONDUCTANC) (Siemens)	B
1 2	9 2	.24 .25	126.0	2.0 -124.0	0	13.63	
ALL P	ARAMET	ERS ARE F	REE				
PARA	METER	BOUNDS FR	OM EQUIVA	LENCE ANALYSI	S		
LAYE	R	MINIMUM	BESI	MAXIMUM			
RHO	1 2	9.068 2.094	9.24 2.25	4 9.448 9 2.451			
THICK	1	121.161	1.00	0 130.360			
DEPTH	1	121.161	126.05	9 130.360			
CUR FREQU	RENT: ENCY:	16.00 A 7.50 H	MPS EM- Iz GAIN:	57 COIL AR 2 RAMP TI	EA: 100 ME: 192.0	.00 sq m. 0 muSEC	
No.	TI (m	ME S)	en DATA	nf (nV/m sqrd) SYNTHE	D	IFFERENCE (percent)	
1 2 3 4 5 6 7	0. 0. 0. 0. 1. 1.	346 427 550 698 869 10 40	65738.9 46489.1 29202.6 18714.0 12204.0 7981.4 5017.3	49556.9 38962.7 27569.0 18859.5 12763.9 7999.3 4998.2		24.61 MA 16.18 MA 5.59 -0.777 -4.58 -0.224 0.380	SKED SKED
ST. JOH	INS RIV	ER WATER			TDEM S NEW	OUNDING D SOUNDING SMYRNA E	ATA TABLE 15 BEACH
PALATKA	, FLOR	IDA		DETECTION INVESTIGATIONS INCORPORATED	PROJECT TABLE	NO.:	93742 5.16-1

فجري

No.	TIME (ms)	em DATA	f (nV/m sqrd) SYNTHE	DIFFEF TIC (perc	RENCE ent)
8 9	1.75	3278.9 2079.3	3214.6 2020.0	1.9 2.8)6 35
10	2.79	1321.5	1318.9	0.1	98
12	4.26	609.0	615.0	-0.9	95
13	5.49	388.6	392.3	-0.9	43
14	6.96	252.7	254.9	-0.9	00
15	11.06	109.2	10.1	-0.8	105
17	14.00	68.37	66.2	7 3.0	
18	17.47	42.58	41.7	5 1.9	4
20	22.23 28.10	25.09 13.99	24.5 14.3	9 1.9 7 -2.7	18 13
CUR FREQU	RENT: 16.00 ENCY: 3.00	AMPS EM- Hz GAIN:	57 COIL AR 8 RAMP TI	EA: 100.00 s ME: 192.00 muS	g m. EC
No.	TIME (ms)	em DATA	f (nV/m sqrd) SYNTHE	DIFFEF FIC (perc	ENCE ent)
21	0.857	12544.9	13109.4	-4.4	9
22	1.06	8496.2	8642.2	-1.7	1
23	1.74	3348.4	3262.6	-0.0	1250
25	2.17	2161.5	2120.9	1.8	37
26	2.77	1362.0	1347.0	1.1	.0
27	3.50	881.2	883.9	-0.3	10
20	4.37	386.4	597.4 397 5	-1.5	15 17
30	6.98	261.7	261.9	-0.0	467
31	8.56	177.6	181.4	-2.0	9
32	10.64	119.1	121.4	-1.9	3
33	13.70	68.9/ 38.53	75.0	5 -8.8 5 _21 g	O MASKED
PARAM "F" I P 1 P 2 T 1	ETER RESOLUTI NDICATES FIXE 1.00 0.00 0.96 0.00 0.01 0 P 1 P 2	ON MATRIX: D PARAMETER .99 T 1	2		
	<u></u>		SUI	TDEM SOUNDIN	IG DATA TABLE
ST. JOHN	NS RIVER WATE	R	SDI	SOUND	ING 15 NA BEACH
MANAGEN	IENT DISTRICT	1	SUBSURFACE		
PALATKA,	FLORIDA		INVESTIGATIONS	PROJECT NO .:	93742
			INCORPORATED	TABLE	5.16-1

5.17 TDEM SITE 16 - LAKE ASHBY SITE

5.17.1 Location Description and Geoelectrical Section

The site is located in south-central Volusia County, Florida (Figure 5.17-1) within an operating sod farm. No obvious signs of interference sources were present. Together, Sites 1, 15, 16, 17, 18, 19 and 20 each lie within the area where the Hawthorn Group is either thin or absent. The top of the Floridan aquifer begins at approximately 85 ft below msl (Scott et al., 1991). The top of the Lower Floridan aquifer is approximately 700 ft below msl (Miller, 1986). Results from QC soundings performed 100 ft north, south, east and west of the initial Rx coil position indicated that no sources of interference were present (Figure 5.17-4).

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The resistivity sounding data and best-fit model inversion are presented on Figure 5.17-2. The interpreted geoelectrical section consists of a two-layer subsurface.

5.17.2 Geological Interpretation of Geoelectrical Model

There is insufficient electrical resistivity contrast between the surficial aquifer system, the Hawthorn Group and the underlying Floridan aquifer to distinguish the three. Fixing the thickness of the upper layer does not resolve this dilemma; therefore it can be interpreted that there exists a two-layer geoelectrical section with a relatively thick (234.4 m = 769 ft) surface layer of intermediate resistivity (41.8 ohm-m) overlying a low resistivity layer (4.6 ohm-m). It can be interpreted that the surficial aquifer system, the Hawthorn Group and part of the Floridan aquifer system exist as a combined but indistinguishable (geoelectrical) layer, overlying a saltwater saturated Floridan aquifer at a depth of 769 ft bls.

5.17.3 Depth to Occurrence of Salt Water

The bottom (second) layer of the geoelectrical model, with a resistivity of 4.6 ohm-m, is interpreted to represent salt water. It occurs at a depth of 769 ft (-743 ft msl). Because the resistivity of layer 1 (41.8 ohm-m) is less than 80 ohm-m, the interpreted depth to the 5,000 mg/L isochlor is equal to the depth of the geoelectrical interface, or at 769 ft depth (-743 ft msl). The resistivity of layer 2 (4.6 ohm-m) corresponds to a chloride content of 6,839 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2.

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5.17.4 Depth of Occurrence of the 250 mg/L Isochlor

Because of the inability to segregate the Floridian aquifer from the overlying surficial aquifer system and the Hawthorn Group, the effective chloride concentration of layer 1 cannot be calculated. Rutledge (1985) mapped a chloride concentration of less than 250 mg/L in the area of the site.

5.17.5 Accuracy of Measurement and Interpretation

Figure 5.17-3 is the equivalence analysis at this site, and the inversion table (Table 5.17-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about $\pm 8 \text{ m}$ (26 ft), which is 3% of the total depth. The resistivity of this layer has an equivalence range of from 3.8 - 5.6 ohm-m. This corresponds to a range in interpreted chloride content of from 8,311 mg/L to 5,590 mg/L, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence range of the resistivity of layer 2 is constrained from 40.4 - 43.5 ohm-m. A corresponding chloride concentration cannot be determined because layer 1 is in part comprised of the Hawthorn Group and surficial sediments. Accordingly, equation (4) may not be valid.

5.17.6 Summary of TDEM Sounding at Lake Ashby (Site 16)

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• The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 769 ft (-743 ft msl) and occur in the Lower Floridan aquifer. The measured layer resistivity at this level yields a chloride concentration of 6,839 mg/L.

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• The ground water within the Floridan aquifer at this site cannot be interpreted because analysis of the TDEM data does not allow the surficial aquifer system and the Hawthorn Group to be distinguished from the Floridan Aquifer System.








CLI LOCAT COU PROJ LOOP S COIL SOUNDI	ENT: S. ION: L. NTY: V ECT: S. IZE: LOC: NG COO	JWRMD AKE ASHBY OLUSIA CO ALTWATER 411.000 0.000 RDINATES:	UNTY, FLO INTERFAC) m by m (X), E:	ORIDA E DETECTION 198.000 m 0.000 m 0.0000 N	SOU ELEY EQUI AI (Y)	DATE: 09 UNDING: 1 VATION: IPMENT: Ge ZIMUTH: 0.0000	9-05-93 8.00 m eonics PRO	TEM
		FITTIN	G ERROR:	4.803	PERCEI	NТ		
L #	RESIS (ohi	riviry m-m)	THICKNES (meters)	SS ELEV) (me	ATION ters)	CONI (Si	OUCTANCE Lemens)	
1 2	41 4	.84 .61	234.4	8 -226	.00 .4	5	5.60	
ALL P.	ARAMETI	ERS ARE F	REE					
PARA	METER I	BOUNDS FR	OM EQUIVA	ALENCE ANALY	SIS			
LAYE	R I	MINIMUM	BEST	r MAXIMU	M			
RHO	1 2	40.366 3.829	41.8 4 4.6 1	43.49 13 5.56	7 3			
THICK	1	226.224	1.00	0 242.47	6			
DEPTH	1	226.224	234.47	76 242.47	6			
CUR FREQU	RENT: ENCY:	16.10 A 7.50 H	MPS EM- z GAIN:	-57 COIL : 4 RAMP	AREA: TIME:	100.00 192.00 n) sq m. NUSEC	
No.	TII (MS	ME 5)	en DATA	nf (nV/m sqr A SYNT	d) HETIC	DIFE (pe	ERENCE ercent)	
1 2 3 4 5 6 7	0.9 0.6 1.7 1.7 2.7	550 598 369 10 40 75 22	6083.0 3141.9 1779.4 1066.6 654.7 438.8 295.4	5263 3109 1885 1096 665 432 285	.9 .6 .5 .8 .9 .2	13 -5 -2 -1 1 3	.46 .02 .97 .80 .69 .33 .42	
				. v [*]				
ST. JOHI MANAGEN	NS RIVE	R WATER			-	TDEM SOUL SOULAK	NDING DATA JNDING 16 E ASHBY	TABLE
PALATKA	, FLORI	DA		DETECTION INVESTIGATION INCORPORATE	NS ED	PROJECT NO. TABLE	: 93 5.1	742 7-1

No.	TIME (ms)	DZ	emf (ATA	nV/m sqra SYNTI	d) HETIC	DIFFERE) (perce	NCE nt)
8 9 10 11 12 13 14 15	2.79 3.42 4.26 5.49 6.96 8.66 11.06 14.00 17.47	201 146 102 67 44 29 20 13	.9 .1 .17 .18 .99 .56 .87	196 142 101 67 46 32 20 13	.2 .3 .4 .88 .27 .00 .80 .45	2.82 2.61 1.25 -1.04 -4.72 -6.70 -1.16 3.01	
17	22.23	6	.24	8 5	. 73	14.62	MASKED
CUR FREQU	RENT: 16 ENCY: 3	.10 AMPS H .00 Hz GAI	SM-57 [N: 8	COIL A RAMP 1	AREA: FIME:	100.00 sq 192.00 muSE(m. C
No.	TIME (ms)	Dž	emf (ATA	nV/m sqra SYNTI	d) HETIC	DIFFEREI (percei	NCE nt)
18 19 20 21 22 23 24 25	0.857 1.06 1.37 1.74 2.17 2.77 3.50 4.37	1838 1147 692 446 302 211 142 98	4 6 6 6 1 0 03	1949 1196 701 440 298 201 139 99	.0 .2 .6 .2 .9 .0 .7 .78	-6.01 -4.23 -1.35 1.41 1.21 4.81 1.62 -1.78	
26 27	5.56 6.98	65. 45.	54 29	68. 48.	.65 .01	-4.74 -6.00	
PARAM "F" I P 1 P 2 - T 1	ETER RESOL NDICATES F 1.00 0.01 0.93 0.00 0.01 P 1 P	UTION MATRIX IXED PARAMET 1.00 2 T 1					ρατά ταβι έ
ST. JOHN MANAGEM	S RIVER WA	TER T				DEM SOUNDING SOUNDING LAKE ASH	DATA TABLE 16 BY
PALATKA,	FLORIDA		INVE	ESTIGATIONS	F	ROJECT NO.:	93742 5.17-1

5.18 TDEM SITE 17 - LAKE HELEN SITE

5.18.1 Location Description and Geoelectrical Section

The site is situated in south-central Volusia County, Florida (Figure 5.18-1) and was located within an abandoned orange grove. No obvious signs of interference sources were present. Together, Site 1, 15, 16, 17, 18, 19 and 20 each lie within the area where the Hawthorn Group is thin to absent. The top of the Floridan aquifer begins at approximately 55 ft below msl (Scott et al., 1991). The top of the Lower Floridan aquifer is approximately 425 ft below msl (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.18-2. The interpreted geoelectrical section consists of a three-layer subsurface.

5.18.2 Geological Interpretation of Geoelectrical Model

There is sufficient electrical resistivity contrast to distinguish two geoelectric layers above the third, saltwater saturated layer. The depth to the contact between the first and second layers occurs at 66 m (217 ft) and not at the hydrostratigraphic contact (130 ft bls) between the Hawthorn Group and the Floridan Aquifer System. The upper layer has an intermediate value of resistivity (40 ohm-m) and a modeled thickness of 66 m (217 ft). This overlies a 258 m (846 ft) thick layer of high resistivity (497 ohm-m). The base layer in the geoelectrical model is a very low resistivity layer (2 ohm-m) situated at a depth of 324 m (1,063 ft). It can be interpreted that these layers represent the combined Hawthorn Group and surficial aquifer system and upper portion of the Upper Floridan aquifer (layer 1), overlying the Floridan aquifer (layer 2) at a depth of 217 ft. Layer 3 is the salt water portion of the Floridan aquifer; this interface occurs at a depth of 1,064 ft.

5.18.3 Depth to Occurrence of Salt Water

The bottom (third) layer of the geoelectrical model, with a resistivity of 2 ohm-m, is interpreted to represent salt water. It occurs at a depth of 1,064 ft (-989 ft msl). Because the resistivity of layer 2 (497 ohm-m) is interpreted to represent freshwater within the Floridan aquifer (i.e., is greater than 80 ohm-m), the interpreted depth to the 5,000 mg/L isochlor is taken at 50 ft below the depth of the geoelectrical interface, or at 1,114 ft depth (-1,039 ft msl). The resistivity of layer 3 (2 ohm-m) corresponds to a chloride content of greater than 10,000 mg/L (15,929 mg/L) assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2.

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5.18.4 Depth of Occurrence of the 250 mg/L Isochlor

The resistivity of layer 2, 497 ohm-m, corresponds to a chloride content of less than 50 mg/L, assuming a 25% porosity and the validity and applicability of equation (4) of Section 4.2. Using the criteria established in Section 4.3 (this is a Class A type resistivity distribution), the position of the 250 mg/L isochlor is placed 50 ft above the depth to the low resistivity interface, or at a depth of 1,014 ft (-939 ft msl).

5.18.5 Accuracy of Measurement and Interpretation

Figure 5.18-3 is the equivalence analysis at this site, and the inversion table (Table 5.18-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about ± 1.5 m (5 ft), which is less than 1% of the total depth. The resistivity of this layer has an equivalence range of from 1.8 - 2.3 ohm-m. This corresponds to a range in interpreted chloride content of from 17,716 mg/L to 13,831 mg/L, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence range of the resistivity of layer 2 is poorly constrained from 321 - 994 ohm-m which corresponds to chloride concentrations which are all less than 50 mg/L (assuming 25% porosity). The chloride-to-sulfate ratio at the site (Table 5.1-4) is 1:1. Accordingly, equation (4) may not be valid. Rutledge (1985) mapped a chloride concentration of less than 250 mg/L in the area of the site.

The estimated depth to the 250 mg/L isochlor from the TDEM study (1,014 ft) does not agree with the estimated depth to the freshwater/saltwater interface of approximately 1,400 ft bls determined by Ghyben-Herzberg analysis (Rutledge, 1982). Rutledge (1982) defines the depth to the freshwater/seawater interface as forty times the groundwater table elevation above msl. The freshwater/saltwater interface is assumed to occur at the transition to unmixed seawater with a chloride concentration of 19,000 mg/L. The difference between the estimated depth of the 250 mg/L isochlor (1,014 ft bls -- from TDEM) and the estimated depth to the 19,000 mg/L isochlor (1,400 ft -- from Ghyben-Herzberg analysis) may indicate the presence of a thick transition zone from freshwater to saltwater in the area of the site.

5.18.6 Summary of TDEM Sounding at Lake Helen (Site 17)

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 1,114 ft (-1,039 ft msl) and occurred within the Lower Floridan aquifer. The measured layer resistivity at this level yields a chloride concentration of greater than 10,000 mg/L.
- The 250 mg/L isochlor is interpreted to occur within the Lower Floridan aquifer at this site at a depth of 1,014 ft (-939 ft msl). The average chloride concentration within the fresh Floridan aquifer section is less than 50 mg/L. The water quality results from the TDEM survey are in agreement with other water quality studies in the area of the site.
- The estimated depth to the 250 mg/L isochlor is not in agreement with the estimated depth from Rutledge (1982) who estimated a depth of approximately 1,400 ft bls to the freshwater/saltwater interface.







CLI LOCAT COU PROJ LOOP S COIL SOUNDI	ENT: (ION:) NTY: Y ECT: (IZE: LOC: NG COO	SJWRMD LAKE HELEN VOLUSIA CO SALTWATER 305.000 0.000 ORDINATES: FITTIN	UNTY, FLOR INTERFACE m by m (X), E: G ERROR:	5 IDA EI DETECTION EC 305.000 m 0.000 m (Y) 0.0000 N: 2.838 PERC	DATE: 11-05-93 SOUNDING: 1 DEVATION: 23.00 m QUIPMENT: Geonics PROTEM AZIMUTH: 0.0000
L #	RESIS (ol	STIVITY nm-m)	THICKNESS (meters)	ELEVATIO (meters	ON CONDUCTANCE (Siemens)
1 2 3	39 490	9.82 5.9 2.00	66.35 258.0	23.00 -43.35 -301.4	1.66 0.519
ALL P	ARAME	TERS ARE F	REE		
PARA	METER	BOUNDS FR	OM EQUIVAL	ENCE ANALYSIS	
LAYE	R	MINIMUM	BEST	MAXIMUM	
RHO	1 2 3	32.748 320.892 1.773	39.827 496.934 2.010	47.230 993.659 2.265	
THICK	1 2	50.494 242.284	-3.785 1.000	84.265 275.388	
DEPTH	1 2	50.494 323.501	66.352 324.427	84.265 326.917	
CUR FREQU	RENT: ENCY:	16.50 A 7.50 H	MPS EM-5 z GAIN:	7 COIL AREA 6 RAMP TIME	100.00 sq m. 212.00 muSEC
No.	Т: (1	IME ns)	emf DATA	(nV/m sqrd) SYNTHET]	DIFFERENCE C (percent)
1 2 3	0 0 0	.346 .427 .550	26391.5 10965.4 4083.6	9986.1 5784.4 2837.7	62.16 MASKED 47.24 MASKED 30.51 MASKED
ST. J		RIVER WATE	R	SUBSUBEACE	TDEM SOUNDING DATA TABLE SOUNDING 17 LAKE HELEN
PALA	TKA, FL	_ORIDA		DETECTION INVESTIGATIONS	PROJECT NO.: 93742 TABLE 5.18-1

No.	TIME (ms)	em DATA	f (nV/m sqrd) SYNTHEI	DIFFER	ENCE ent)
4 5 6 7 8 9 10 11 12 13 14 15 16 17 18 19 20	0.698 0.869 1.10 1.40 1.75 2.22 2.79 3.42 4.26 5.49 6.96 8.66 11.06 14.00 17.47 22.23 28.10	$1601.4 \\704.7 \\348.8 \\191.1 \\127.4 \\90.47 \\66.72 \\52.68 \\40.88 \\30.09 \\22.39 \\16.69 \\12.04 \\8.02 \\5.39 \\3.42 \\2.02 \\$	1406.1 715.8 356.6 195.3 126.3 88.18 66.47 53.23 41.54 31.41 23.66 18.06 13.06 9.35 6.68 4.51 2.99	$ \begin{array}{c} 12.1\\ -1.5\\ -2.2\\ -2.1\\ 0.8\\ 2.5\\ 0.3\\ -1.0\\ -1.6\\ -4.4\\ -5.6\\ -8.2\\ -8.4\\ -16.5\\ -23.8\\ -31.8\\ -47.7\\ \end{array} $	9 MASKED 6 2 5 92 2 75 4 0 0 MASKED 9 MASKED 1 MASKED 1 MASKED 5 MASKED 2 MASKED 2 MASKED 4 MASKED
CUR FREQU	RENT: 16.50 ENCY: 3.00	AMPS EM- Hz GAIN:	57 COIL ARE 8 RAMP TIM	:A: 100.00 so IE: 212.00 muS	q m. EC
No.	TIME (ms)	em DATA	f (nV/m sqrd) SYNTHEI	DIFFER	ENCE ent)
21 22 23 24 25 26 27 28 29 30 31 32 33 34 35 PARAM "F" I P 1 P 2 P 3 T 1 - T 2	0.857 1.06 1.37 1.74 2.17 2.77 3.50 4.37 5.56 6.98 8.56 10.64 13.70 17.40 21.70 ETER RESOLUTION NDICATES FIXEN 0.85 0.00 0.02 0.05 -0.03 0 0.20 -0.09 0 0.05 0.03 -0 P 1 P 2	759.5 398.1 213.6 136.2 94.96 70.16 53.54 41.04 31.42 26.24 21.74 15.52 11.99 8.09 5.92 ON MATRIX: D PARAMETER .80 .08 0.72 .02 0.07 P 3 T 1	749.9 398.6 209.8 130.3 93.95 69.78 54.55 42.96 33.38 25.86 20.46 15.71 11.37 8.21 5.97	$ \begin{array}{c} 1.2\\ -0.1\\ 1.7\\ 4.2\\ 1.0\\ 0.5\\ -1.8\\ -4.6\\ -6.2\\ 1.4\\ 5.8\\ -1.2\\ 5.1\\ -1.4\\ -0.8\\ \end{array} $	5 34 8 7 6 43 8 6 2 2 6 4 7 7 40
ST. JO MANAG PALATI	DHNS RIVER WATI GEMENT DISTRICT KA, FLORIDA	ER	SUBSURFACE DETECTION INVESTIGATIONS INCORPORATED	TDEM SOUNDING SOUNDIN LAKE HI PROJECT NO.: TABLE	G DATA TABLE G 17 ELEN 93742 5.18-1

5.19 TDEM SITE 18 - DELTONA SITE

5.19.1 Location Description and Geoelectrical Section

The site is situated in southwestern Volusia County, Florida (Figure 5.19-1) and is located within a wooded area. A buried pipeline was present approximately 100 ft north of the Tx loop. QA soundings were collected 100 ft to the south and north of the initial Rx coil position. Results of the QA soundings indicated that the data was unaffected by the pipeline (Figure 5.19-4). Together, Sites 1, 15, 16, 17, 18, 19 and 20 each lie within the area where the Hawthorn Group is thin to absent. The top of the Floridan aquifer begins at approximately 85 ft below msl (Scott et al., 1991). The top of the Lower Floridan aquifer is approximately 700 ft below msl (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.19-2. The interpreted geoelectrical section consists of a two-layer subsurface.

5.19.2 Geological Interpretation of Geoelectrical Model

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There is insufficient electrical resistivity contrast between the surficial aquifer system, the Hawthorn Group and the underlying Floridan aquifer to distinguish the three. Fixing the thickness of the upper layer does not resolve this dilemma; therefore, it can be interpreted that there exists a two-layer geoelectrical section with a relatively thick (300.5 m = 986 ft) surface layer of intermediate (68.4 ohm-m) resistivity overlying a very low resistivity layer (1.4 ohm-m). It can be interpreted that the surficial aquifer system, the Hawthorn Group and part of the Floridan aquifer exist as a combined but indistinguishable (geoelectrical) layer, overlying a saltwater saturated Floridan aquifer at a depth of 986 ft.

5.19.3 Depth to Occurrence of Salt Water

The bottom (second) layer of the geoelectrical model, with a resistivity of 1.4 ohm-m, is interpreted to represent salt water. It occurs at a depth of 986 ft (-940 ft msl). Because the resistivity of layer 1 (68.4 ohm-m) is less than 80 ohm-m, the interpreted depth to the 5,000 mg/L isochlor is taken at the depth of the geoelectrical interface, or at 986 ft depth (-940 ft msl). The resistivity of layer 2 (1.4 ohm-m) corresponds to a chloride content of 22,821 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2.

5.19.4 Depth of Occurrence of the 250 mg/L Isochlor

Because of the inability to segregate the Floridan aquifer from the overlying surficial aquifer system and Hawthorn Group, the effective chloride concentration of layer 1 cannot be calculated. Rutledge (1985) mapped a chloride concentration of less than 250 mg/L in the area of the site.

5.19.5 Accuracy of Measurement and Interpretation

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Figure 5.19-3 is the equivalence analysis at this site, and the inversion table (Table 5.19-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about $\pm 4 \text{ m}$ (13 ft), which is 1% of the total depth. The resistivity of this layer has an equivalence range of from 1.1 - 1.7 ohm-m. This corresponds to a range in interpreted chloride content of from 29,087 mg/L to 18,767 mg/L, again subject to the same assumptions of porosity and validity of equation (4). The estimated depth of the saltwater layer (986 ft bls) is in good agreement with the estimated depth of the freshwater/saltwater interface (approximately 1000 ft bls) based on Ghyben-Herzberg analysis (Rutledge, 1982). Rutledge (1982) defines the depth to the freshwater/saltwater interface as forty times the groundwater elevation above msl. The freshwater/saltwater interface is assumed to occur at the transition to unmixed seawater with a chloride concentration of 19,000 mg/L. The difference between the estimated depth to the 5,000 mg/L isochlor (986 ft bls) - from TDEM) and the estimated depth to the 19,000 mg/L isochlor

(1,000 ft bls -- from Ghyben-Herzberg analysis) may indicate the presence of a thin transition zone from fresh to salt water in the area of the site.

The equivalence range of the resistivity of layer 1 is from 65.2 - 71.6 ohm-m. A corresponding chloride concentration cannot be determined because layer 1 is in part comprised of the Hawthorn group and surficial sediments. Accordingly, equation (4) may not be valid.

5.19.6 Summary of TDEM Sounding at Deltona (Site 18)

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 986 ft (-940 ft msl) and occur within the Lower Floridan aquifer. The results from the TDEM studies are in agreement with the results from another study in the area (Rutledge, 1982). The measured layer resistivity at this level yields a chloride concentration of greater than 10,000 mg/L.
- The quality of groundwater within the Floridan aquifer at this site cannot be interpreted because analysis of the TDEM data does not allow the surficial aquifer system and the Hawthorn Group to be distinguished from the Floridan Aquifer System.

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CLII LOCAT COU PROJI LOOP S COIL SOUNDI	ENT: S ION: DE NTY: V(ECT: S IZE: LOC: NG COOE	JWRMD SLTONA DLUSIA CO ALTWATER 274.000 0.000 RDINATES:	OUNTY, FLO INTERFACE) m by) m (X), : E:	RIDA E DETECTION E 152.000 m 0.000 m (Y) 0.0000 N:	DATE: 10-05-93 SOUNDING: 1 ELEVATION: 14.00 EQUIPMENT: Geonics AZIMUTH: 0.0000) m PROTEM
		FITTIN	IG ERROR:	4.364 PER	RCENT	
L #	RESIST (ohr	TIVITY n-m)	THICKNES (meters)	S ELEVATI (meter	CON CONDUCTANC (Siemens)	CE
1 2	68 1	.36 .39	300.5	14.00 -286.5	4.39	
ALL P.	ARAMETI	ERS ARE F	REE?			
PARA	METER I	BOUNDS FF	ROM EQUIVA	LENCE ANALYSIS	3	
LAYE	R 1	MINIMUM	BEST	MAXIMUM		
RHO	1 2	65.162 1.087	68.36 1.39	8 71.573 5 1.709		
THICK	1	296.325	1.00	0 304.465		
DEPTH	1	296.325	300.55	2 304.465		
CUR FREQU	RENT : ENCY :	21.30 A 7.50 H	MPS EM- Iz GAIN:	57 COIL ARE 5 RAMP TIM	2A: 100.00 sq m. 1E: 192.00 muSEC	
No.	TII (ms	1E 3)	em DATA	f (nV/m sqrd) SYNTHET	DIFFERENCE CIC (percent)	5
1 2	1.1	LO 40	281.6	299.8	-6.45	
- 3 4	1.7	75	116.2	108.6	6.52	
5	2.7	79	55.48	54.34	2.05	
7	4.2	26	42.03 31.93	42.43 33.16	-0.954 -3.84	
ST. JOH MANAGE	INS RIV	ER WATER		SUBSURFACE	TDEM SOUNDING SOUNDING DELTONA	DATA TABLE 18
PALATKA	, FLOR	IDA		DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: TABLE	93742 5.19-1

	No.	TIME (ms)		emf DATA	(nV/m sqrd) SYNTH) DII ETIC (1	FFERENCE percent)
:	8	5 4 9		23 80	21 4	59	2 26
	ğ	6.96		18 24	24.0 10 4	59 · ·	-3.30
	10	8.66		14 09	14 1	50 · · ·	-1.00
	11	11.06		10 72	10 3	24	3 51
	12	14.00		7 79	±0.3	14	J. 54 A 57
	13	17.47		5 39	/•* 5	34	
	14	22.23		3.63	3.6	53	0.00114
	15	28.10		2.51	2.4	12	3 35
					2		5.55
	CUI	RRENT:	21.30 AMP	S EM-5	7 COIL AN	REA: 100.	00 sq m.
	FREQ	UENCY:	3.00 Hz	GAIN:	8 RAMP TI	IME: 192.00	muSEC
	No	M		-			
	NO.	TIME		emf	(nV/m_sqrd)) DI	FFERENCE
		(ms)		DATA	SYNTH	STIC ()	percent)
	16	1 00		220 1	220 /	·	2.00
	17	1 27		32U.1	332.8	5 - -	-3.96
	10 10	1 74		124 2	103.5	כ ו	3.6Z
	10	1.14 1.14		124.2 Q2 77	112.1	L 76	9.69 MASKED
	20	4.1/ 9 77		0J.22 50 72	/ 8 • /	10	J.JD 2 EE
	21	2.77		10 20	5/•. / 5	12 53	3.33
	22	4.37		32 36	40.0	34	-0.02
	23	5.56		25 30	34.3 96 3) <u>-</u> ·	-0.II -0.II
	24	6-98		20.99	20.3	50 · 51	-3.1J 9 97 Wigwer
		0.00		20.33	20	× ±	2.2/ FIADRED
	PARAI "F" P 1 P 2 - T 1	METER RES INDICATES 0.98 -0.04 0. 0.00 -0.0 P 1 1	OLUTION MA FIXED PAN 74 01 1.00 P 2 T 1	ATRIX: RAMETER			
S N F	ST. JOHN MANAGEM PALATKA,	NS RIVER MENT DISTR FLORIDA	WATER RICT	S D	SDII UBSURFACE ETECTION WESTIGATIONS	TDEM SOUN SOL DE PROJECT NO.:	IDING DATA TABLE INDING 18 ELTONA 93742
				I IN	CORPORATED	TABLE	5.19-1
			ويتعاد والمحادث والتكري			L	

5.20 TDEM SITE 19 - BLUE SPRINGS STATE PARK SITE

5.20.1 Location Description and Geoelectrical Section

The site is situated in southwestern Volusia County, Florida (Figure 5.20-1) and is located at Blue Springs State Park within an area of known karst activity. Together, Sites 1, 15, 16, 17, 18, 19 and 20 each lie within the area where the Hawthorn Group is thin to absent. The top of the Floridan aquifer begins at approximately 125 ft below msl (Scott et al., 1991). The top of the Lower Floridan aquifer is approximately 600 ft below msl (Miller, 1986). Water quality data supplied by SJRWMD from surface water collected from Blue Springs on February 3, 1993 showed that chloride concentrations equalled 443 mg/L and sulfate concentrations equaled 67 mg/L.

The resistivity sounding data and best-fit model inversion are presented on Figure 5.20-2. The interpreted geoelectrical section consists of a two-layer subsurface.

5.20.2 Geological Interpretation of Geoelectrical Model

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There is insufficient electrical resistivity contrast between the surficial aquifer system, the Hawthorn Group, and the upper part of the Floridan aquifer to distinguish the three. Fixing the thickness of the upper layer does not resolve this dilemma; therefore, it can be interpreted that there exists a two-layer geoelectrical section with a relatively thick (150.5 m = 494 ft) surface layer of intermediate (27.9 ohm-m) resistivity overlying a low resistivity layer (4.4 ohm-m). It can be interpreted that the surficial aquifer system, the Hawthorn Group and part of the Floridan aquifer exist as a combined but indistinguishable (geoelectrical) layer, overlying a saltwater saturated Floridan aquifer at a depth of 494 ft.

5.20.3 Depth to Occurrence of Salt Water

The bottom (second) layer of the geoelectrical model, with a resistivity of 4.4 ohm-m, is interpreted to represent salt water. It occurs at a depth of 494 ft (-448 ft msl). Because the resistivity of layer 1 (27.8 ohm-m) is less than 80 ohm-m, the interpreted depth to the 5,000 mg/L isochlor is taken at the depth of the geoelectrical interface, or at 494 ft depth (-448 ft msl). The resistivity of layer 2 (4.4 ohm-m) corresponds to a chloride content of 7,157 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2.

5.20.4 Depth of Occurrence of the 250 mg/L Isochlor

Because of the inability to segregate the Floridan aquifer from the overlying surficial aquifer system and the Hawthorn Group, the effective chloride concentration of layer 1 cannot be calculated. Water quality results supplied by SJRWMD indicated that chloride concentrations in water discharging from Blue Springs exceed 250 mg/L. Rutledge (1985) mapped a chloride concentration of above 250 mg/L in the area near the site.

5.20.5 Accuracy of Measurement and Interpretation

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Figure 5.20-3 is the equivalence analysis at this site, and the inversion table (Table 5.20-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about 6 m (20 ft), which is 4% of the total depth. The resistivity of this layer has an equivalence range of from 3.9 - 5.0 ohm-m. This corresponds to a range in interpreted chloride content of from 8,094 mg/L to 6,280 mg/L, again subject to the same assumptions of porosity and validity of equation (4).

The equivalence range of the resistivity of layer 1 is constrained from 26.8 - 29.2 ohm-m. A corresponding chloride concentration cannot be determined because layer 1 is in part comprised of the Hawthorn group and surficial sediments. Accordingly, equation (4) may not be valid.

5.20.6 <u>Summary of TDEM Sounding at Blue Springs State Park</u> (Site 19)

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 494 ft (-448 ft msl). The measured layer resistivity at this level yields a chloride concentration of 7,157 mg/L.
- The quality of groundwater within the Floridan aquifer at the site cannot be interpreted because analysis of the TDEM data does not allow for the surficial aquifer and the Hawthorn Group to be distinguished from the Floridan Aquifer System.

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CLI LOCAT COU PROJ LOOP S COIL SOUNDI	ENT: S ION: E NTY: V ECT: S IZE: LOC: NG COC	JWRMD BLUE SPRI OLUSIA C BALTWATER 229.00 0.00 ORDINATES	NGS STATE OUNTY, FLO INTERFACE 0 m by 0 m (X), : E:	PARK RIDA E DETECTION E 76.000 m 0.000 m (Y) 0.0000 N:	DATE: 11-05 SOUNDING: 1 ELEVATION: 14 QUIPMENT: Geonic AZIMUTH: 0.0000	-93 .00 m cs PROTEM
		FITTI	NG ERROR:	3.306 PER	CENT	
L #	RESIS (oh	TIVITY m-m)	THICKNES: (meters)	S ELEVATI (meter	ON CONDUCTA	ANCE ns)
1 2	27 4	.85	150.5	14.00 -136.5	5.40	
ALL P	ARAMET	ERS ARE	FREE			
PARA	METER	BOUNDS F	ROM EQUIVA	LENCE ANALYSIS	1	
LAYE	R	MINIMUM	BEST	MAXIMUM		
RHO	1 2	26.797 3.928	27.85	8 29.164 6 4.992		
THICK	1	144.390	1.00	0 156.402		
DEPTH	1	144.390	150.51	2 156.402		
CUR FREQU	RENT : ENCY :	24.00 7.50	AMPS EM-: Hz GAIN:	57 COIL ARE 4 RAMP TIM	A: 100.00 sq E: 160.00 muSE(m. C
No.	TI (m	ME IS)	em: DATA	f (nV/m sqrd) SYNTHET	DIFFEREN IC (percer	NCE ht)
1 2 3 4 5 6 7	0. 0. 1. 1. 2.	550 698 869 10 40 75 22	3508.2 2133.2 1360.3 898.5 556.9 381.7 256.8	3615.1 2172.8 1381.0 863.9 565.3 386.3 258.9	-3.04 -1.85 -1.52 3.85 -1.51 -1.18 -0.793	3
ST. JOI	HNS RI	VER WATE	R	SDII	TDEM SOUNDIN SOUNDI	G DATA TABLE NG 19
MANAGE PALATK	EMENT A, FLO	DISTRICT RIDA		SUBSURFACE DETECTION INVESTIGATIONS INCORPORATED	BLUE SPRINGS PROJECT NO.: TABLE	STATE PARK 93742 5.20-1

No.	TIME	emf	(nV/m_sqrd)	DIFFERENCE	
	(ms)	DATA	SYNTHETIC	C (percent)	
8	2.79	174.3	176.8	-1.43	
ğ	3.42	119.1	124.9	_4_84	
10	4.26	81.37	85.52	-5.10	:
11	5.49	52.18	54.35	_4.16	
12	6.96	34.58	35.07	-1.41	
13	8.66	23.99	23.09	3.72	
14	11.06	14.99	14.21	5.20	
CUF FREQU	RENT: 24.00 ENCY: 3.00	AMPS EM-57 Hz GAIN: 8	COIL AREA RAMP TIME	: 100.00 sq m. : 160.00 muSEC	
No	ПТ МФ	5f	(n)//m ~~~~~	NTRABALISA	
NO.	(me)	emi Dama	(IIV/M SQTO)		14
	(DATA	SINTRETI(c (herceur)	
15	0.857	1423.0	1421.4	0.109	
16	1.06	984.4	930.4	5.48	
17	1.37	601.3	590.6	1.76	
18	1.74	405.6	391.5	3.48	
19	2.17	278.1	270.2	2.82	
20	2.77	155.2	180.3	-16.21 MASK	ŒD
"F" 1 P 2 - T 1	0.99 -0.01 0.94 0.00 0.02 0 P 1 P 2	U PARAMETER .99 T 1		·	
ST. JOHN MANAGEN PALATKA,	NS RIVER WATER IENT DISTRICT FLORIDA		SDII UBSURFACE ETECTION WESTIGATIONS ICORPORATED	TDEM SOUNDING DATA SOUNDING 19 BLUE SPRINGS STATE PROJECT NO.: 93 TABLE 5.	TABLE PARK 3742 20-1

5.21 TDEM SITE 20 - DE LAND SITE

5.21.1 Location Description and Geoelectrical Section

The site is situated in southwestern Volusia County, Florida (Figure 5.21-1) and is located within a woodland with no obvious sources of interference. Together, Sites 1, 15, 16, 17, 18, 19 and 20 each lie within the area where the Hawthorn Group is thin to absent. The top of the Floridan aquifer begins at approximately 65 ft below msl (Scott et al., 1991). The top of the Lower Floridan aquifer is approximately 750 ft below msl (Miller, 1986).

The resistivity sounding data and best-fit model inversion are presented on Figure 5.21-2. The interpreted geoelectrical section consists of a two-layer subsurface.

5.21.2 Geological Interpretation of Geoelectrical Model

There is insufficient electrical resistivity contrast between the surficial aquifer system, the Hawthorn Group and the underlying Floridan aquifer to distinguish the three. Fixing the thickness of the upper layer does not resolve this dilemma; therefore, it can be interpreted that there exists a two-layer geoelectrical section with a relatively thick (276.7 m = 908 ft) surface layer of intermediate resistivity (42.3 ohm-m) overlying a low resistivity layer (3.3 ohm-m). It can be interpreted that the surficial aquifer system, the Hawthorn Group, and part of the Floridan aquifer exist as a combined but indistinguishable (geoelectrical) layer, overlying a saltwater saturated Florida aquifer at a depth of 908 ft.

5.21.3 Depth to Occurrence of Salt Water

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The bottom (second) layer of the geoelectrical model, with a resistivity of 3.3 ohm-m, is interpreted to represent salt water. It occurs at a depth of 908 ft (-869 ft msl). Because the resistivity of layer 1 (42.3 ohm-m) is less than 80 ohm-m, the interpreted depth to the 5,000 mg/L isochlor is taken at the depth of the geoelectrical interface, or at 908 ft depth (-869 ft msl). The resistivity of layer 2 (3.3 ohm-m) corresponds to a chloride content of 9,594 mg/L assuming a porosity of 25% and the validity and applicability of equation (4) of Section 4.2.

5.21.4 Depth of Occurrence of the 250 mg/L Isochlor

Because of the inability to segregate the Floridan aquifer from the overlying surficial aquifer system and the Hawthorn Group, the effective chloride concentration of layer 1 cannot be calculated. According to Rutledge (1982), the chloride concentration in the Upper Floridan aquifer in the area of this site is less than 250 mg/L.

5.21.5 Accuracy of Measurement and Interpretation

Figure 5.21-3 is the equivalence analysis at this site, and the inversion table (Table 5.21-1) lists the upper and lower bounds of the inverted parameters of the geoelectrical model.

The range of equivalence in determining the depth to the low resistivity layer is about $\pm 7 \text{ m}$ (23 ft), which is 3% of the total depth. The resistivity of this layer has an equivalence range of from 2.8 - 4.0 ohm-m. This corresponds to a range in interpreted chloride content of from 11,334 mg/L to 7,888 mg/L, again subject to the same assumptions of porosity and validity of equation (4). The estimated depth of the saltwater layer (908 ft bls) does not concur with the estimated depth of the freshwater/saltwater interface (approximately 600 ft bls) based on Ghyben-Herzberg analysis (Rutledge, 1982). Rutledge (1982) defines the depth to the freshwater/saltwater interface as forty times the groundwater table elevation above msl. The freshwater /saltwater interface is assumed to occur at the transition to unmixed seawater with a chloride concentration of 19,000 mg/L.

The equivalence range of the resistivity of layer 1 is from 41 - 44 ohm-m. A corresponding chloride concentration cannot be determined because layer 1 is in part comprised of the Hawthorn Group and surficial sediments. Accordingly, equation (4) may not be valid.

5.21.6 Summary of TDEM Sounding at De Land (Site 20)

- The depth of occurrence of salt water (5,000 mg/L isochlor) is interpreted to be 908 ft (-869 ft msl) and occur within the Lower Floridan aquifer. The measured layer resistivity at this level yields a chloride concentration of 9,594 mg/L. The results of the TDEM study do not agree with the results from other studies in the area. Rutledge (1982) estimated a depth of approximately 600 ft bls to the freshwater/saltwater interface.
- The quality of groundwater within the Floridan aquifer at the site cannot be interpreted because analysis of the TDEM data does not allow for the surficial aquifer and the Hawthorn Group to be distinguished from the Floridan Aquifer System.






CLI LOCAT COU PROJ LOOP S COIL SOUNDI	ENT: 9 ION: 1 NTY: V ECT: 9 IZE: LOC: NG COO	JWRMD DELAND VOLUSIA CO SALTWATER 305.000 0.000 DRDINATES:	OUNTY, FL INTERFAC m by m (X), E:	ORID E DE 34	A TECTION 5.000 m 0.000 m 0.0000 N	SOU ELE EQU A: (Y)	DATE: 1 UNDING: 1 VATION: IPMENT: (ZIMUTH: 0.00(12-05-93 1 12.00 m Geonics PF 00	N ROTEM			
		FITTIN	IG ERROR:		4.656	PERCEI	NT					
L #	RESIS (oł	STIVITY 1m-m)	THICKNE (meters	SS)	ELEV (me	/ATION sters)	CO1 (5	NDUCTANCE Siemens)				
1 2	42	2.33 3.30	276.7		12 -264	.00 .7		6.53				
ALL P	ALL PARAMETERS ARE FREE											
PARA	METER	BOUNDS FR	OM EQUIV	ALEN	CE ANALY	SIS						
LAYE	R	MINIMUM	BES	т	MAXIMU	IM						
RHO	1 2	40.901 2.759	42.3 3.3	36 09	43.84 4.02	3 13						
THICK	1	269.115	1.0	00	283.29	5						
DEPTH	1	269.115	276.7	43	283.29	5						
CUR FREQU	RENT: ENCY:	19.50 A 7.50 H	MPS EM Z GAIN	-57 : 4	COIL RAMP	AREA: TIME:	100.0 257.00)0 sq m. muSEC				
No.	TI (I	IME as)	e DAT	mf (: A	nV/m sqr SYNI	d) HETIC	DIF (I	FFERENCE percent)				
1 2 3 4 5 6 7	0 0 1 1 2	550 698 869 10 40 75 22	7275.6 4295.6 2622.1 1600.7 935.7 579.8 356.6		7531 4548 2772 1567 901 551 339	.1 .8 .2 .5 .3 .3 .0	-	-3.51 -5.89 -5.72 2.07 3.67 4.90 4.93				
ST. JOHN		ER WATER		S	DII	-	TDEM SOU	NDING DATA UNDING 20 DELAND	TABLE			
MANAGEN PALATKA,	FLOR	DA		DET	ECTION STIGATION	s	PROJECT NO.	: 9:	3742			
				INC	JKPUKAIE	ע		J.,	- 1 - 1			

. A

No. T. (1	IME ms) Di	emf (nV/m sqr ATA SYNT	d) DIF HETIC (P	FERENCE ercent)
$\begin{array}{cccccccccccccccccccccccccccccccccccc$.79229.42161.26115.4979.9656.6640.0628.0019.4712.238.104	.0 224 .3 161 .0 115 .77 79 .38 55 .78 39 .96 26 .06 18 .87 12 .87 4	.5 .4 – .2 – .33 .71 .83 .99 .11 .24 .75 .85	1.95 0.0922 0.124 0.548 1.18 2.32 6.81 4.97 4.84 5.67 0.386
CURRENT: FREQUENCY:	19.50 AMPS H 3.00 Hz GA	EM-57 COIL IN: 8 RAMP	AREA: 100.0 TIME: 257.00	0 sq m. muSEC
No. T: (1	IME ns) D2	emf (nV/m sqr ATA SYNT	d) DIF HETIC (P	FERENCE ercent)
19 0 20 1 21 1 23 2 24 2 25 3 26 4 27 5 28 6 29 8 30 10 PARAMETER 1 "F" INDICA P 1 1.00 P 2 -0.01 T 1 0.00 P 1	.857 2695 .06 1708 .37 985 .74 590 .17 368 .77 239 .50 155 .37 109 .56 76 .98 55 .56 39 .64 28 RESOLUTION MATRIX FES FIXED PARAMET 0.89 0.01 1.00 P 2 T 1	.3 2867 .4 1722 .4 956 .0 562 .1 358 .1 231 .8 159 .5 114 .78 81 .85 58 .96 43 .71 31	1 - 4 - 3 4 5 3 6 - 5 - 35 - 55 - 46 -	6.37 0.819 2.95 4.67 2.61 3.26 2.38 4.56 5.94 5.09 8.98 9.58
ST. JOHNS RIVE MANAGEMENT DI	R WATER STRICT	SUBSURFACE	TDEM SOUND SOUN DE	DING DATA TABLE IDING 20 LAND
PALATKA, FLORIE	A	DETECTION INVESTIGATIONS INCORPORATED	PROJECT NO.: TABLE	93742 5.211

6.0 SUMMARY AND CONCLUSIONS

A time domain electromagnetic (TDEM) survey was performed at 19 sites in the St. Johns River Water Management District and one site in Georgia during the months of April-May, 1993. The principal findings of this survey can be summarized as follows:

TDEM is a geoelectrical method which can be used to estimate the vertical variation or resistivity of subsurface formations and/or hydrostratigraphic units. Translating the geophysical measurement of electrical resistivity into a model of geology and water quality depends upon comparison to other available subsurface data, consistency of data sets from nearby soundings from this and prior years, and application of empirical relationships to produce interpreted water-quality results. As outlined in Section 4, the conversions to water quality values (chloride concentrations) are based upon the relationships established using Kwader's (1982) data for Seminole County, as used for SJRWMD in previous studies (Blackhawk, 1990 and CEES, 1992). The formulae employed use assumptions of a 25% porosity, similar water chemistry (specifically, a 5:1 chloride-to-sulfate ratio) as Kwader's data, and that the saltwater interface occurs within the Floridan Aquifer System. With regards the latter point, chloride concentration values are presented only for those portions of the geoelectrical section which correspond to the Floridan aquifer and not for the Hawthorn Group and surficial sediments or for surficial sediments where the Hawthorn Group is missing.

Under circumstances where there is little contrast in resistivity between the surficial sediments and/or Hawthorn Group and the Upper Floridan aquifer, the chloride concentration of the ground water above the freshwater/saltwater interface cannot be determined. This is because of the assumptions implicit in equation (4) are not valid.

Finally, because the freshwater/saltwater boundary is not an abrupt interface but a transition zone, criteria relating to the relative resistivities above and below the geoelectrical interface were used to establish an empirical definition of depths to the 250 and 5,000 mg/L isochlors. Again, these were the same criteria as used in past years' TDEM surveys (Blackhawk, 1990 and CEES, 1992) in order to maintain consistency from year to year.

6-1

6.1 DETERMINING THE DEPTH OF THE INTERFACE BETWEEN FRESH WATER AND GROUND WATER OF HIGH CHLORIDE CONCENTRATION (GREATER THAN 1,450 mg/L)

As stated in previous years' reports (Blackhawk, 1990 and CEES, 1992), "ground water with a chloride content greater than 1,450 mg/L is characterized in the Floridan aquifer by resistivities less than 20 ohm-m when the aquifer has a porosity of about 25%." In accordance with this statement, a deep layer with a resistivity of less than 20 ohm-m was detected at 18 of the 20 sites surveyed. At the Picolata sounding, Site 6, the basal resistivity layer had a resistivity of 27.4 ohm-m. The resistivity value was sufficiently close to 20.0 ohm-m that the results of this sounding were considered similar to the other 18 sites. Only the Cumberland Island sounding, Site 2, failed to detect the basal, low resistivity layer. A forward-modeling/sensitivity analysis indicates that if such a layer were present, it must be at least 2,200 ft deep. The remaining 19 sites show variation in depth to this interface to range from approximately 413 - 2,427 ft. All the interpreted depths place the saltwater interface within the Floridan aquifer System.

6.2 WATER QUALITY IN THE FLORIDAN AQUIFER AND DEPTH OF OCCURRENCE OF THE 250 mg/L ISOCHLOR

Based on the assumptions that: (a) The Floridan aquifer has a porosity of 25%, (b) ground water within the study area have a chemistry similar to those analyzed by Kwader (1982), and (c) equation (4) in Section 4.2 is valid, ground water having chloride concentrations of less than 250 mg/L correspond to geoelectrical layers having resistivities in excess of 80 ohm-m. The distribution of resistivities of the Upper Floridan aquifer show, for the most part, high resistivities and, therefore, fresh waters of less than 250 mg/L are present in the Floridan aquifer over much of the survey area. There was one site, St. Augustine #1 (Site 4), where the resistivity of the Upper Floridan aquifer water is interpreted to be present. When a layer with a chloride concentration of less than 250 mg/L is interpreted, the position of the 250 mg/L isochlor is fixed by the relative resistivities of the deep, conductive layer and the fresh (resistive) layer above - generally placing it 50 ft above the geoelectrical interface. When the resistivity of the Upper Floridan aquifer is such that the interpreted chloride concentration exceeds 250 mg/L, a depth to the 250 mg/L isochlor was not determined as the entire system is considered to be brackish.

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